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BUREAU OF ENGINEERING

Report  
on  
Absorption Coefficients of Supersonic  
Sound in Open Sea Water.

NAVAL RESEARCH LABORATORY  
ANACOSTIA STATION  
WASHINGTON, D.C.

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## INTRODUCTION

1. This expedition was in part a continuation of a program of surveying strategic areas to study the water conditions affecting underwater sound transmission and reception. Previous expeditions had covered the North and Central Atlantic and Guantanamo-Cuba areas. This expedition covered the areas adjacent to the Panama Canal. Three areas were selected on the Atlantic side and three on the Pacific side, all beyond the 100 fathom curves, and so distributed as to divide the approaches into three approximately equal areas. Each area was occupied for a period of 24 hours and an average of nine runs was made in each period.

2. The theoretical basis for this work and the experimental routine are given in detail in previous reports on absorption coefficients (Reference 1). Certain new items in technique may be noted here.

3. An automatic recorder was used on the S-20 which gave a continuous record of the intensity of the signal transmitted by the SEMMES as she steamed towards the S-20 from a maximum range of about 14,000 yards. The zero decibel level on the recorder was maintained constant on all runs at 0.1 microvolts delivered by the D-2 type pick-up on the S-20 to the input of the amplifier. By the use of an attenuation box in the input to the amplifier, and by taking the recorder lead from the amplifier detector grid, the signal strength was always kept within the linear range of the amplifier and recorder. The zero range level was also measured in February and March, but is not used in calculating  $\alpha$ , which is based on the slope of the range-intensity curve.

4. A platinum resistance thermometer with 250 feet of cable was used to measure the temperatures. The bridge was calibrated to read directly in degrees Fahrenheit to 0.01 degree. At the conclusion of each run temperatures were measured at 0, 0.5, 1, 2, 3, 4, 5, 6, 7, 8, 9, 10, 12, 14, 16, 18, 20, 25, 30 and 35 fathoms. There is some uncertainty about the depths below 20 fathoms due to the drift of the ship while lying to.

5. An attempt was made to use a newly developed bathythermograph to record automatically both pressure and temperature, but it did not prove satisfactory under the working conditions.

6. In calculating the ship's speed during the approach runs in a given area, all the RAR data for the same r.p.m. were averaged and used on all runs at that speed, instead of using the RAR values for each run singly. This materially reduced the labor of plotting the curves and improved the average accuracy of the absorption coefficients. The velocity of sound for the RAR calculations was taken as that for the average surface temperature during the period (Reference 2).

Ref. 1 NRL Report No. S-1204 of 16 October 1935.

NRL Report No. S-1466 of 12 August 1938.

Ref. 2 Underwater Sound in Northern Waters, NRL file C-S68/86.

## RESULTS

7. The original data are plotted on Plates 1 to 58 as range-intensity curves and temperature-depth curves with notations as to the number of the run, the date, time, sea, sky and wind conditions and the calculated values of the loss coefficient,  $\alpha$ , for each frequency. Each point on a range-intensity curve represents the average value of the sound intensity, as read from the recorder tape, for a period of one minute.

8. The loss coefficient,  $\alpha$ , is taken as the slope in decibels per kiloyard of the "best" straight line between 1000 and 10,000 yards on the Atlantic side. The drawing of this particular line is somewhat arbitrary and involves considerable judgment in certain cases. An attempt was made to be consistent for comparative purposes and the 1000 to 10,000 yard interval gave the most consistent results.

9. On Plates 29 to 30 the temperature gradients at different depths and  $\alpha$  are plotted against clock time to show the relation between gradient and  $\alpha$ .

## DISCUSSION OF RESULTS ON THE ATLANTIC SIDE

10. The data for the Atlantic side are generally comparable with those obtained in the Guantanamo area in 1938. It is usually possible to get an average value of  $\alpha$  from 1 to 10 kiloyards, though nearly all curves show considerable fluctuations from the mean value. A summary of the loss coefficient data is given in Table 1.

TABLE 1

SUMMARY OF LOSS COEFFICIENT DATA  
ATLANTIC SIDE

<u>Area</u>	<u>Date</u>	<u>Median Time</u>	<u>Loss Coefficient</u> <u>OC = db/kyd.</u>		<u>Ratio</u> <u>OC 23.6/</u> <u>/OC 17.7</u>	
			<u>17.7 Kcs.</u>	<u>23.6 Kcs.</u>		
B	30 Jan.	1617	3.5			
		1819	4.6			
		2026	4.4			
		2224	6.2	7.8	1.26	
	31 Jan.	0026	7.6	9.1	1.20	
		0245	6.8	7.2	1.06	
		0613	5.0	5.1	1.02	
		0818	4.0			
		1020	5.0			
	Area	Average -		5.23	7.30	
	C	31 Jan.	1238	3.5	4.2	1.20
			1611	5.7		
			1845	4.0	5.4	1.35
			2247	3.5	6.2	1.77
1 Feb.		0214	5.6	8.4	1.50	
		0534	4.5	6.0	1.33	
		0810	5.6	6.7	1.20	
		1016	5.5	7.4	1.35	
Area		Average -		4.74	6.33	
A		2 Feb.	1041	7.6		
	1324		6.0			
	1542		5.5	6.1	1.11	
	1907		5.7	6.9	1.21	
	2139		6.8	8.0	1.18	
	3 Feb.	0012	5.3			
		0210	5.8			
		0406	6.6			
		0541	5.1			
		0920	<u>5.3</u>			
	Average		5.97	7.00	1.27	
	Grand Average		5.33	6.74	1.27	
			±0.60	±0.50		
Guantanamo Feb.-Mar. 1938		4.06	6.00	1.48		
Ratio Panama to Guantanamo		1.31	1.17	.85		

11. For comparison the results at Guantanamo in February - March 1938 and the ratios of the coefficients for 23.6 kilocycles to 17.6 kilocycles in the two areas are also given. The higher average loss coefficient in Panama may be explained as due to much larger temperature gradients and to larger swells and white-caps from the sweep of the wind across the open Caribbean. It is also probable that the water is more disturbed by surface and deep water currents in the Panama area. While the loss coefficient for 23.6 kilocycles is always greater than for 17.6 kilocycles on the same range-intensity run, the ratio is less at Panama than at Guantanamo. The difference would appear to be greater than the experimental error, but there is no obvious explanation.

12. A notable characteristic of nearly all the curves is the wide fluctuation in intensity from the mean value. The fluctuations are a complex combination of short time and long time effects. For example, a common case like Run 17 gives the following values of  $\alpha$  for successive 1000 yard intervals.

TABLE 2

<u>Kiloyard Interval</u>	<u><math>\alpha</math> = db/kyd.</u>
1-2	+ 8
2-3	+10
3-4	-23
4-5	+32
5-6	-14
6-7	+29
7-8	- 2
8-9	-10
9-10	+11
1-10 mean	+ 5.6

where negative  $\alpha$  means an increase in intensity with range. In an extreme case like Run #27 short range values for  $\alpha$  may run as high as  $\pm 100$  db/kyd. On the original tape short time swings of 15 db/sec. are common corresponding to values for  $\alpha$  of  $\pm 5$  db/yd. or 1000 times the average value.

13. These rapid fluctuations lead to the conclusion that the intensity at any point is the resultant of a combination of two or more rays or paths, a direct ray, a reflected ray, and a refracted ray, at times in phase producing reinforcement and at times more or less out of phase producing interference, but at all times suffering loss through

absorption, scattering, reflection, or refraction. These effects may be different for the different paths. This opinion is confirmed by a detailed study of the continuous record. For short periods, either the high values of intensity or the low values may be steady with only a few decibels fluctuation, but the intermediate values fluctuate rapidly and irregularly from second to second by as much as  $\pm 20$  decibels. This is evident on the short sections of tape shown on Plate 59. Since the usual ship's speed was 3 to 4 yards per second and the wavelength of underwater sound at 17.6 kilocycles is 3.3 inches, it is obvious that rapid fluctuations are possible if there are two or more waves traveling by different paths. Two paths are always present - the direct path and the path of the ray reflected from the air-water surface. Due to the roughness of this surface from ripples, swells or white-caps, this sound is generally diffused or scattered, but for short periods of seconds or less may give specular reflection. There is also a  $180^\circ$  phase displacement for this type of reflection.

14. A third path may also be present if there is a temperature inversion as was common on the Atlantic side, e.g., runs 17 and 27. Since sound travels faster in warm water than in cold, the lower half of the beam which tends to spread downward even with no gradient will be refracted upward by this warm layer to reinforce or interfere with the direct signal and with the signal reflected from the air-water surface. Where there are large inversions as in runs 11, 17 and 27, the range-intensity curves are generally characterized by relatively high levels, maintained to long ranges and also by relatively large and roughly periodic swings in intensity. The wind velocity is relatively low - 10 kts. or less for these runs. Also the peaks and valleys of the 24.6 kilocycle curve follow quite closely those of the 17.6 kilocycle curve. In run 21 with less marked inversion, the two frequencies are quite different. For runs 24 to 28 inclusive, readings were taken on 17.7 kilocycles only to develop the curve more in detail from the continuous record, but there was no significant change in form.

15. An interesting point to be noted on the temperature versus depth curves is that from 1700 on 31 January to 1050 on 1 February the maximum temperature of the inversion point between 10 and 15 fathoms was always higher than the surface temperature at the given time, and that on run 26 the peak inversion temperature was higher than any temperature measured in the period 31 January - 3 February. Apparently the wind did not reach sufficient velocity effectively to stir the water down to the 15 fathom level. The maximum wind recorded was 14 - 15 kts. and it was generally less than 10 kts. For this warm layer between 10 and 15 fathoms to persist for nearly a week between two colder layers seems to require a compensating change in salinity, but no data are available on this point.

16. The general effect of a temperature inversion of  $0.2^\circ\text{F}$  or more is to keep the major portion of the sound energy in the upper layer of water above the inversion point. The instantaneous values may fluctuate rapidly due to reinforcement or interference, but the general level is high. The average intensity at 10,000 yards is 40 decibels for all runs at 17.7 kilocycles and in several cases it is up to 70 decibels.

17. The subject of reflection coefficients is being covered in a separate report, but it may be stated here that during the period 3-7 February echoes were consistently obtained at 3000 yards and occasionally at 4000 yards or more when the submarine was on the surface or at periscope depth. At 60 feet the maximum range was approximately 1500 yards and at 100 feet no echoes were obtained at all. The temperature depth curves on runs 27 and 28 show large temperature inversions on 3 February. The temperature-depth data taken on 6 and 7 February are given on Plate 31 and show large gradients at the 10 - 15 fathom level with an inversion near 5 fathoms when the best echoes were obtained on 6 February. These data therefore confirm the explanation given in the preceding paragraph and indicate the tactical importance of this type of information.

#### DISCUSSION OF RESULTS ON THE PACIFIC SIDE

18. Runs 29 to 54 were made in three areas on the Pacific side beyond the 100 fathom curve approximately on a line between Capa Mala and Pinas Bay during the period 28 February to 9 March 1939. The range-intensity curves are generally characterized by an initially low level, a rapid drop in the first 2000 or 3000 yards, and a very slow decrease beyond this point. In general separate values of  $\alpha$  are shown for 1000 to 3000 yards and for 3000 to 10,000 yards. A summary of the loss coefficients  $\alpha$  and also the decibel level at 1000 and 10,000 yards is given in Table III.

TABLE 3

#### SUMMARY OF AVERAGE, PACIFIC SIDE

Area	Date	Loss Coefficients				Db Level for		Wind Velocity in Knots
		1000 to 3000 yds.		3000 to 7000 yds.		17.6 kcs.		
		17.7 kcs.	23.6 kcs.	17.6 kcs.	23.6 kcs.	1000 Yards	10,000	
A	2-3 Mar.	13.6	16.0	6.1	6.3	93	21.6	12
B	28 Feb.- 1 Mar.	20.0	24.0	2.0	2.9	80	24.0	17
C	8-9 Mar.	24.7	26.6	3.1	3.6	90	22.0	19
Average		19.4	22.2	3.7	4.3	88	22.5	16

19. The temperatures of the waters of Panama Bay are determined in a complicated way by the combination of solar radiation, ocean currents, wind, tides, bottom contours, salinities, animal or vegetable growth and doubtless other factors. The cold Humbolt current running north or north-west along the Pacific coast of south and central America maintained a fairly constant temperature of approximately 60°F. at 20 fathoms depth during the period 28 February - 9 March. The surface temperature followed a daily cycle from a minimum of 66°F. at about 0200 to a maximum of 72°F.

at about 1500 averaging about 68°F. for the period. This is approximately 10°F. colder than the average surface temperature measured on the Atlantic side during the period 28 January - 3 February. The outstanding feature, however, is the high temperature gradients frequently reaching values as high as 1.2°F. per fathom around the 5 fathom curve. It should be noted that the temperature scale on the temperature-depth curves for the Pacific side is five times the scale used on the Atlantic side. Occasionally a temperature inversion was found on one run, but it did not persist until the next one.

20. In spite of the obvious effect of the temperature gradients on sound transmission, there is no consistent simple correlation between the gradients and the loss coefficients as was found on the Atlantic side.

21. The wind velocity was generally high averaging 16 knots and reaching a maximum of 24 knots. The direction was consistently within 15 degrees of North. It was noted, however, that the swells were not so high as on the Atlantic side and that white-caps appeared generally at about 14 knots as compared with 10 knots on the Atlantic side. During runs 29 to 36 while the S-20 was lying to she drifted some 30 miles in 20 hours or at a rate of 1.5 knots in an average wind of 17 knots.

22. It is a matter of common knowledge in Coco Solo that submarines making deep dives in Panama Bay may have to take on 5000 to 10,000 pounds of ballast after they have a good trim at periscope depth before they can submerge to greater depths. This is contrary to the general rule or to experience on the Atlantic side at Coco Solo. Attempts to compute the effects of temperature and pressure on the water and the submarine do not give values corresponding to the experience stated above, so it seems reasonable to assume that the deeper, colder water of the Humbolt current has a higher salinity than the shallow warm layer near the surface. The effect would be enhanced by rainfall, or run off from the adjacent land, and the thermal gradient would delay the mixing. Also the stirring action of the north wind at this season was not as effective as in the open Caribbean, as may be seen by a comparison of the temperature-depth curves in the two areas.

23. The high initial value of  $\alpha$ , averaging 19.4 decibels per kiloyard for 17.6 kilocycles for the 1000 to 3000 yard interval of the range-intensity curve, can be explained as due to the large temperature gradient at the boundary between the shallow warm layer and the colder water below, which effectively drains the sound energy out of the upper layer. The low value of  $\alpha$  averaging 1.5 decibels per kiloyard for 17.6 kilocycles on runs 31 to 35 and 46 to 49 for the 3000 to 10,000 yard range, may be explained on the assumption of relatively high salinity in the deeper, colder water. Since the velocity of sound increases with salinity and depth (reference 2) a certain portion of the energy refracted downward from the upper layer will be refracted upward from the more dense region to reach the receiver on the surface ship at longer ranges. While the intensity level is relatively low at 20 to 30 decibels, it is also relatively uniform, indicating in general a single path rather than a combination of two or more as found on the Atlantic side with a temperature inversion.

24. In addition to the vertical temperature gradients there was considerable evidence of horizontal gradients although these were not specifically measured. The injection water temperature on the SEMMES frequently varied in short periods of time. Readings of the thermometer on raising seldom checked the temperatures found on lowering as the ship was drifting all the time. The S-20 reported great difficulty at times in maintaining an even keel when running submerged, although normally her trim was good. This effect is believed to be due to the strong cold current running over a ridge and being given an upthrust that carried it into the warmer layer. There would be both a thermal and a current effect on the submarine and a horizontal temperature gradient would be found.

25. Certain areas in Panama Bay are conspicuous for the "red" water. These areas are patches of varying sizes up to several square miles. Several range runs were made through such areas, but the loss coefficients were so high and variable throughout the day that it was impossible to determine the effect of the red water itself. It seems probable that the same conditions that produce poor sound conditions are also favorable for the growth of these microscopic plants and that they are an indication rather than a cause of poor transmission.

26. The two temperature-depth curves on Plate 58 were taken during experimental torpedo firing by the S-20 as points about 20 miles west of the Perlas Islands. Listening conditions were fair on 24 February and good on 25 February. On the latter day the torpedoes could be picked up where fired by the S-20 or approximately 4000 yards and followed to the SEMMES which was the target. Since the water was relatively shallow the long range sound may have been reflected from the bottom as the temperature-depth curves would indicate poor direct transmission in deep water. No echoes were obtained at 3000 to 4000 yards from the S-20 submerged at periscope depth.

#### SUMMARY OF EXPERIMENTAL DATA ON THE TRANSMISSION OF SOUND THROUGH WATER.

27. The mass of data collected during the past six years with continuously improving measuring equipment and technique and covering areas with widely different characteristics warrants certain statements concerning the transmission of sound through open sea water that are summarized here for convenience.

- (a) Sea water cannot in general be considered the homogeneous medium which the simple laws of wave transmission assume.
- (b) Temperatures and temperature gradients vary widely in amount and with time. The time variations are seasonal, diurnal, or irregular due to winds and currents.
- (c) Salinities are known to vary with location, depth and time, but little quantitative salinity data have been taken to correlate simultaneously with the sound data.

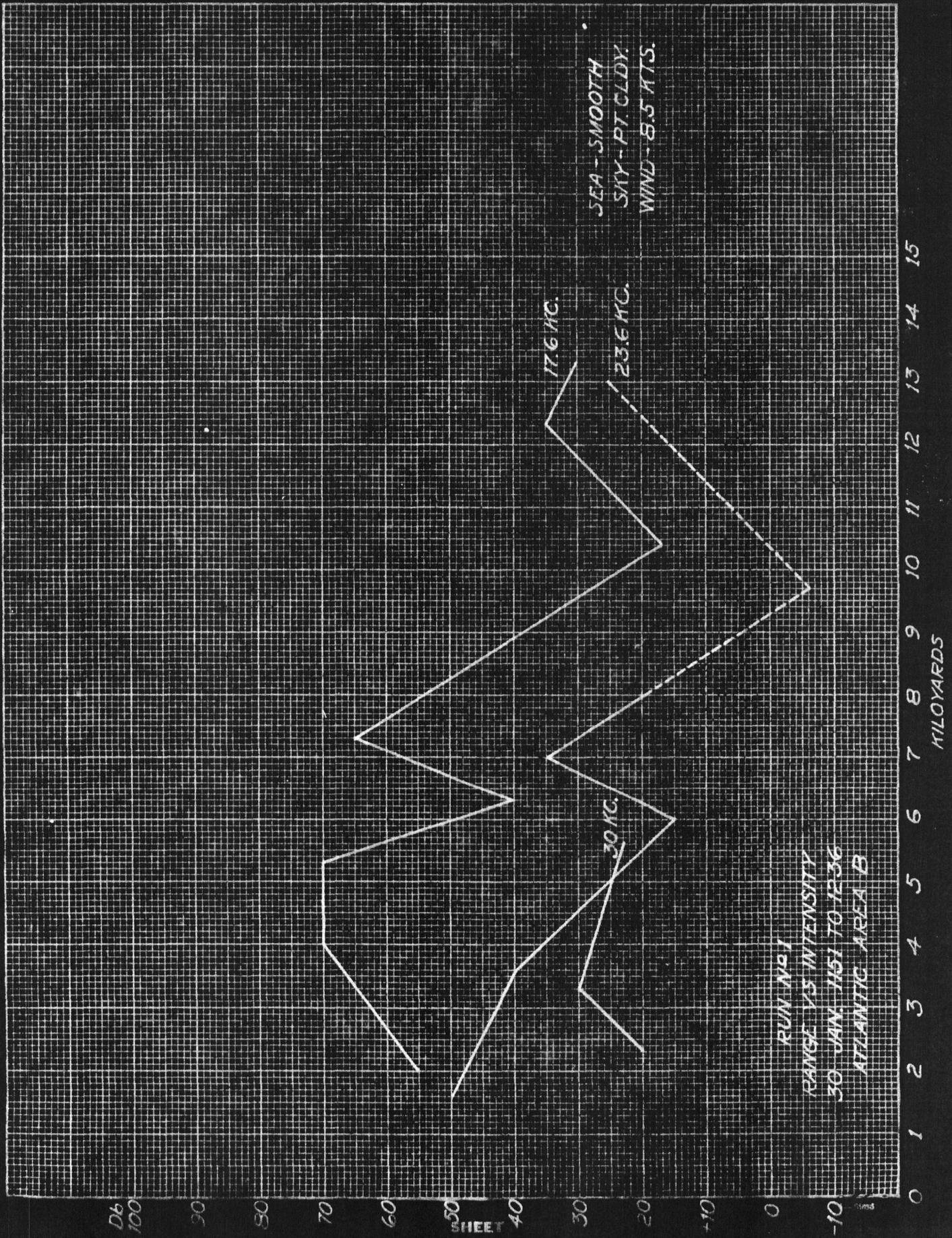
- (d) The wind is an important, though irregular, factor because of its stirring action on the water which may extend to a depth of 200 feet or more and profoundly modify the temperature gradients.
- (e) Animal or vegetable growth, both microscopic and macroscopic, may affect sound transmission, but the effect has not been measured positively or quantitatively.
- (f) Turbulence is common and may be fine grained or coarse grained. It may be caused by the wind or by ocean currents, particularly where there is a relative movement between the surface layer and the deeper water.
- (g) The sound beam is subject to absorption, reflection, refraction and scattering, operating according to the common laws of physics, but combining in complicated ways to determine the instantaneous value of the intensity at a given point.
- (h) The numerical values for the pure absorption coefficient of homogeneous sea water which may be calculated from its physical constants according to classical theory are 100 to 1000 times smaller than the loss coefficients determined experimentally. The other items given in (e) to (g) above are, therefore, important causes of loss.
- (i) Range-intensity curves, where decibels (log intensity) are plotted against range between 500 and 10,000 yards or more, may be divided into three broad or general types for which equally broad or general explanations may be given.
- (j) The first type shows a linear relation between decibels and range with a generally high level and with only small variations. It is obtained where the temperature gradient is small or zero down to 100 feet or more and the surface has a light chop or ripples. The temperature gradient causes little if any downward refraction of the beam and the ripples scatter the sound reflected from the air-water surface so that it does not interfere with the direct signal. Run 14, Plate 14, is a typical example.
- (k) The second type may show an average linear relation between decibels and range, but the instantaneous value may vary widely from the average in a somewhat periodic way. The average level is not quite as high as for the first type, but the peaks may reach surprisingly high values at long ranges. This type of curve is usually found where there is a small <sup>positive</sup> temperature gradient (decrease of temperature with depth) near the surface and a <sup>negative</sup> temperature gradient, or inversion, at lower depths. The surface gradient bends the beam downward, but the inversion bends it upward again and since there may be many paths of different length, there are strong reinforcements and interferences as indicated by the fluctuations in the curve. Run 11, Plate 11, is a typical example.

- (l) The third type of curve shows a very rapid drop at 5 to 100 times the normal rate to a low value in the first few thousand yards and usually a subnormally low decrease thereafter. The first effect is due to a large <sup>negative</sup> positive temperature gradient in the shallow surface layer that effectively drains the sound energy out of this layer, or bends the beam so that it passes under a receiving ship at shallow depths. If there is a considerable increase in salinity or density with depth the beam may be refracted upward and a low level signal will be maintained at the longer ranges. Run 32, Plate 35, is a typical example of this type.
- (m) While three general types of range-intensity curves may be distinguished there are many graduations or combinations between them and no exact or consistent correlation of any type with season and place is possible unless additional data on sea, sky, wind, temperatures and salinities are also known and considered.
- (n) Water conditions are and will be the dominant factors in determining the ranges attained by sound equipment. Increased sound output or increased receiving sensitivity give a proportional gain in range under ideal conditions, but under unfavorable conditions, such as the extreme temperature gradients found on the Pacific side in Panama, they are ineffective because the beam does not strike the target.
- (o) For the installations on the SEMMES and S-20 during 1938 at Guantanamo and 1939 at Panama, broadside echoes were generally obtained by the SEMMES at 10 knots or less on the 17.6 kilocycle equipment from the S-20 on the surface or at periscope depth if the direct signal intensity at the S-20 was  $60 \pm 5$  db. or more. Thus an estimate of the echo ranges may be obtained from the range-intensity curves. This experimental value of  $60 \pm 5$  db. applies accurately only to these two ships under the conditions stated, but should be approximated by any QC equipment in good condition and under the same conditions. If, however, the speed of the searching ship is increased so that the background of noise is raised 10 decibels, then a 70 decibel level is necessary to hear echoes. If the submarine is below a temperature inversion the echo range is extremely short or zero. The target angle is also a big factor.

## CONCLUSIONS

28. From the data presented above, it is concluded:

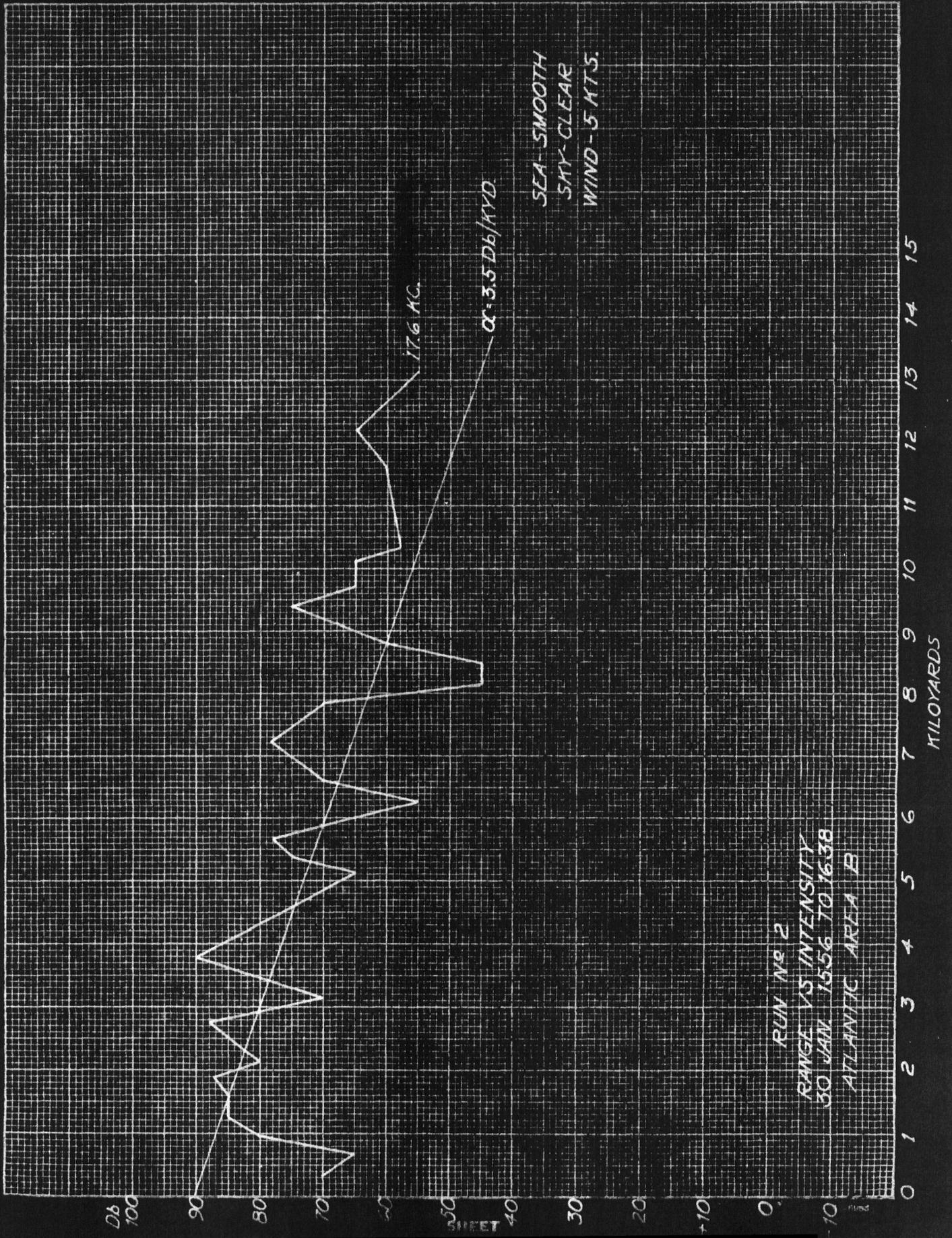
- (a) Water conditions are a major factor in the transmission of sound through water.
- (b) Water conditions vary widely and rapidly with time and place and accurate classifications are not possible.
- (c) Experimental determination at the particular time and place by sound methods is the most practical means of obtaining sufficiently accurate data for planning a sound tactical program.
- (d) In the absence of exact data at the particular time and place, a rough estimate of conditions to be anticipated may be obtained from this and previous reports.



RUN No 1  
RANGE 1/3 INTENSITY  
30 JAN 1151 TO 1236  
ATLANTIC AREA B

SHEET

4-0700 Rev. 15, Apr 51

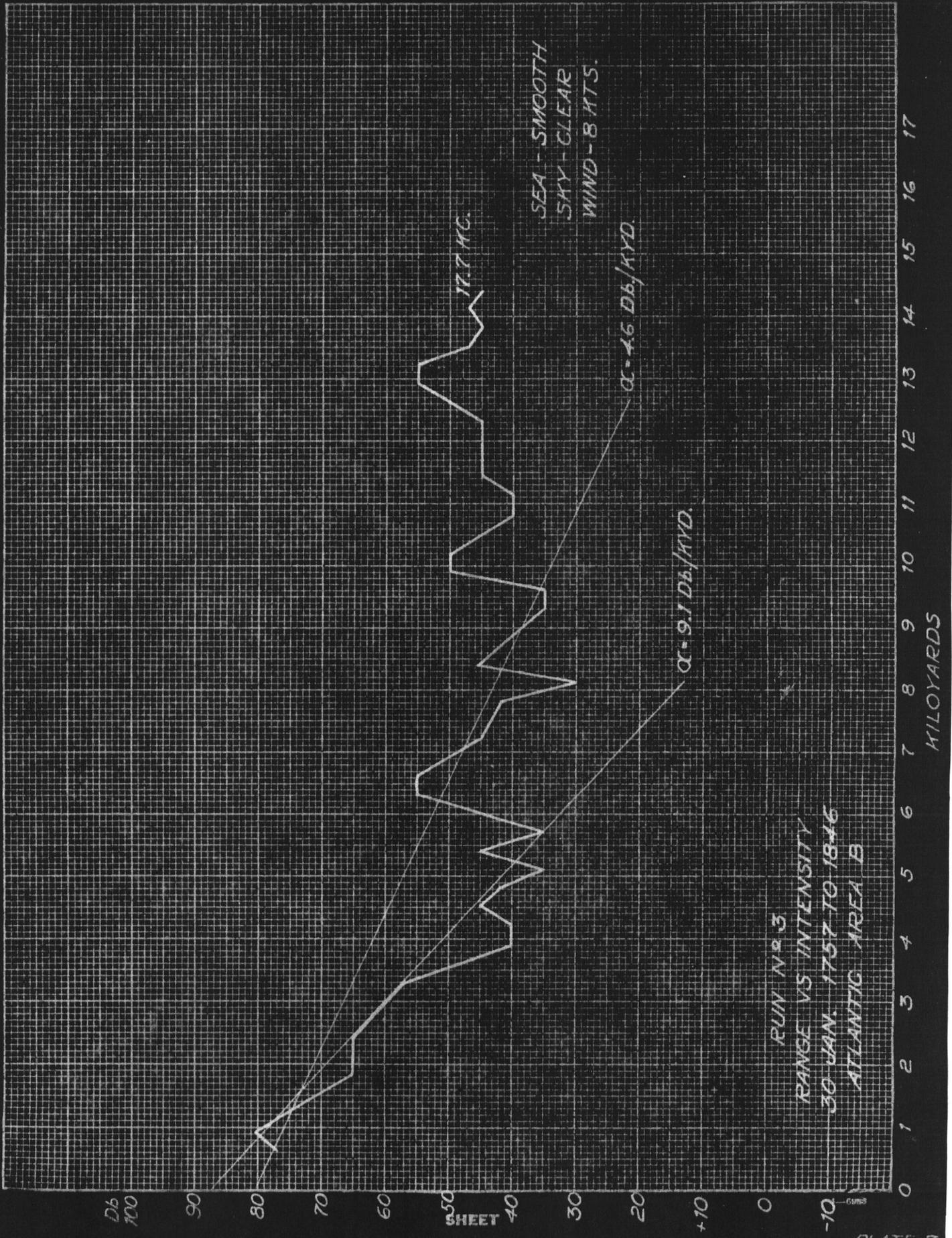


SEA-SMOOTH  
SKY-CLEAR  
WIND-5 KTS.

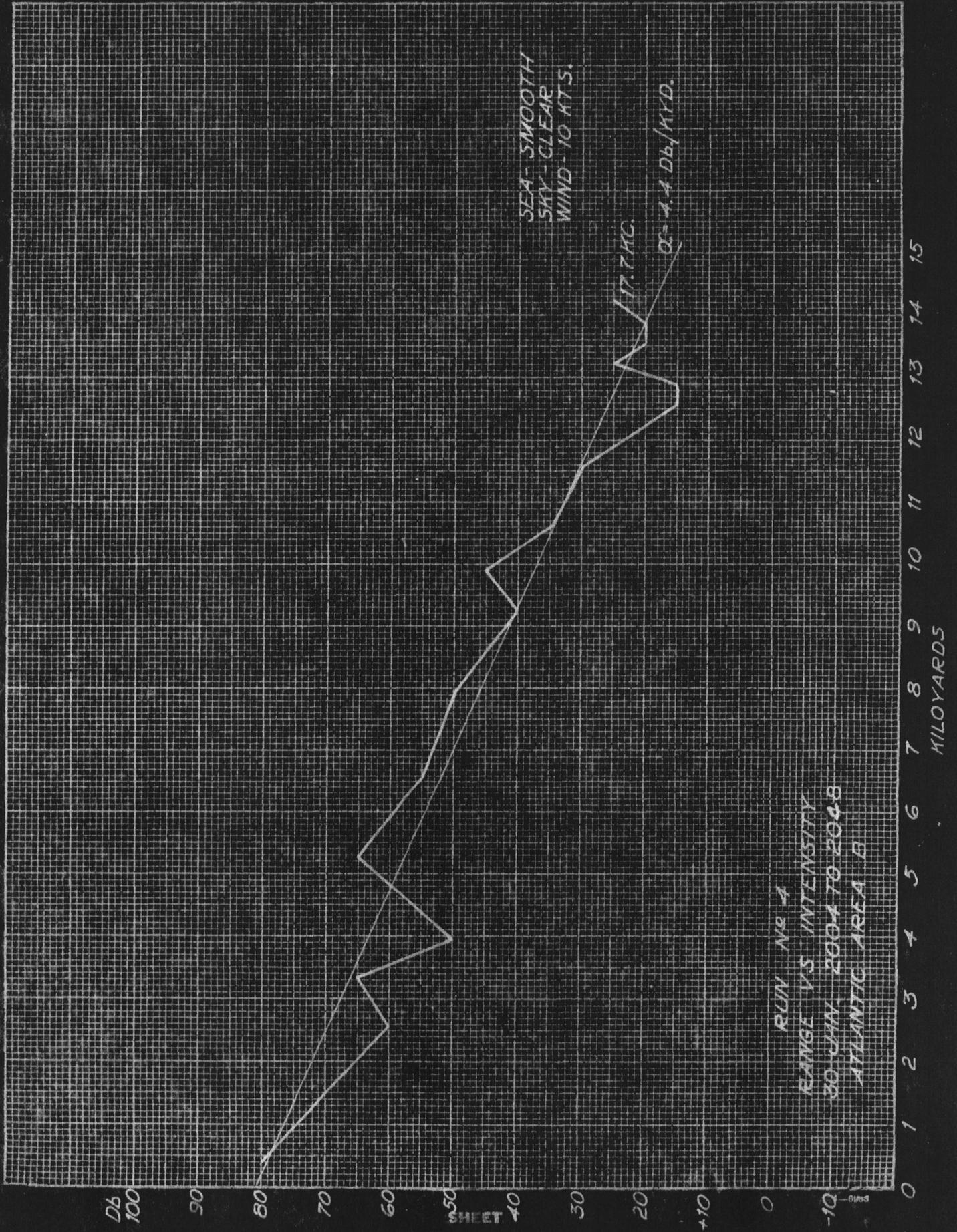
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RANGE VS INTENSITY  
30 JAN. 1956 TO 1638  
ATLANTIC AREA B

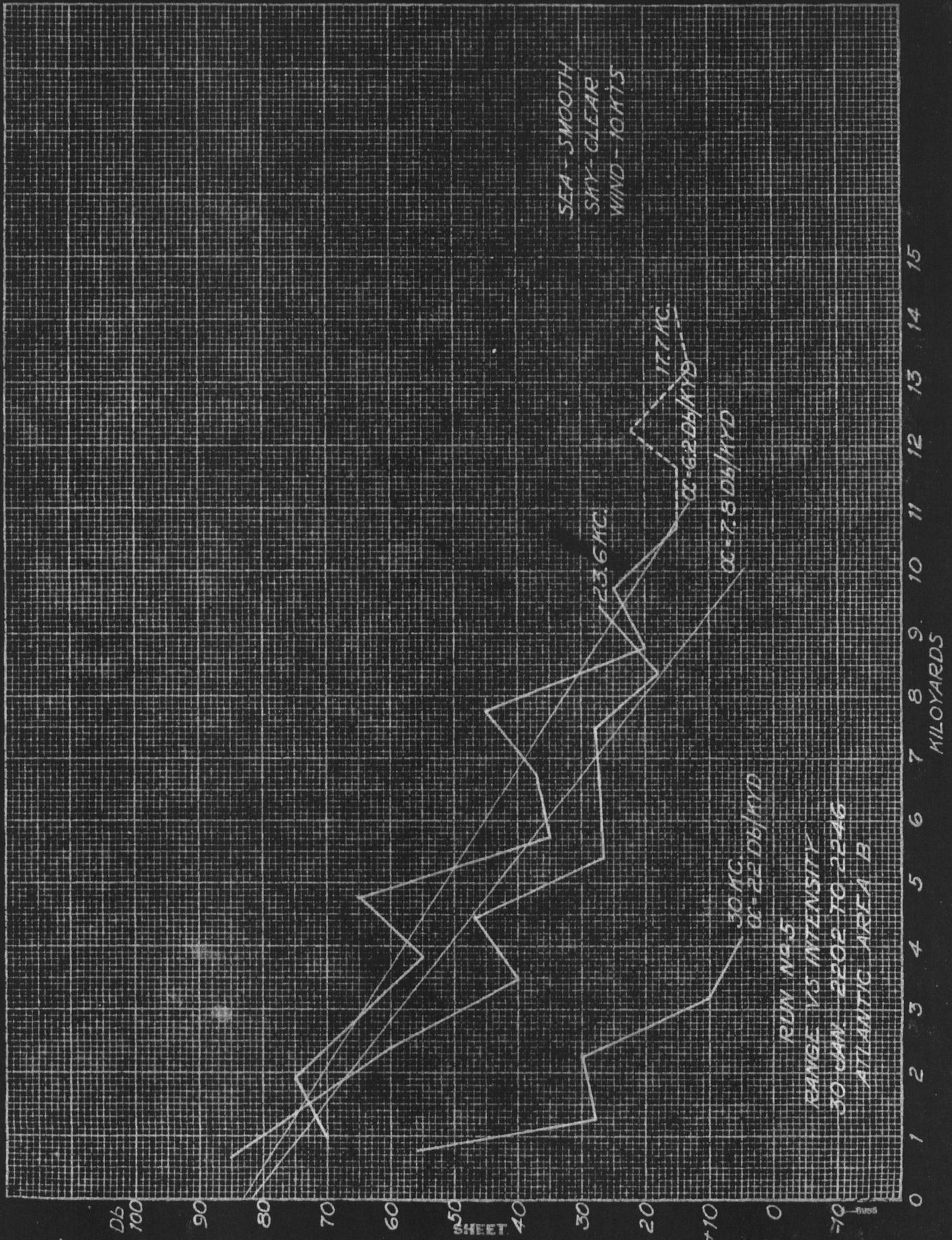
SHEET

PLATE 2

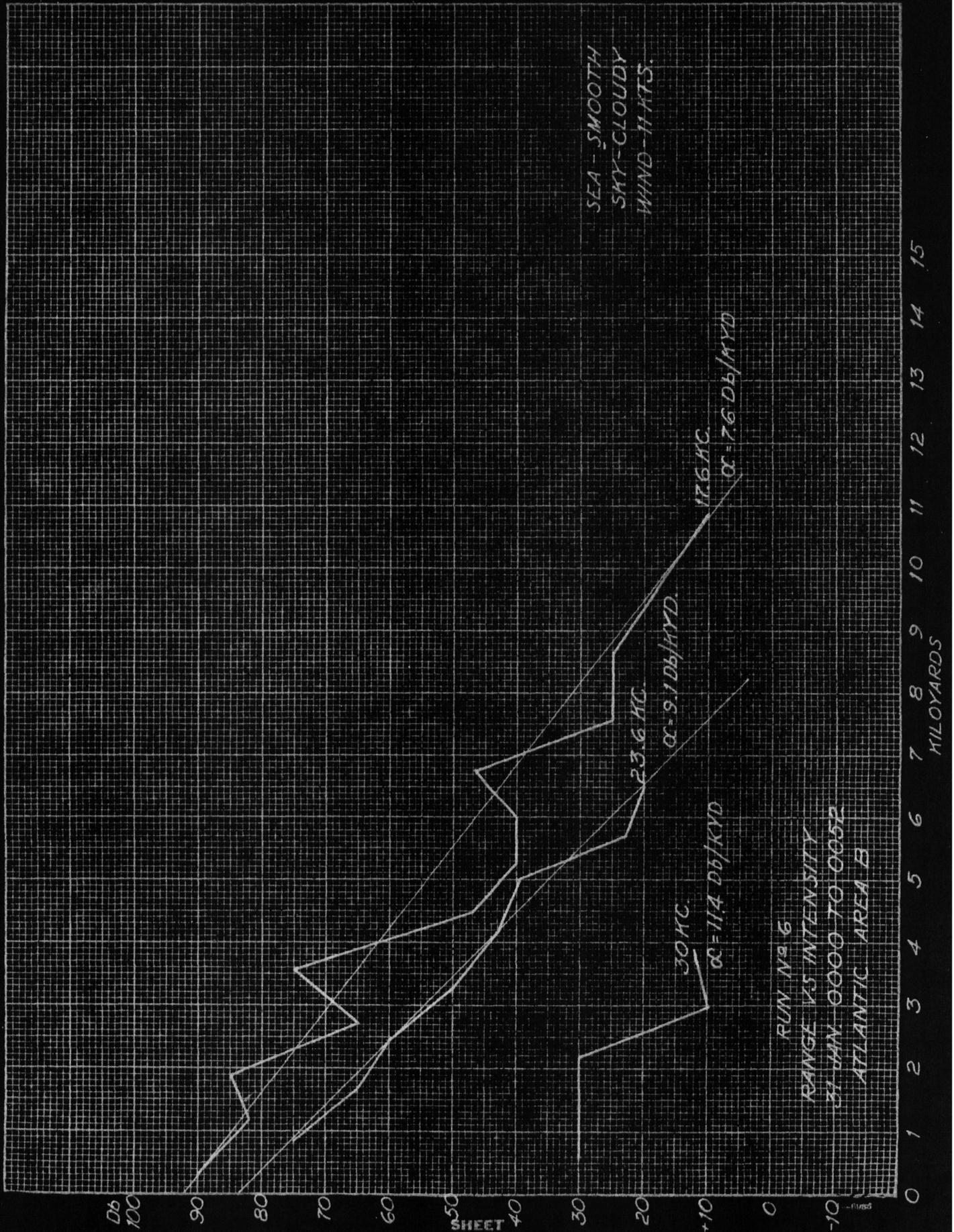


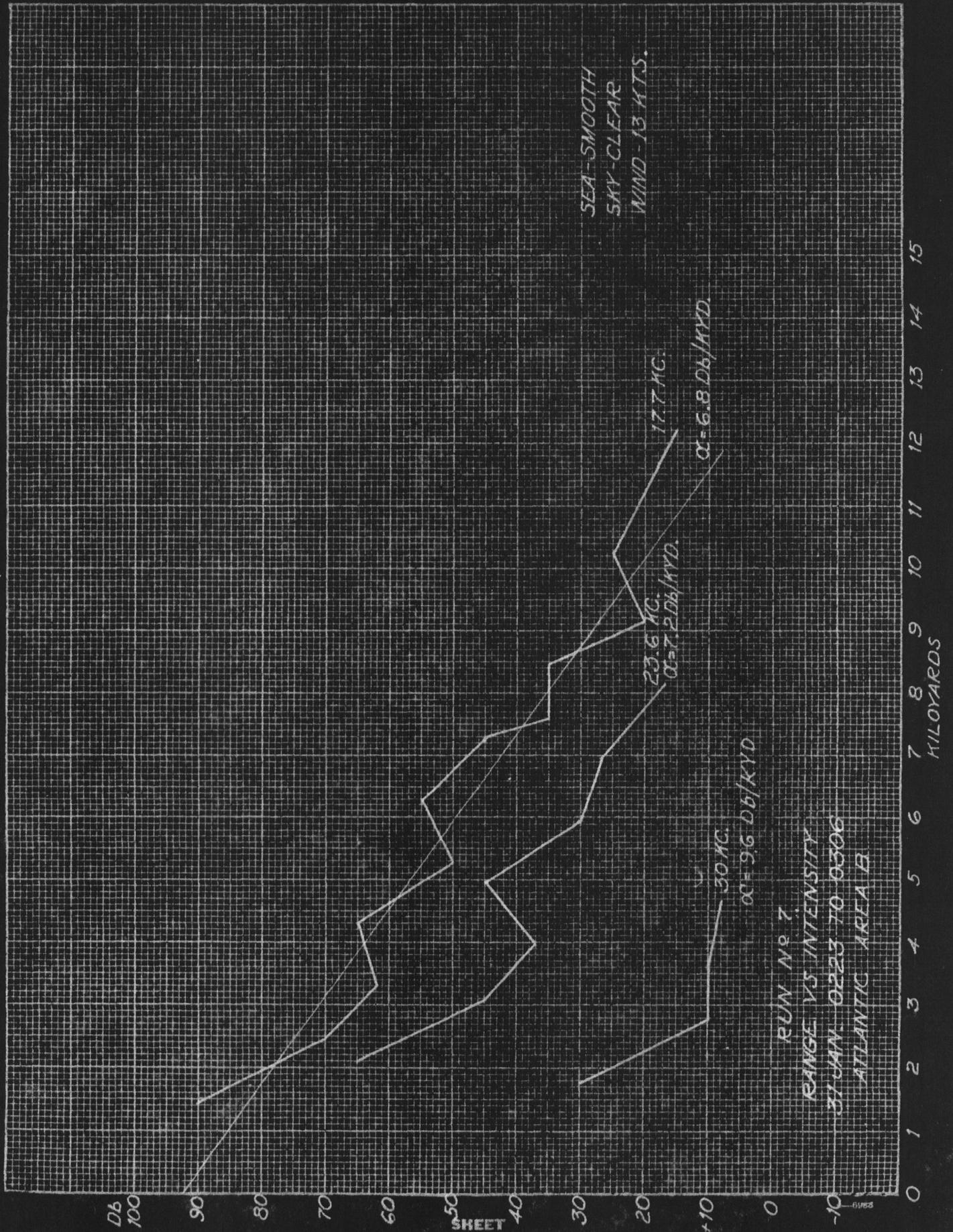
4-9883 N. R. L. 34



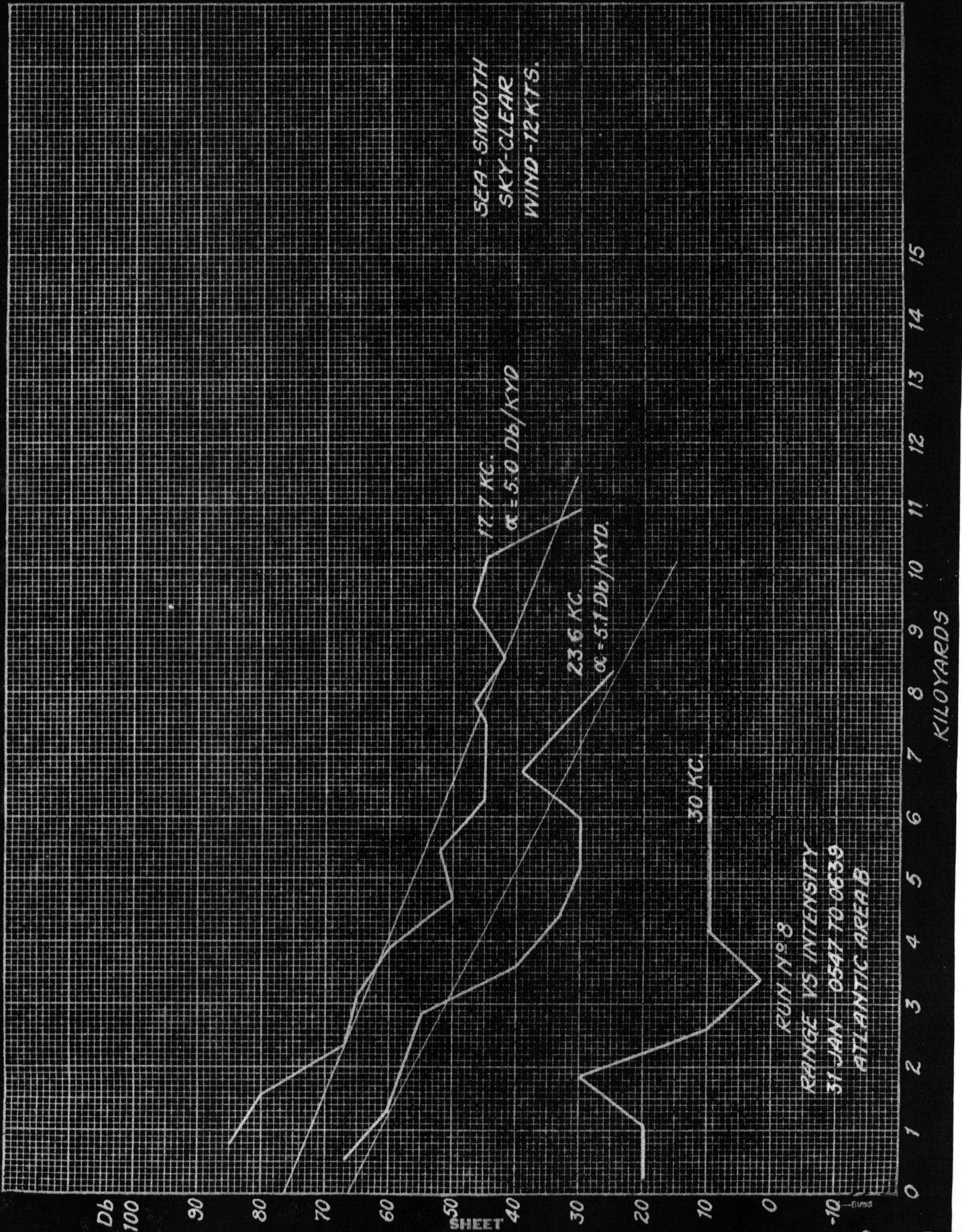


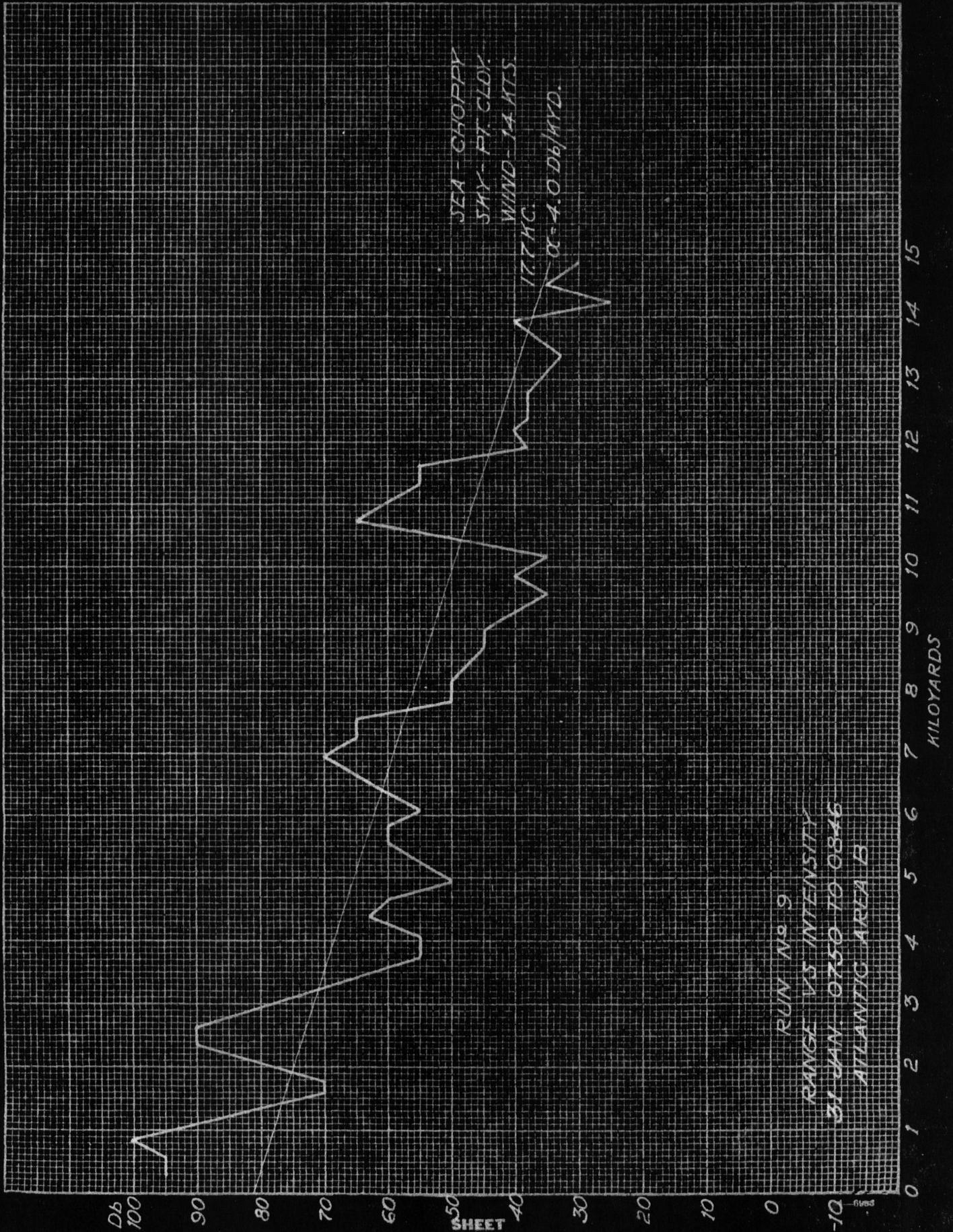
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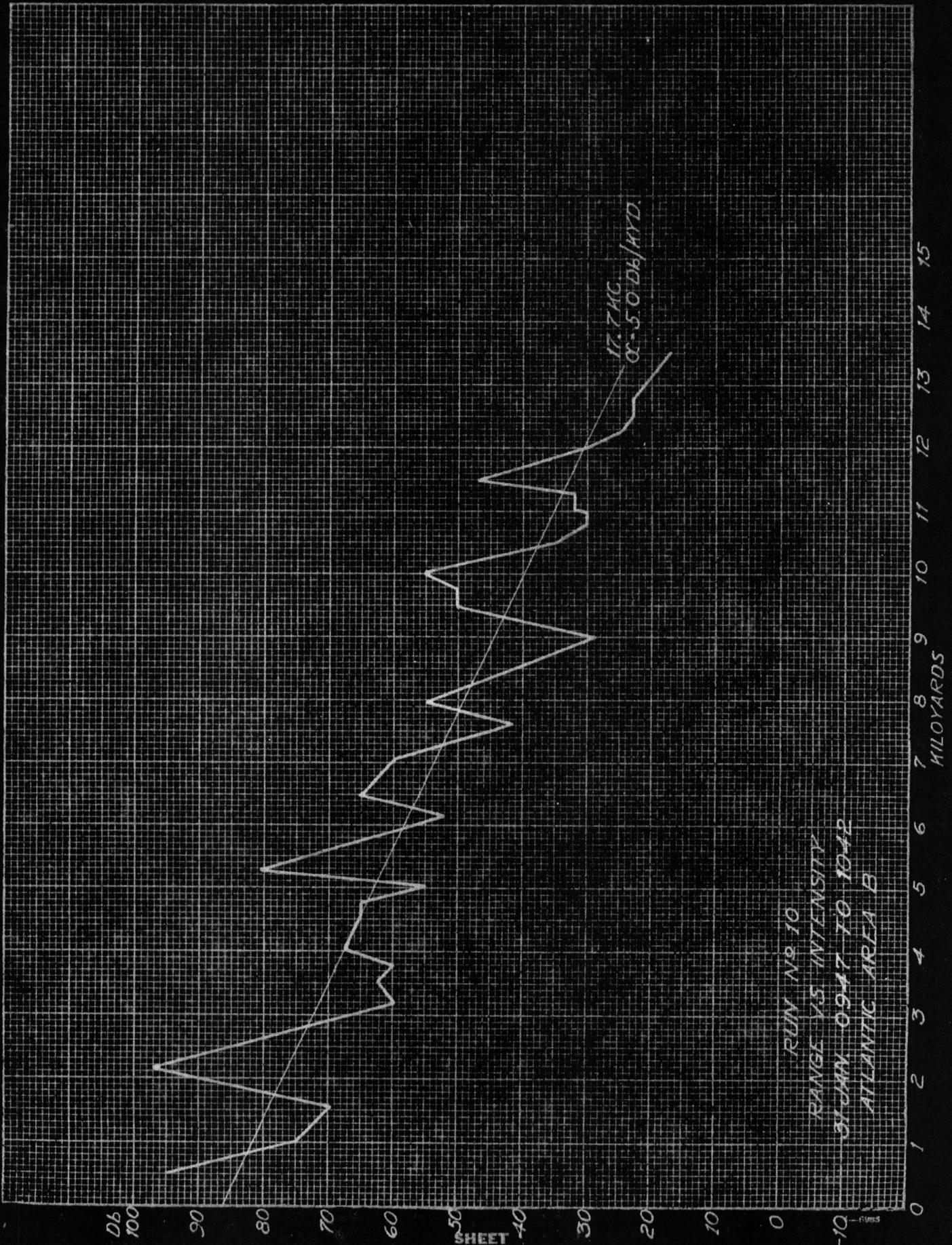


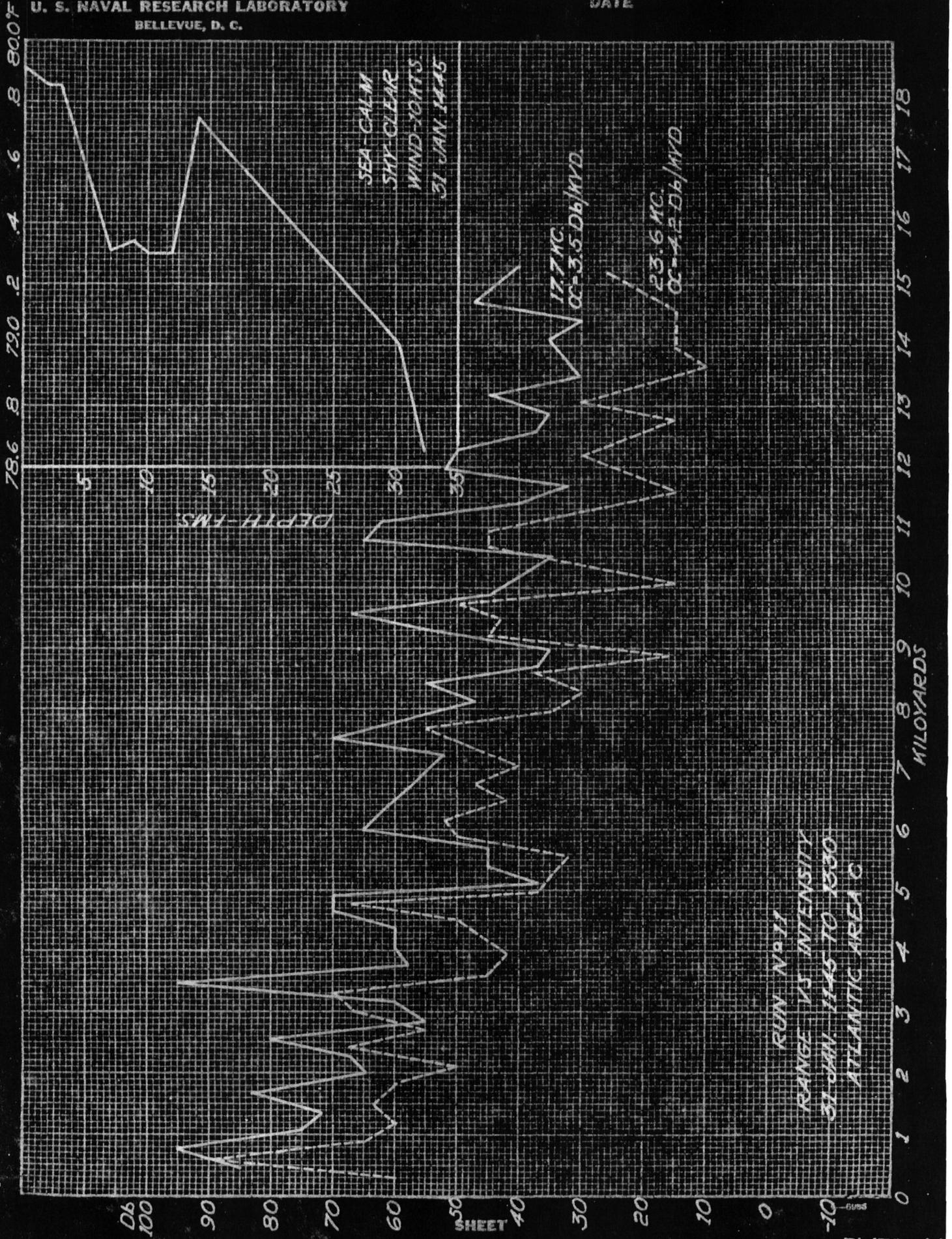


4-0053 N. R. L. 34

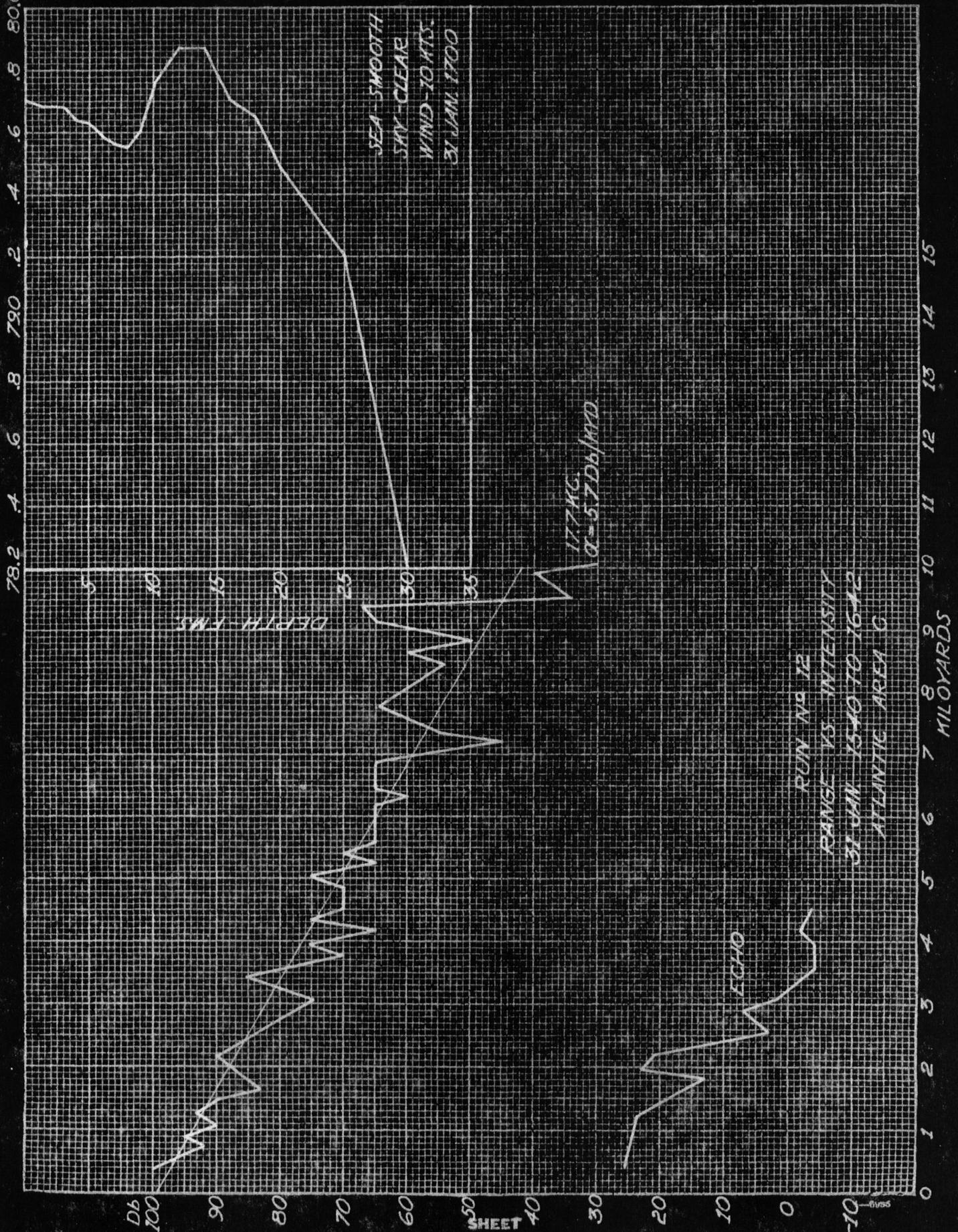


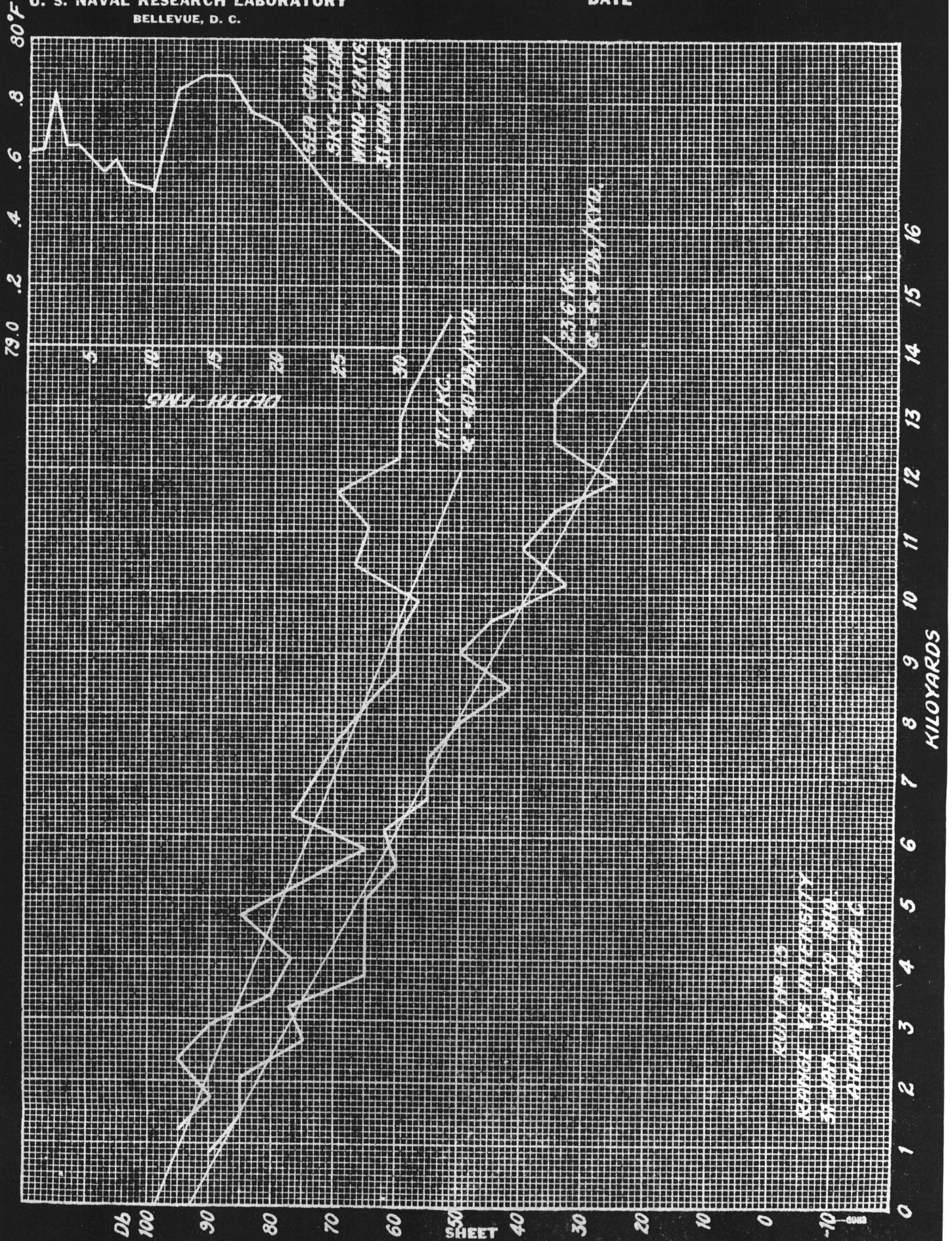




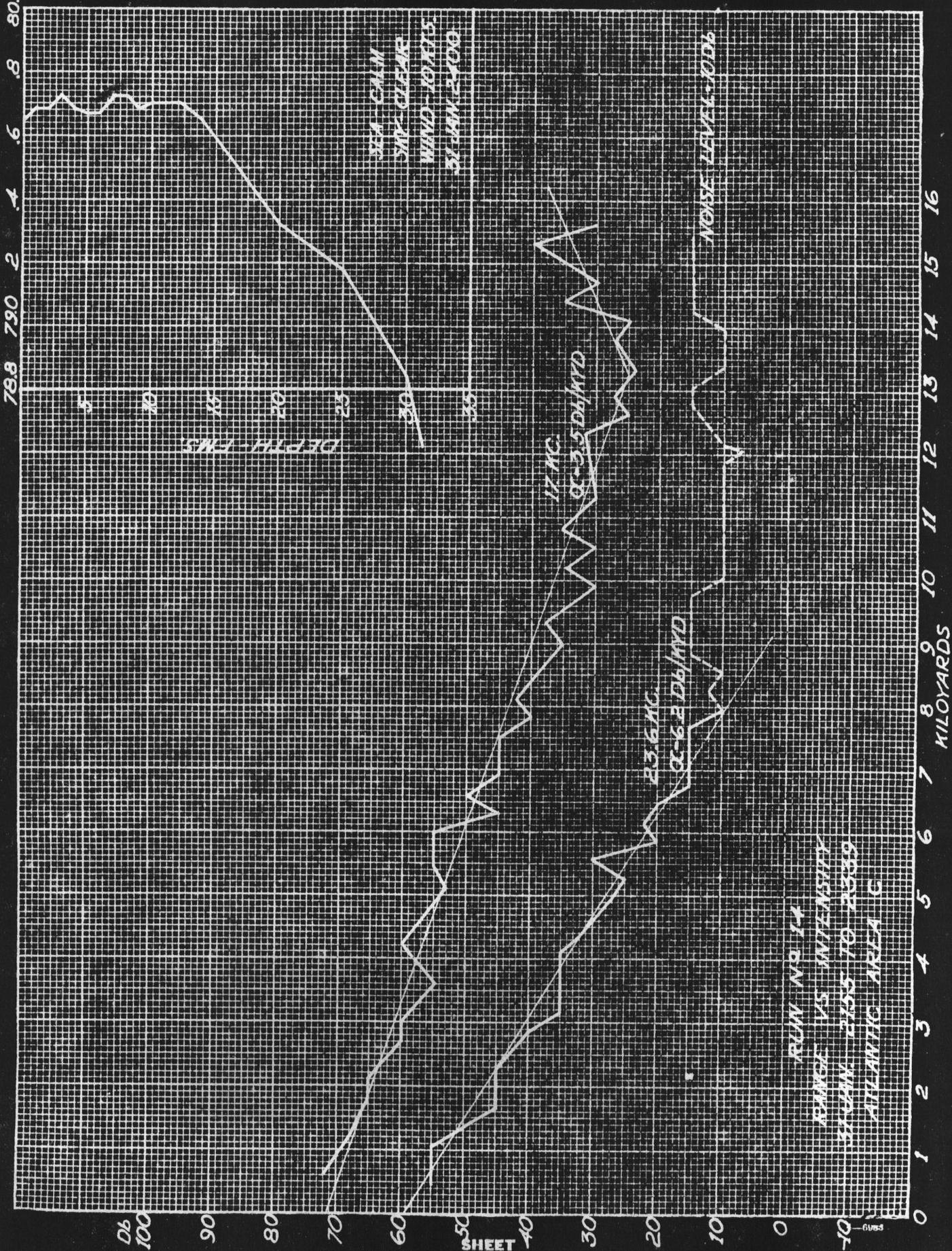


RUN No. 11  
RANGE VS INTENSITY  
31 JAN. 1145 TO 1330  
ATLANTIC AREA C

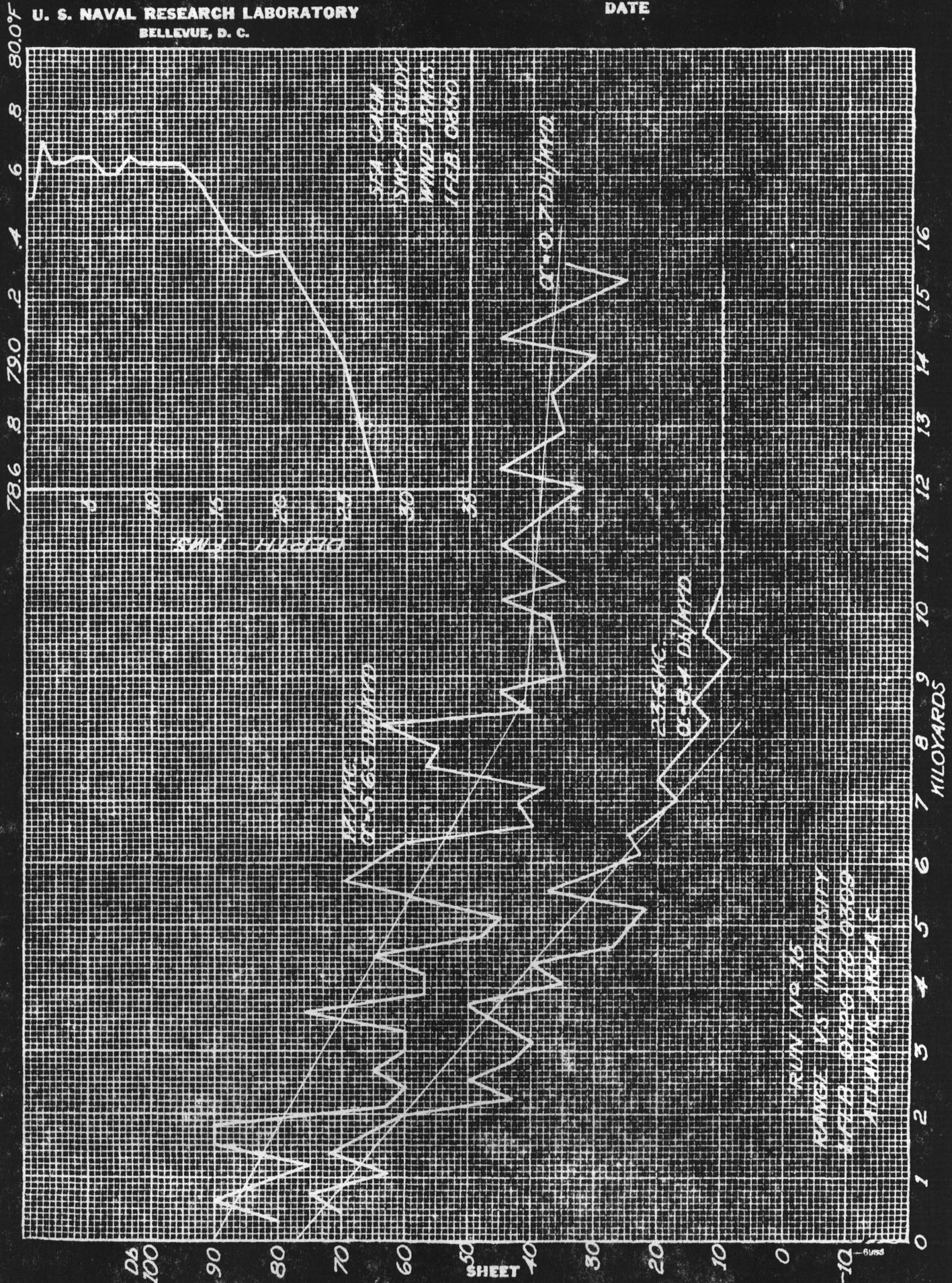




RUN # 13  
RANGE 15 INTENSITY  
31 JAN 1945 10 1945  
ATLANTIC OCEAN C



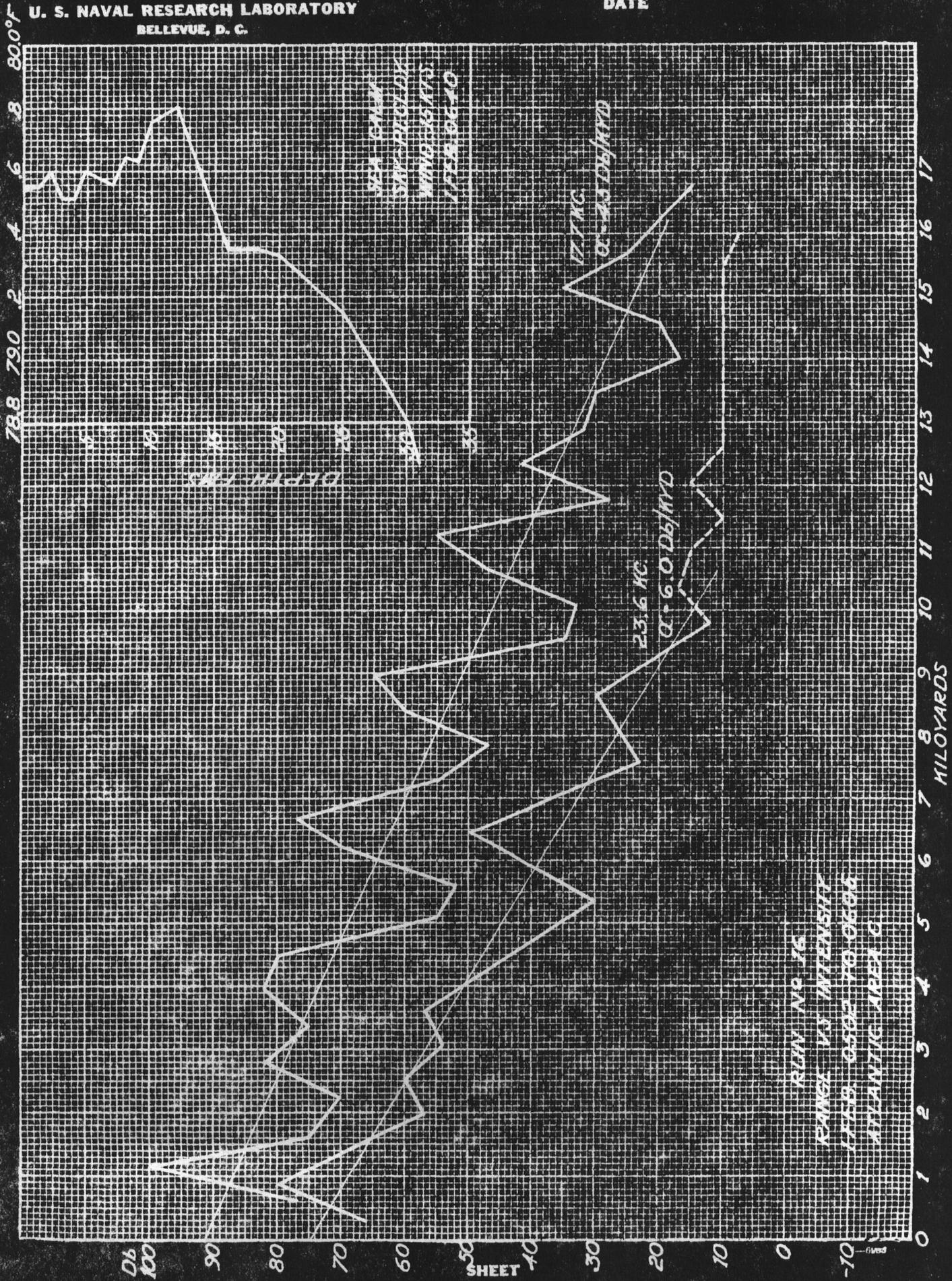
RUN NR 14  
 RANGE VS INTENSITY  
 SLAW 2155 TO 2339  
 ATLANTIC AREA C



RUN No 15  
 RANGE VS INTENSITY  
 17 FEB 02 00 TO 0300  
 ATLANTIC AREA C

Db  
100

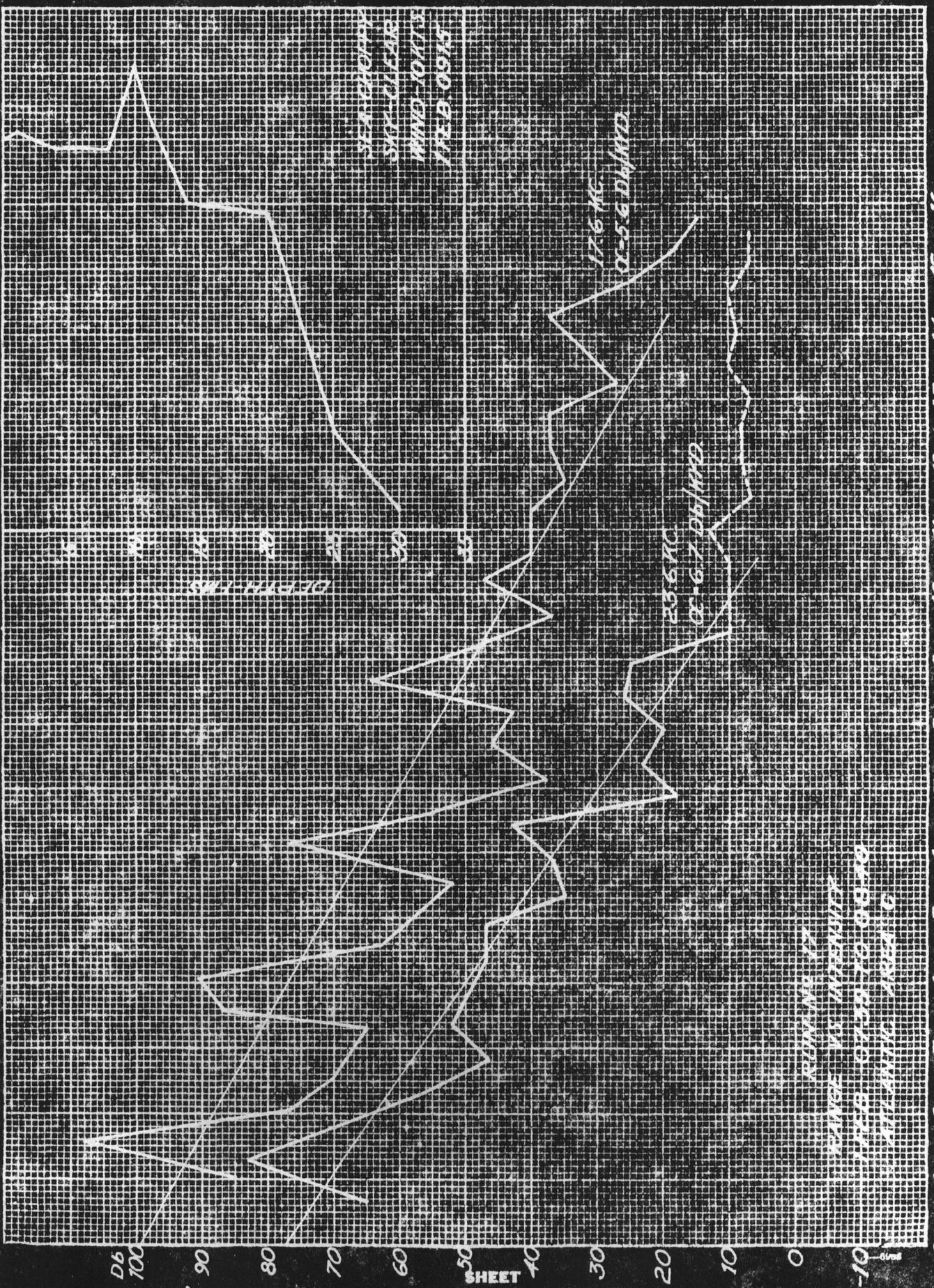
SHEET



78.8 79.0 2 4 6 8 80.0°F

DB  
100  
90  
80  
70  
60  
50  
40  
30  
20  
10  
0  
-10

SHEET

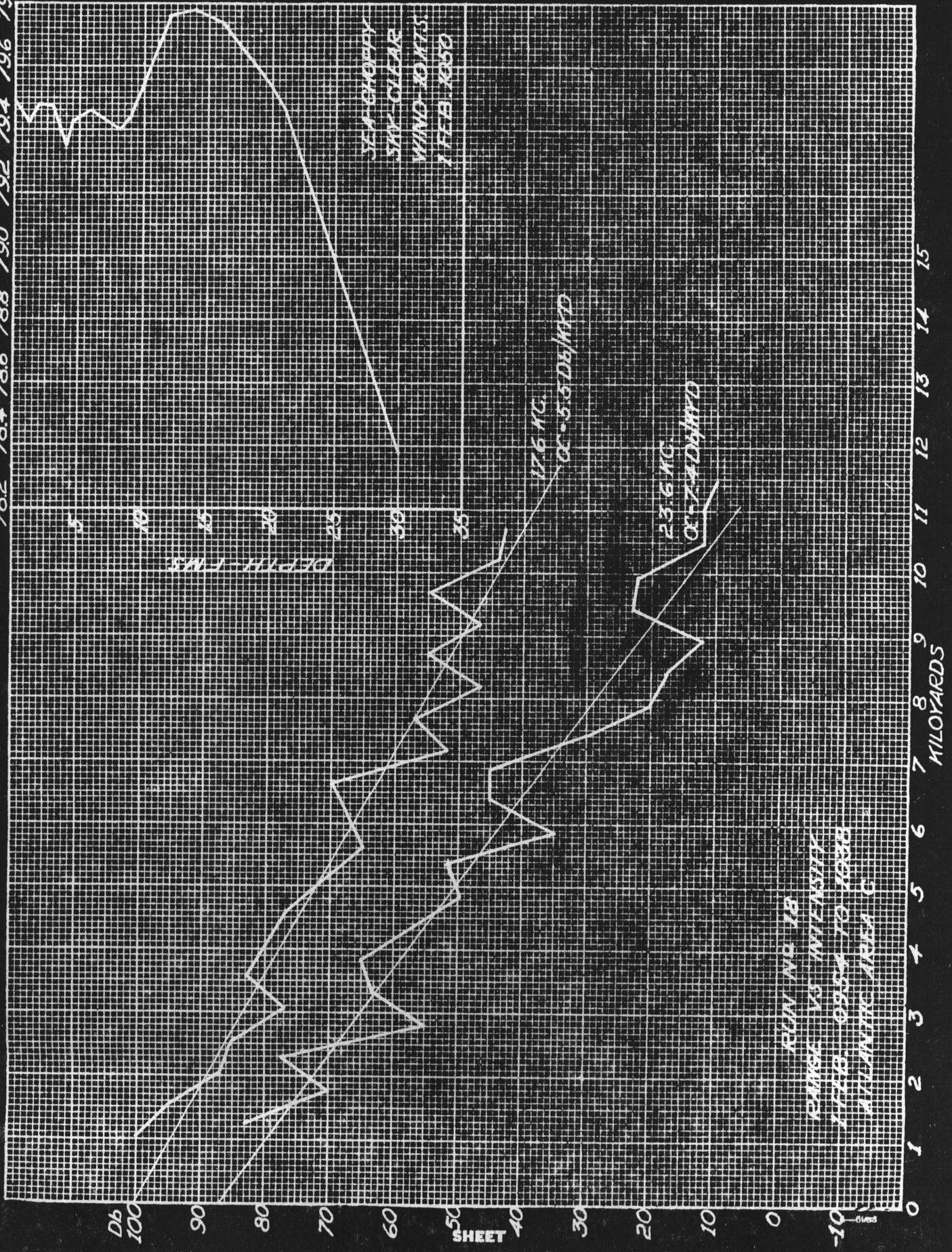


MILES

78.2 78.4 78.6 78.8 79.0 79.2 79.4 79.6 79.8°F

U. S. NAVAL RESEARCH LABORATORY  
BELLEVUE, D. C.

DATE

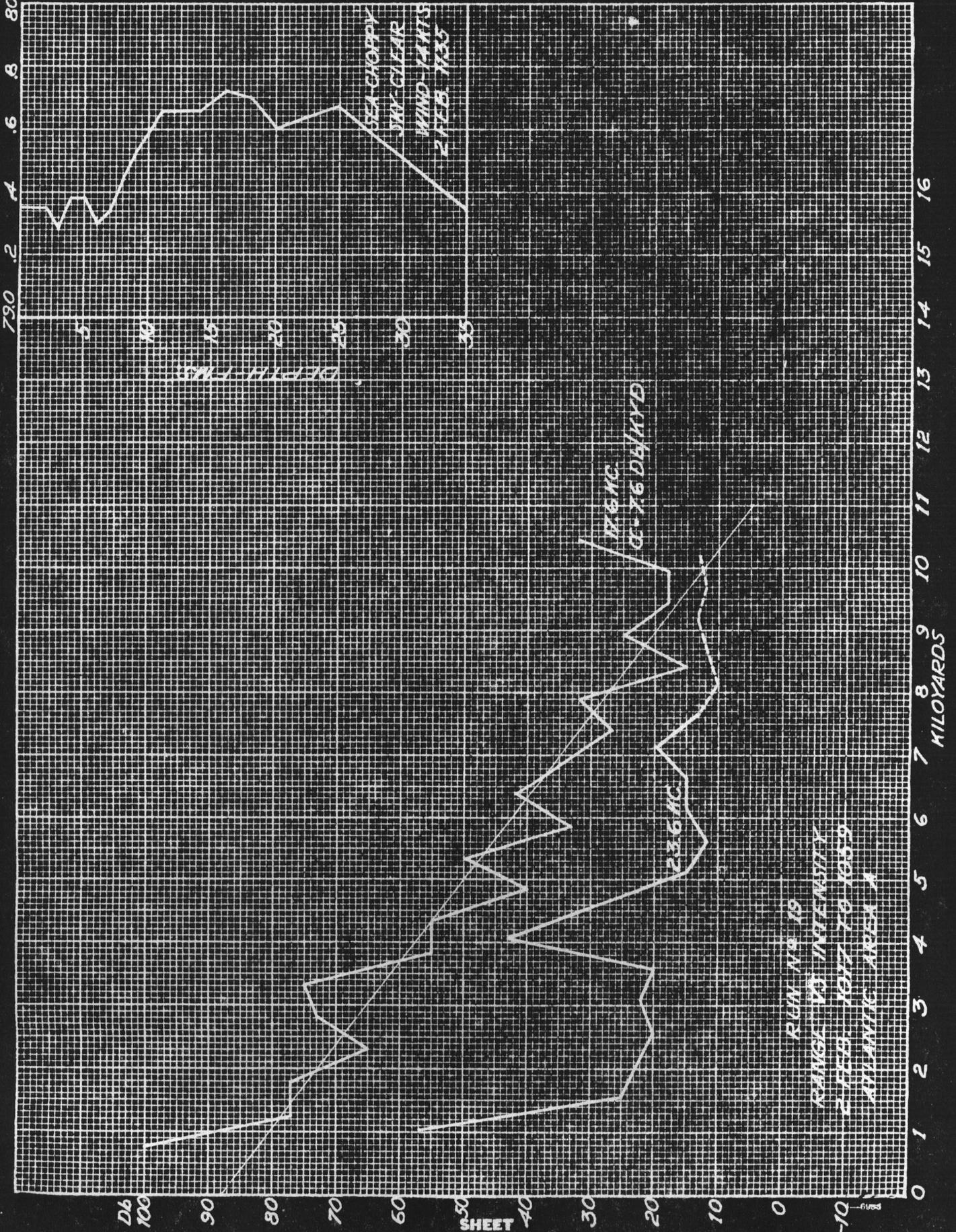


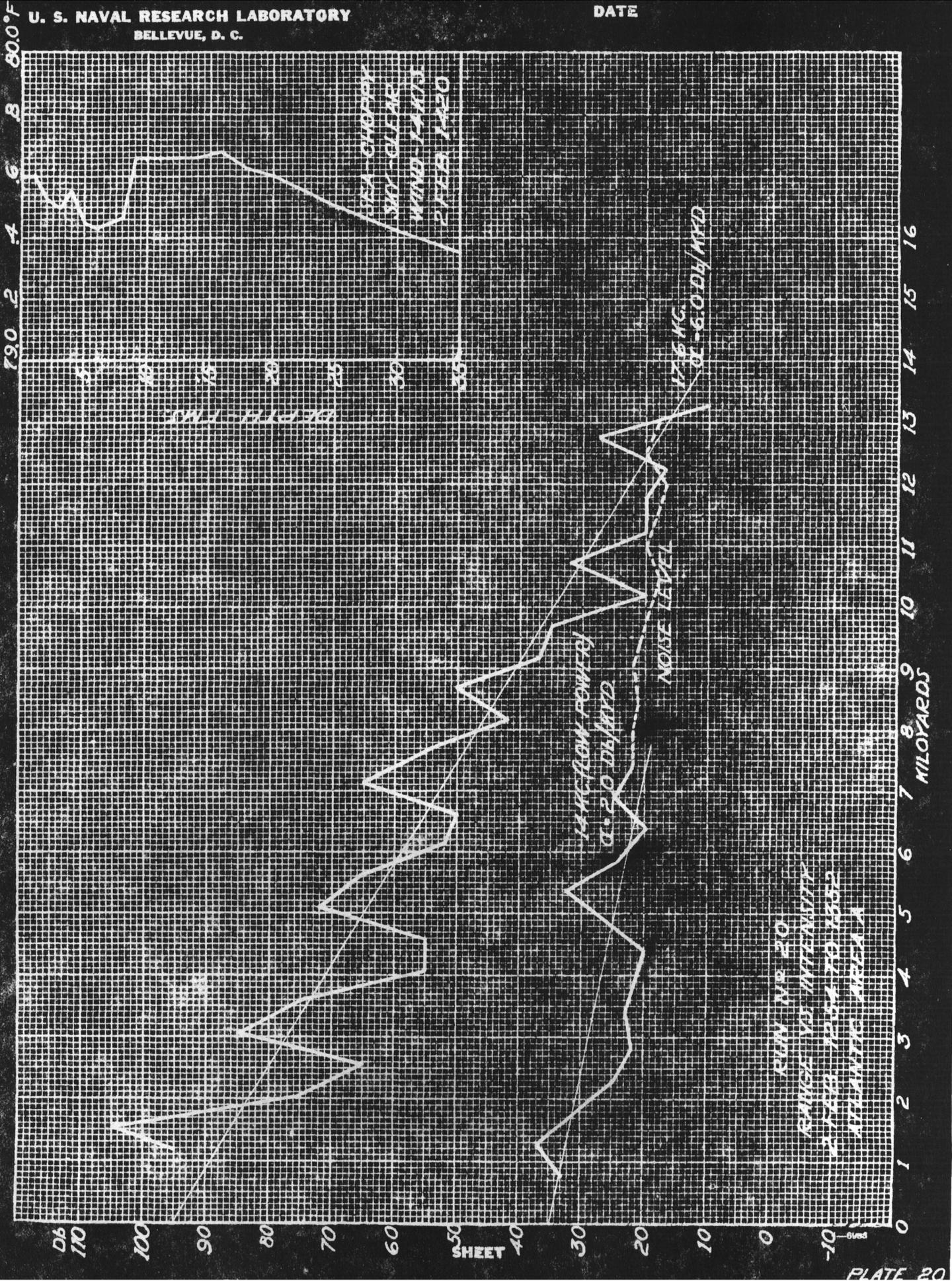
RUN NO. 18  
RANGE VS INTENSITY  
1 FEB. 1950 TO 1950  
ATLANTIC AREA C

Db  
100  
90  
80  
70  
60  
50  
40  
30  
20  
10  
0

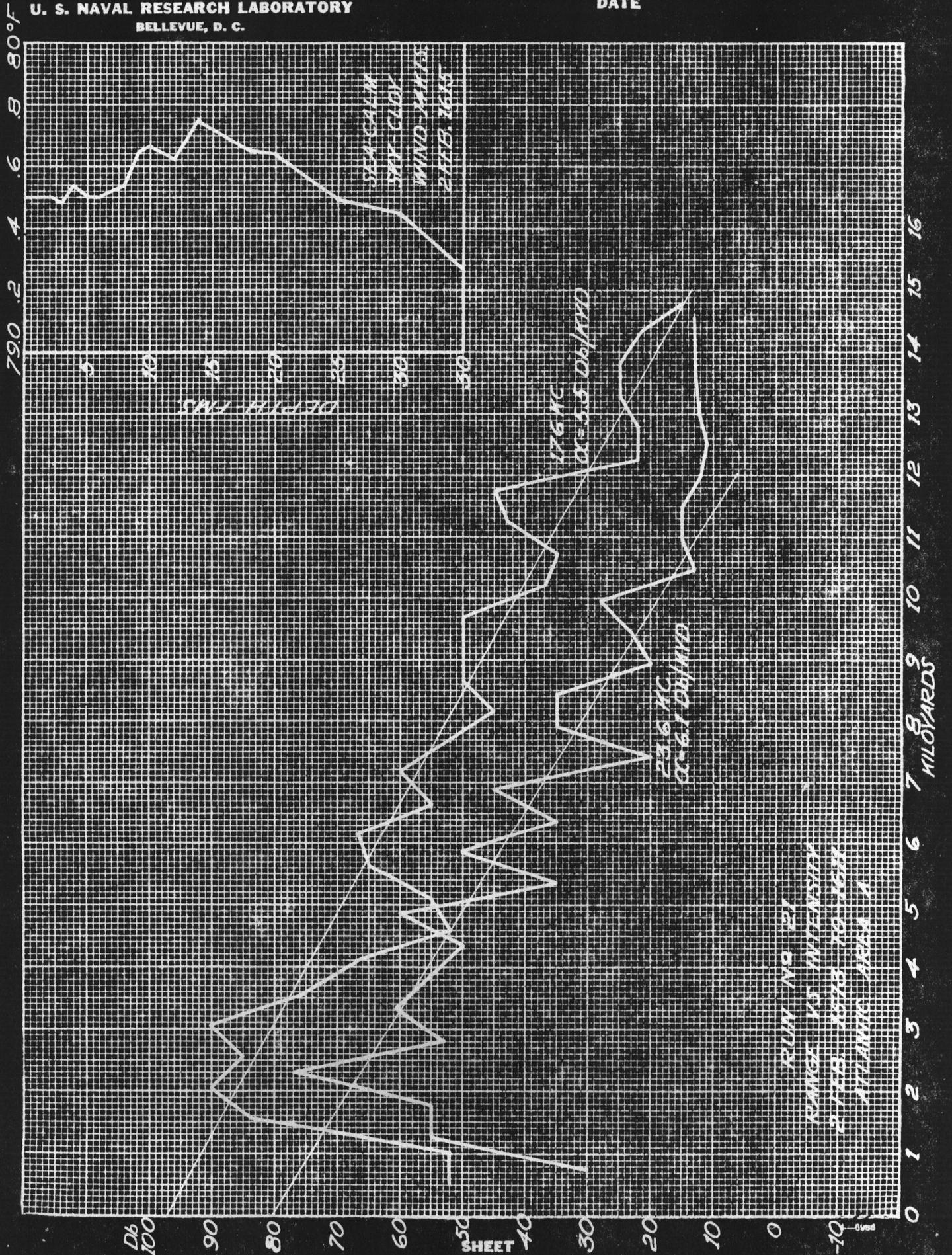
SHEET

0 1 2 3 4 5 6 7 8 9 10 11 12 13 14 15





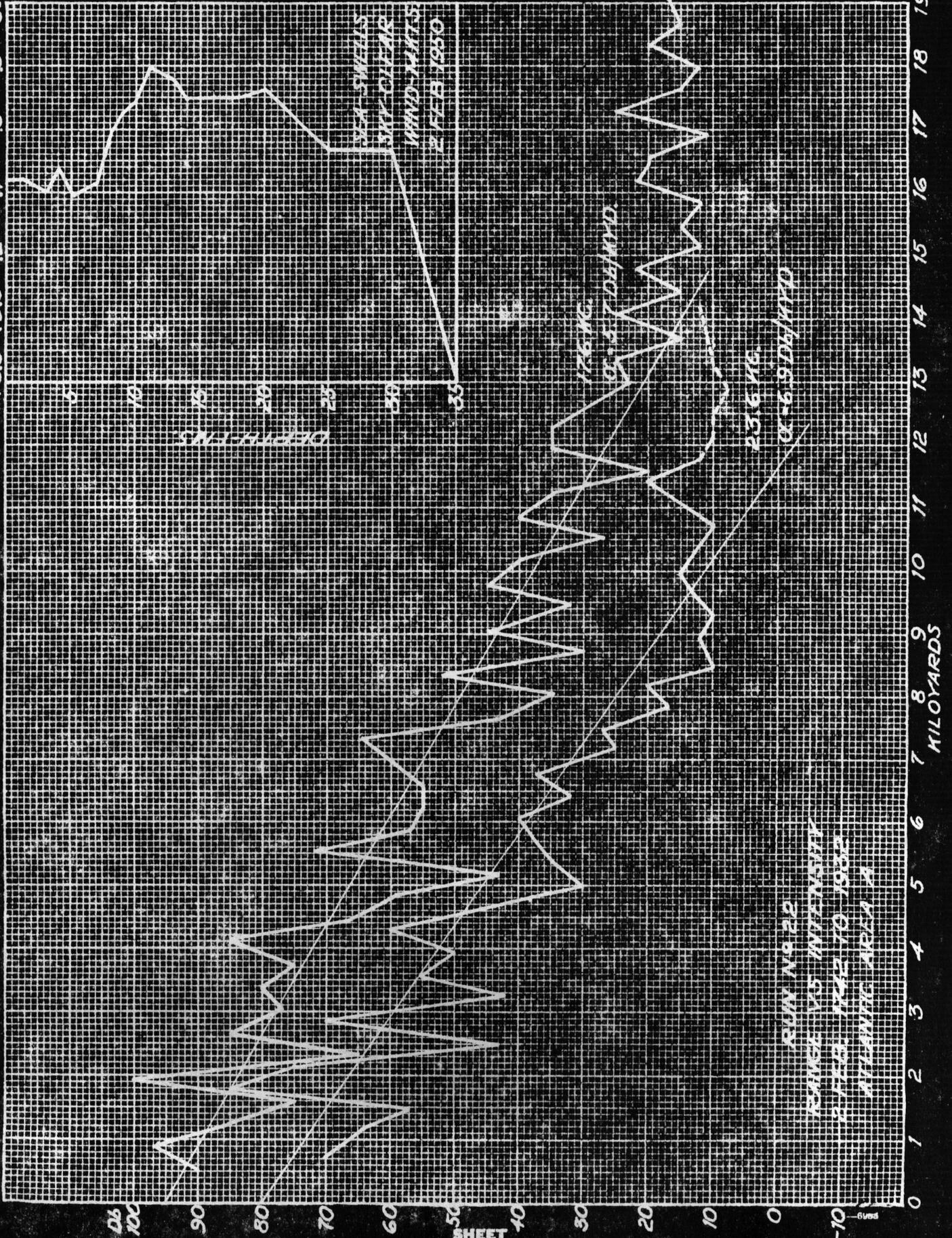
PLAN No 20  
RANGE VS INTENSITY  
2770 FAE0 TO 1852  
ATLANTIC AREA A



788 790 2 4 6 8 80.0°F

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DATE



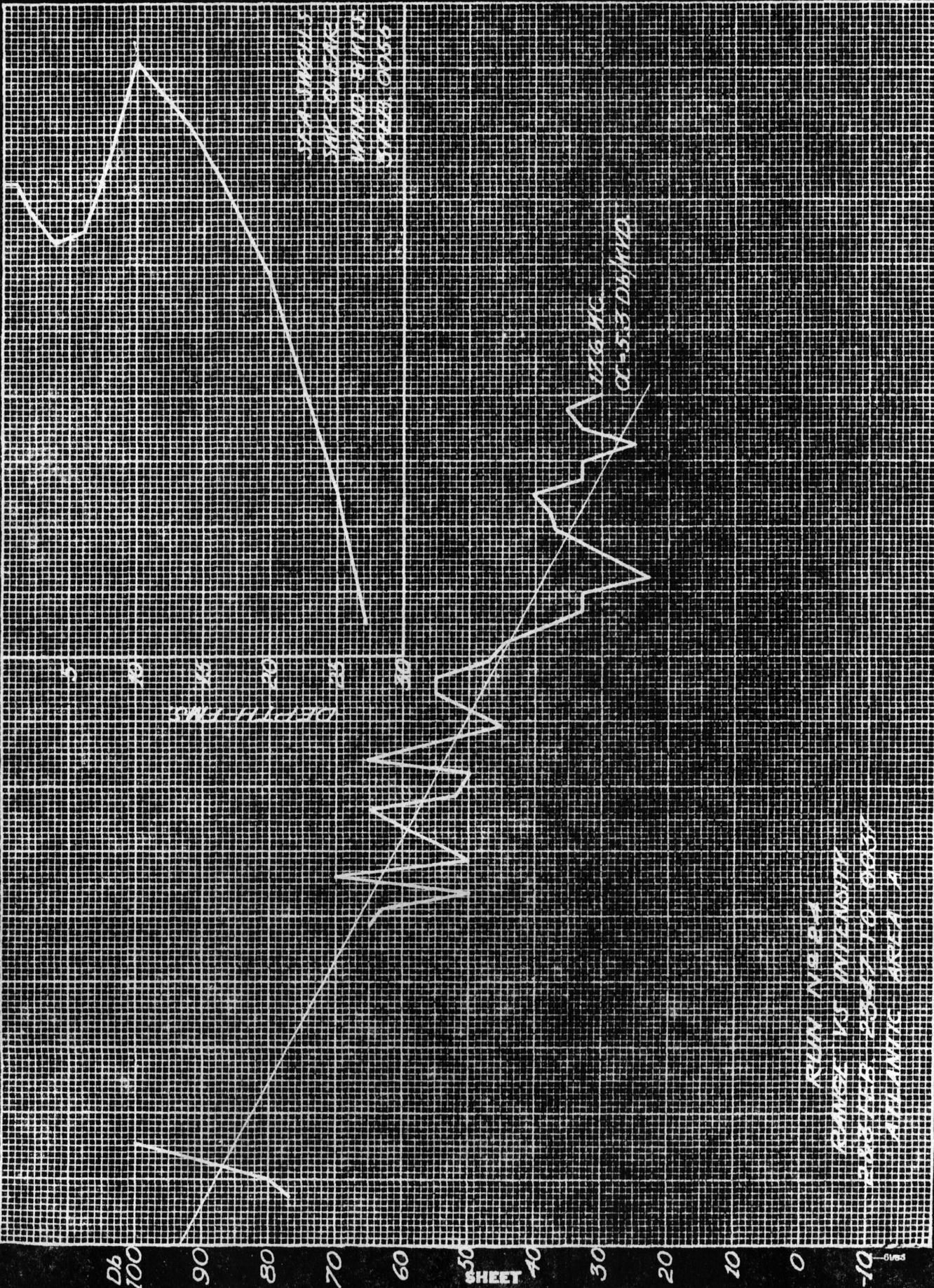
RUN N422  
RANGE VS INTENSITY  
2 FEB 1950 TO 1952  
ATLANTIC AREA A

MILOYARDS

SHEET

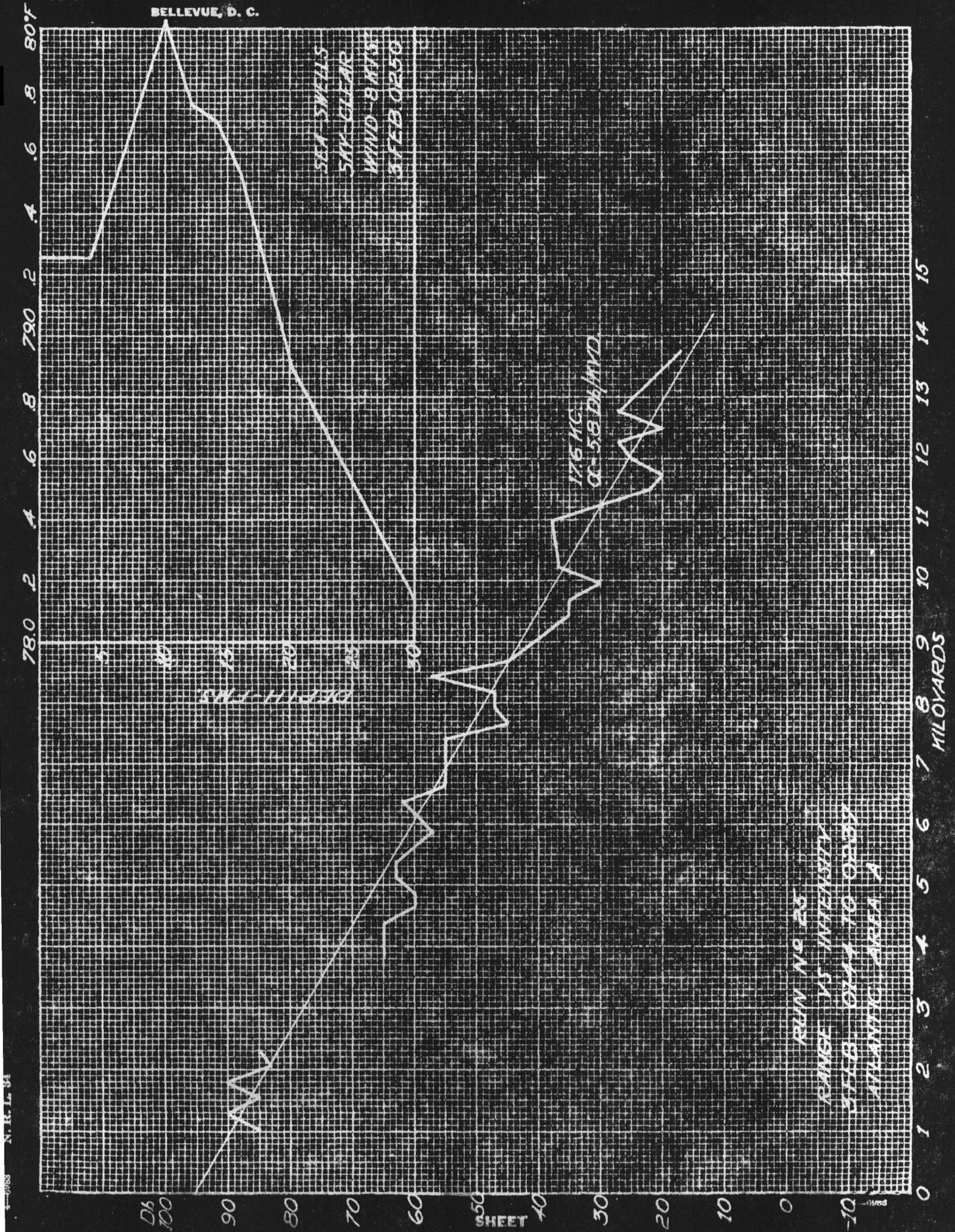


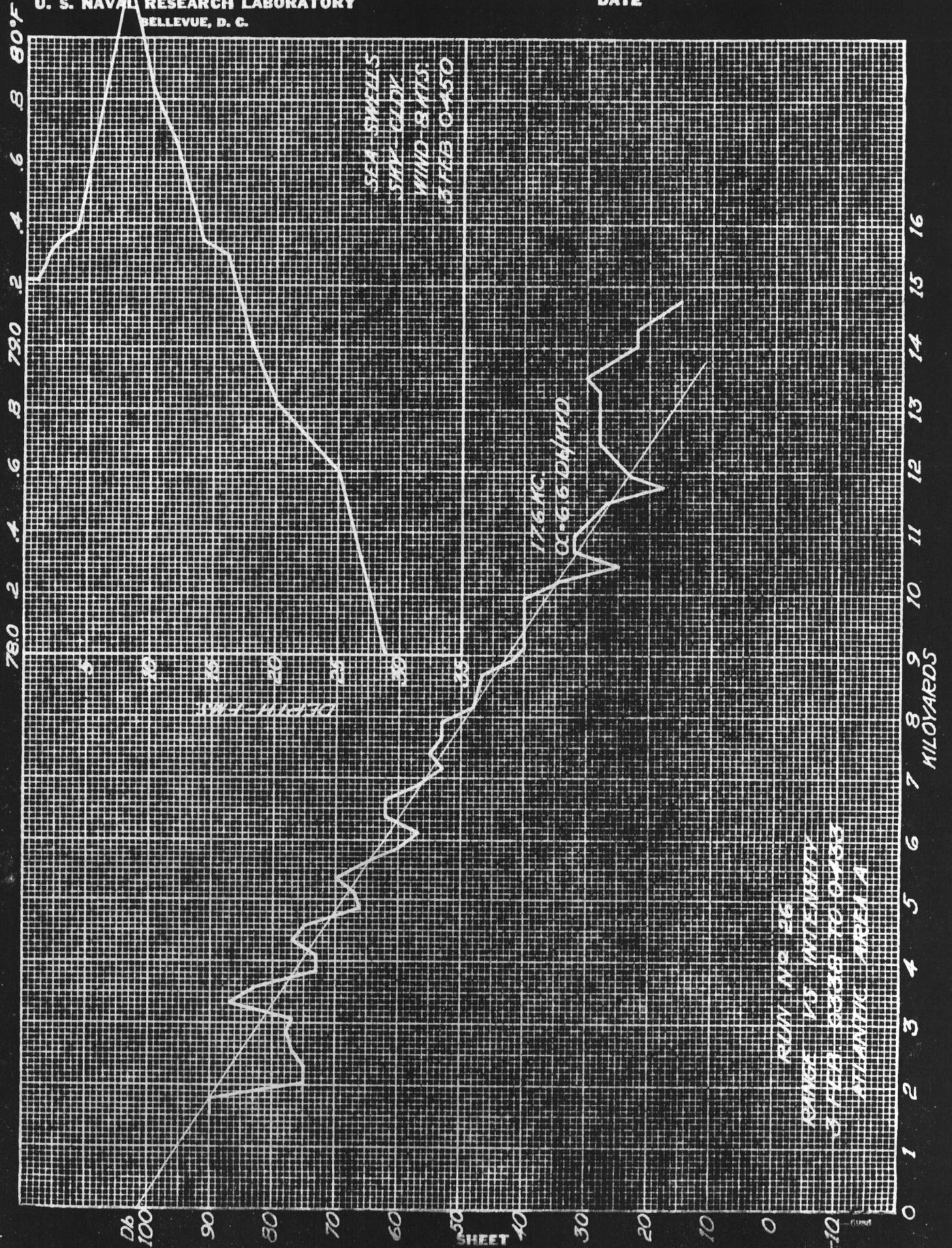
780 2 4 6 8 790 2 4 6 8 80°F



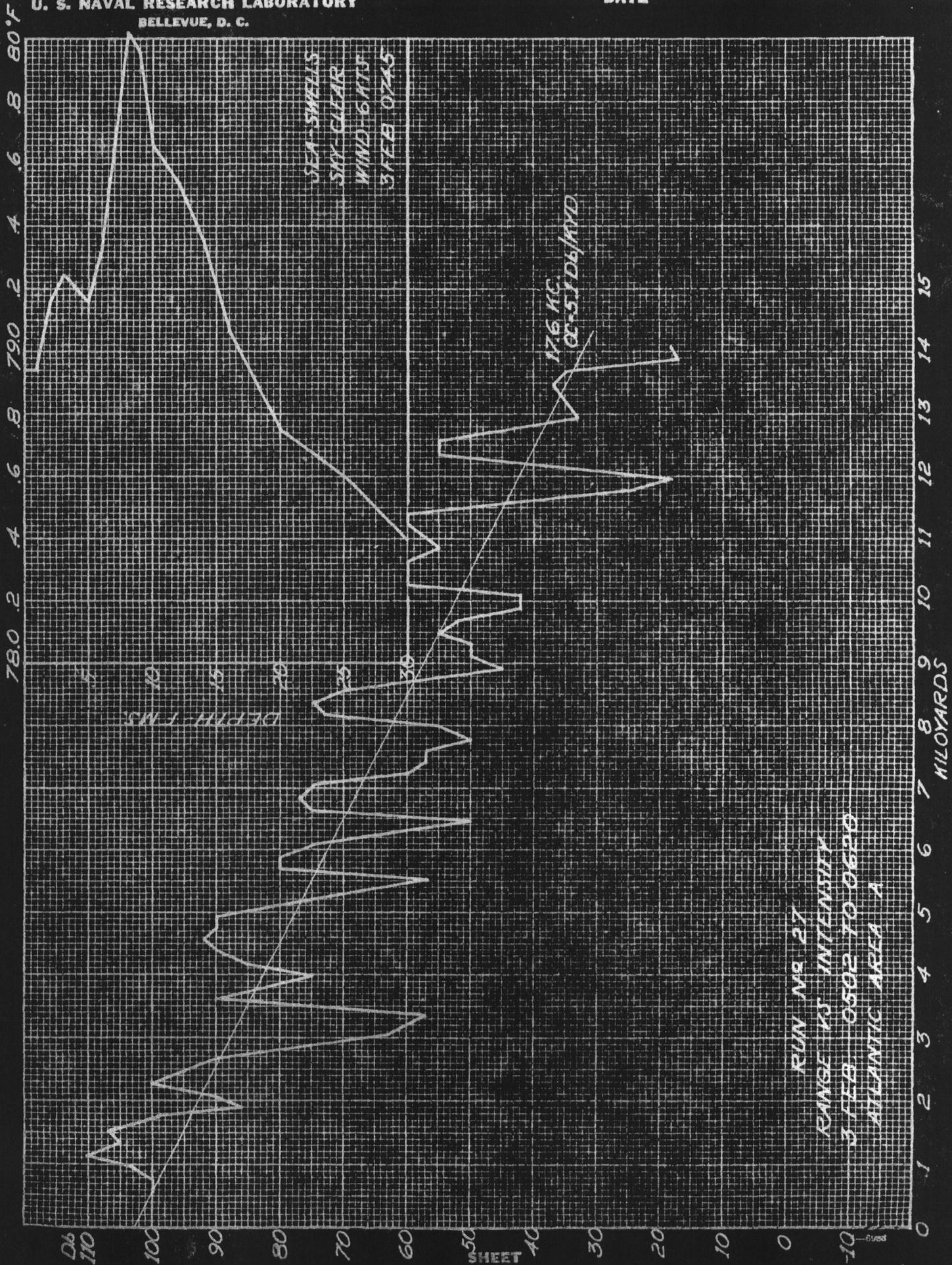
780 2 4 6 8 790 2 4 6 8 80°F

MILYARDS





RUN NO 26  
RANGE VS INTENSITY  
FILE# 0338 100433  
ATLANTIC AREA A

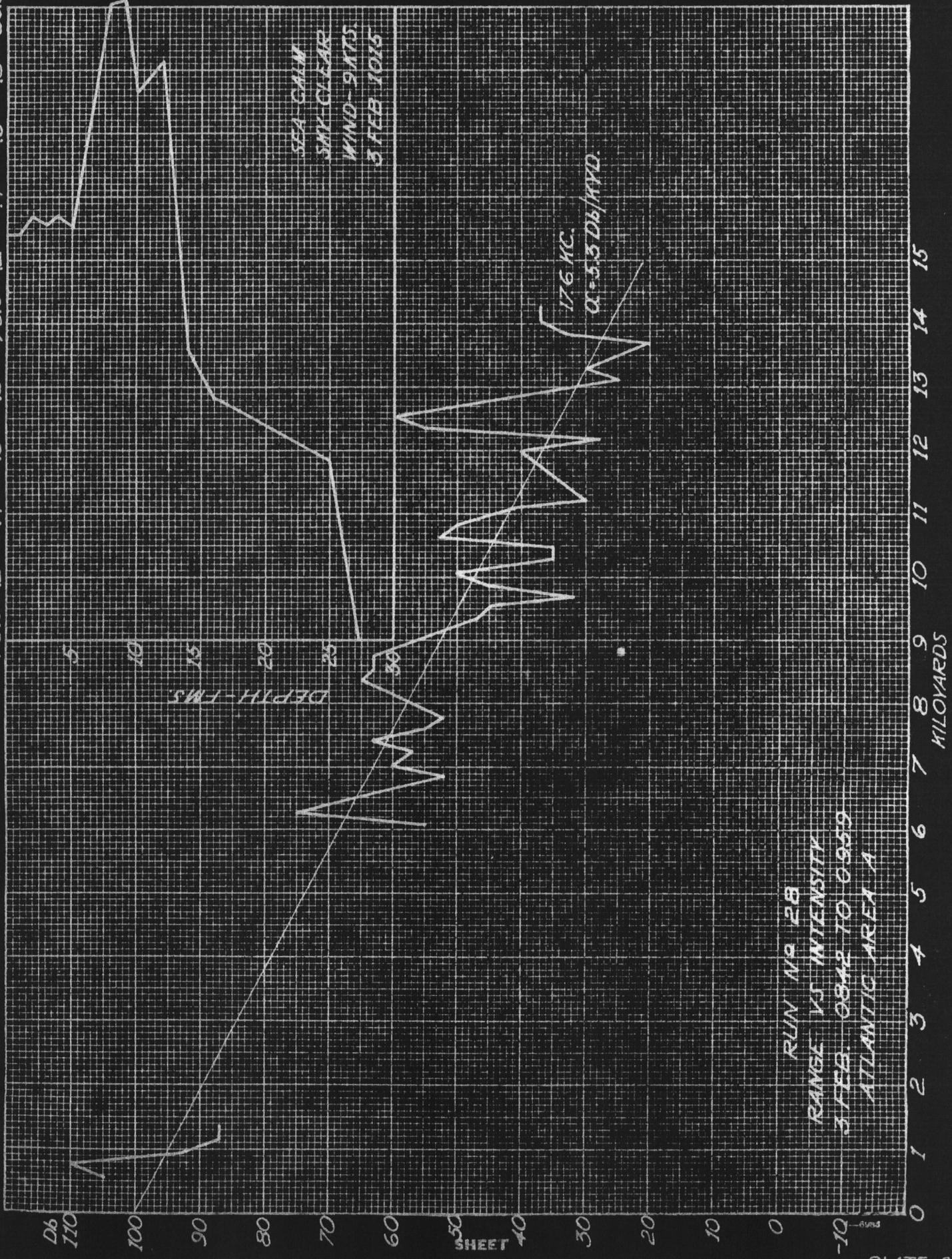


RUN NR 27  
 RANGE 15 INTENSITY  
 3 FEB 0502 TO 0620  
 ATLANTIC AREA A

78.0 2 .4 .6 .8 79.0 2 .4 .6 .8 80.0°F

U. S. NAVAL RESEARCH LABORATORY  
BELLEVUE, D. C.

DATE



0 1 2 3 4 5 6 7 8 9 10 11 12 13 14 15  
MILES

RUN NO 28  
RANGE VS INTENSITY  
3 FEB 0842 TO 0959  
ATLANTIC AREA A

Db  
110  
100  
90  
80  
70  
60  
50  
40  
30  
20  
10  
0

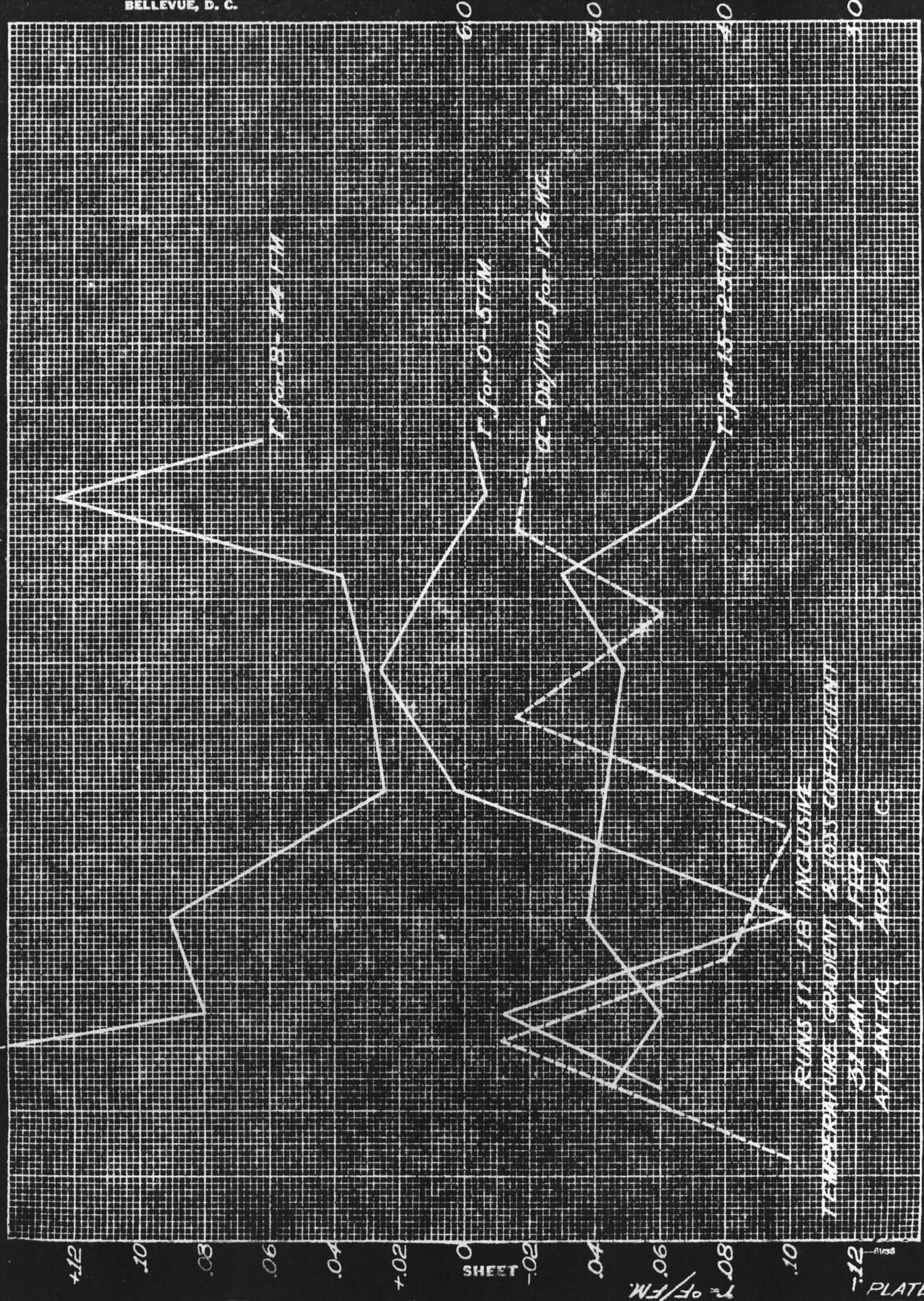
DEPTH - FMS

SHEET

24

L. 34

4-1963



12 14 16 18 20 22 24 :02 04 06 08 10 12  
 31 JAN. 1 FEB.  
 CLOCK TIME

SHEET

PLATE 29

→.24

P 11:00-11:10 AM

T for 07-37M

T for 15-237M

OC-DP/KVD for 17.6 MG.

RUNS 19-28 INCLUSIVE  
TEMPERATURE GRADIENTS & LOSS COEFFICIENTS  
2 & 3 FEB  
ATLANTIC AREA A

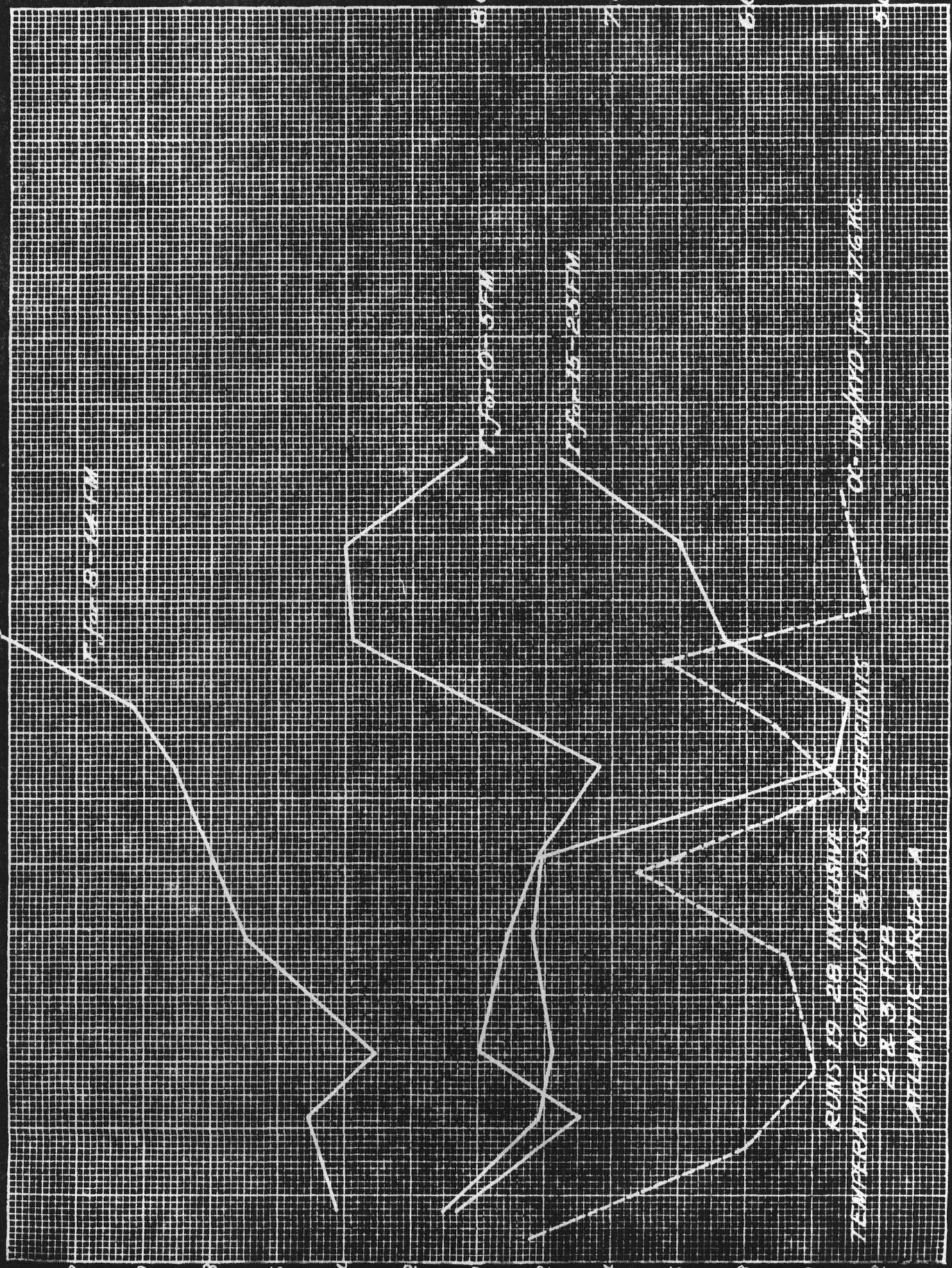
12 14 16 18 20 22 24 02 04 06 08 10 12  
2 FEB.  
3 FEB.  
CLOCK TIME

N. R. L. 34

4-4083

T = °F/M

PLATE 30





58 59 60 61 62 63 64 65 66 67 68 69 70°F



SEA - CALM  
SWF - BRIGHT  
WIND - 8 KTS  
28 FEB. 1500

176 MC  
OC = 4.6 Db/MYD

23.6 MC  
OC = 6.0 Db/MYD

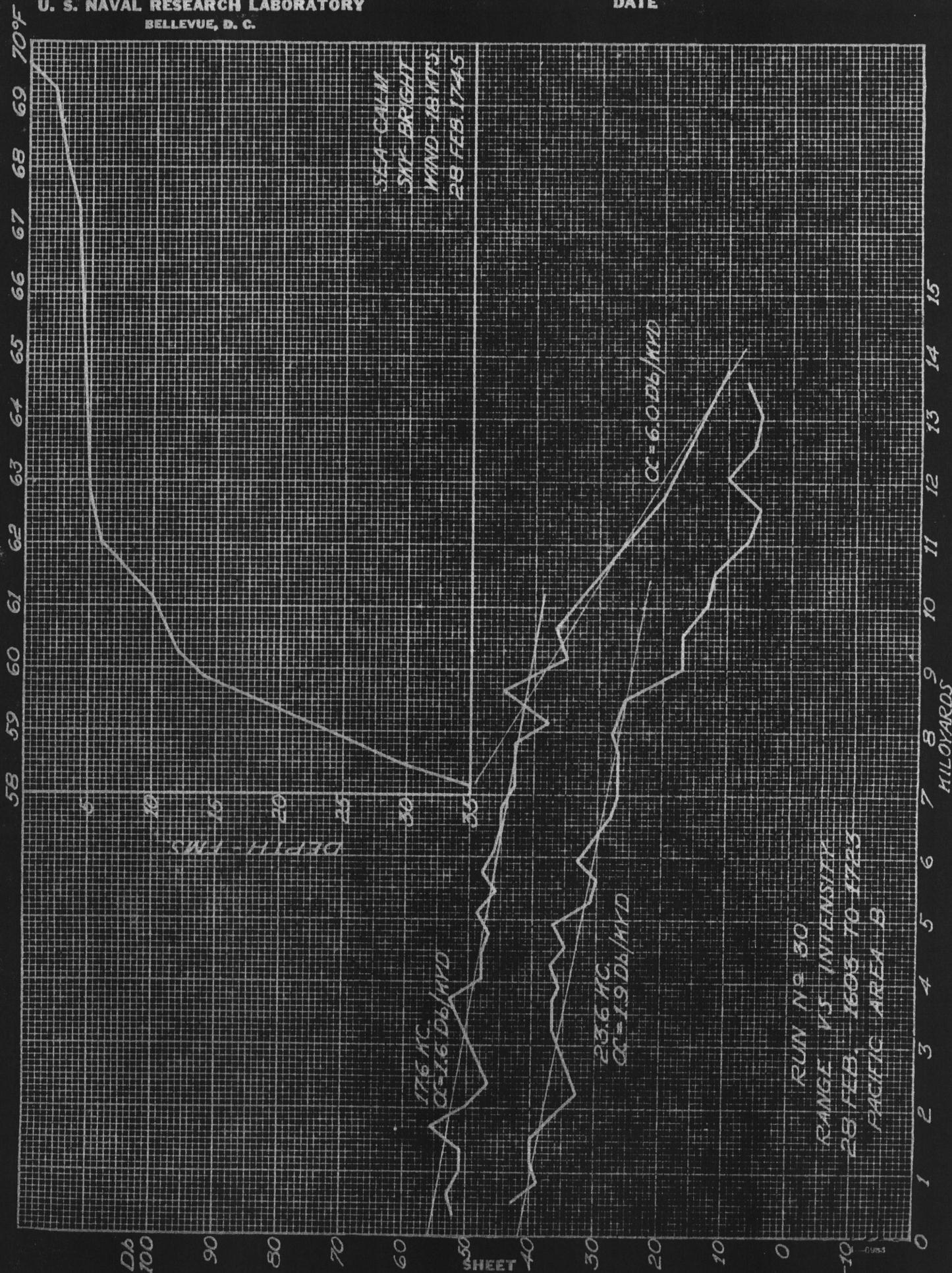
RUN NO 29  
RANGE VS INTENSITY  
28 FEB. 1253 TO 1304  
PACIFIC AREA B

0 1 2 3 4 5 6 7 8 9 10 11 12 13 14 15  
MILYARDS

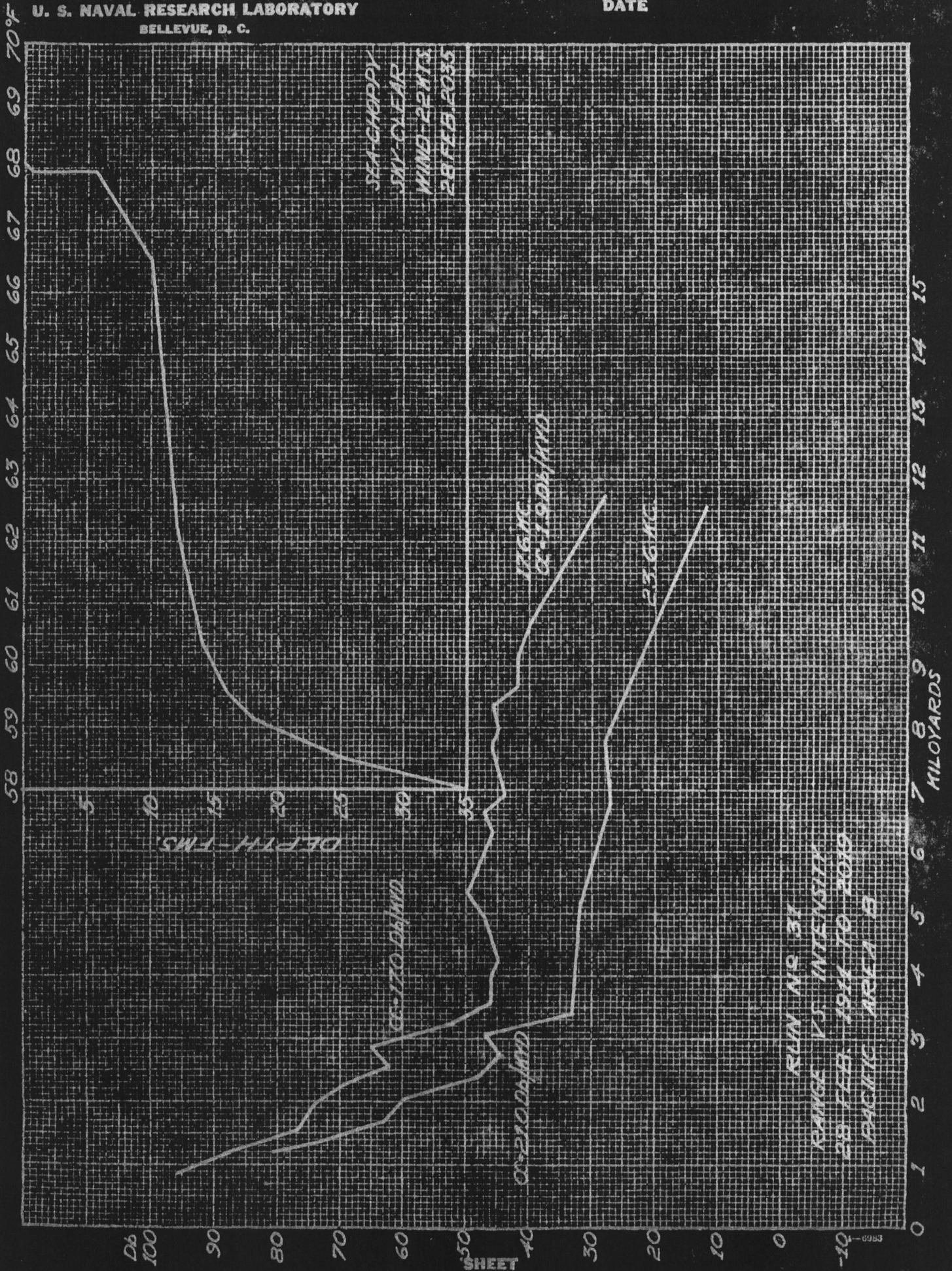
Db  
100  
90  
80  
70

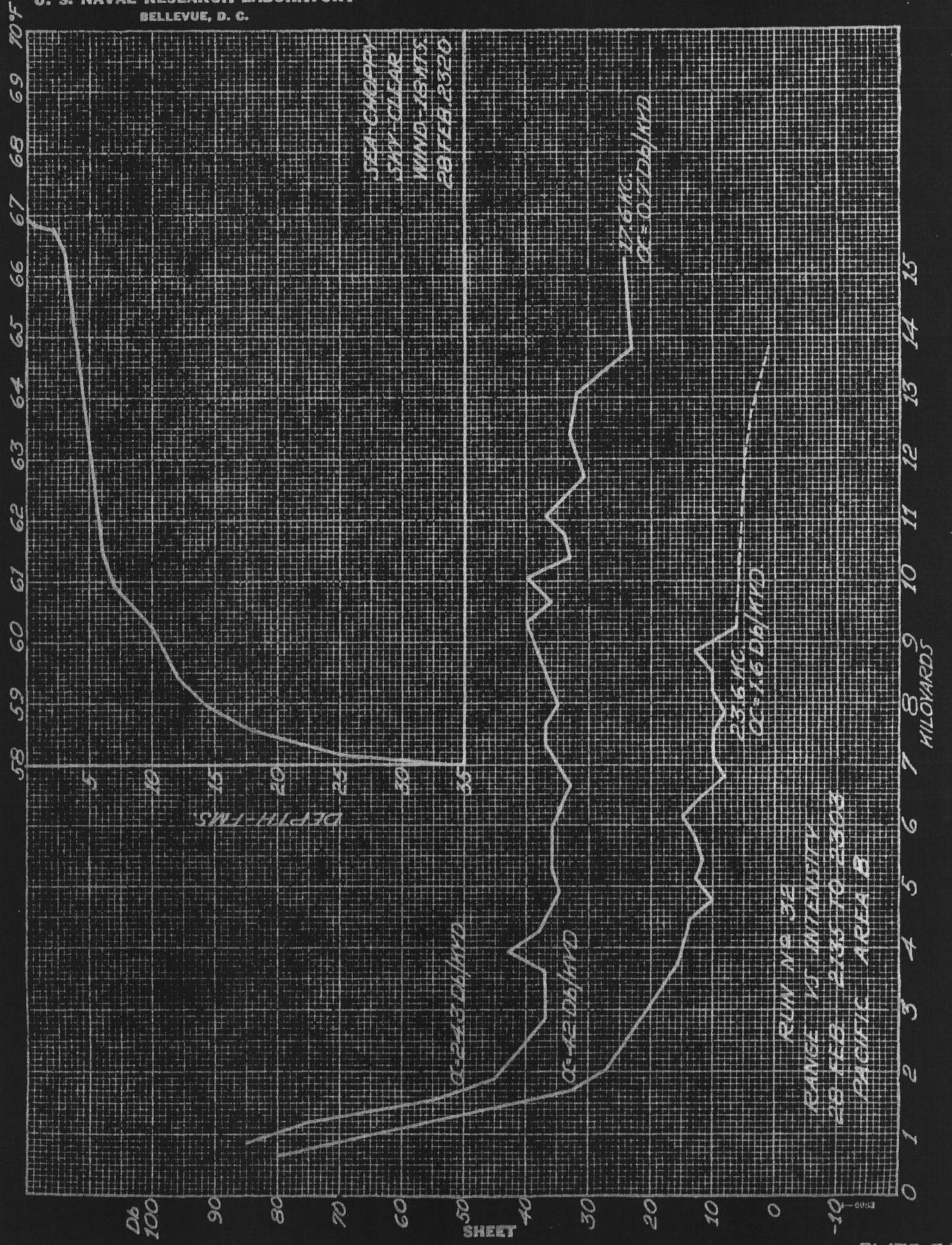
DEPTH - FMS.  
5  
10  
15  
20  
25  
30  
35

SHEET

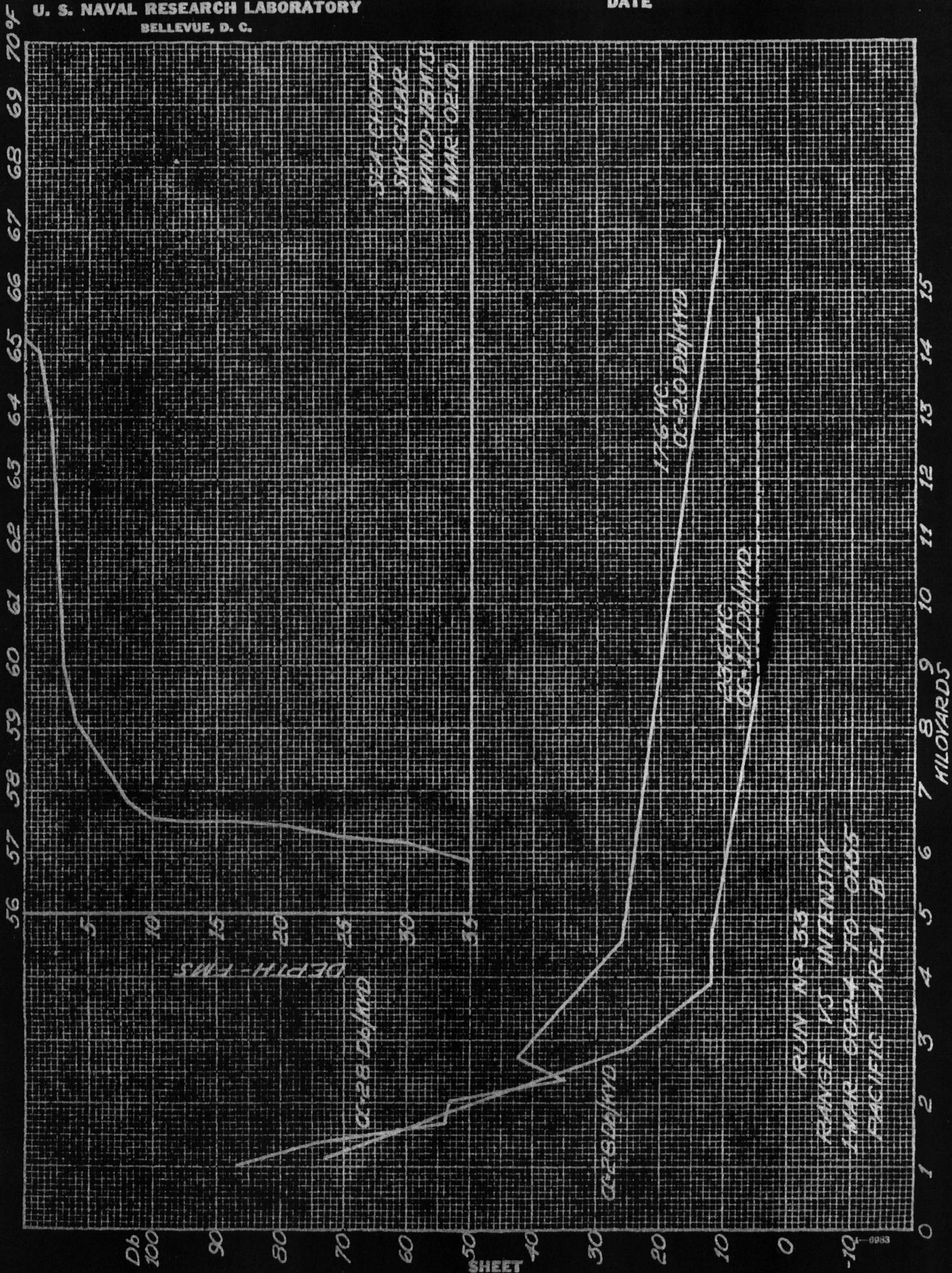


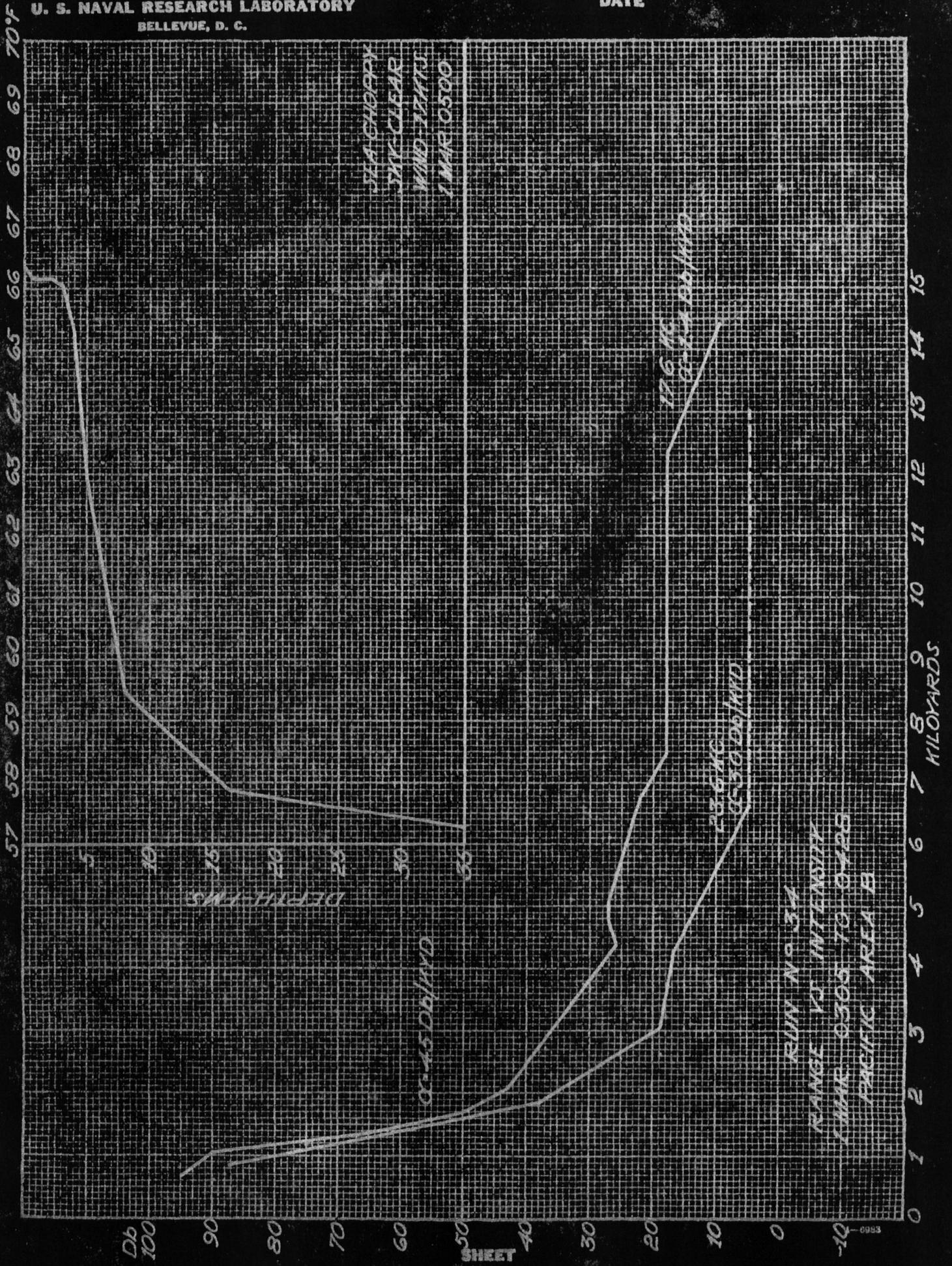
RUN No 30  
RANGE VS INTENSITY  
28 FEB. 1603 TO 1745  
PACIFIC AREA B

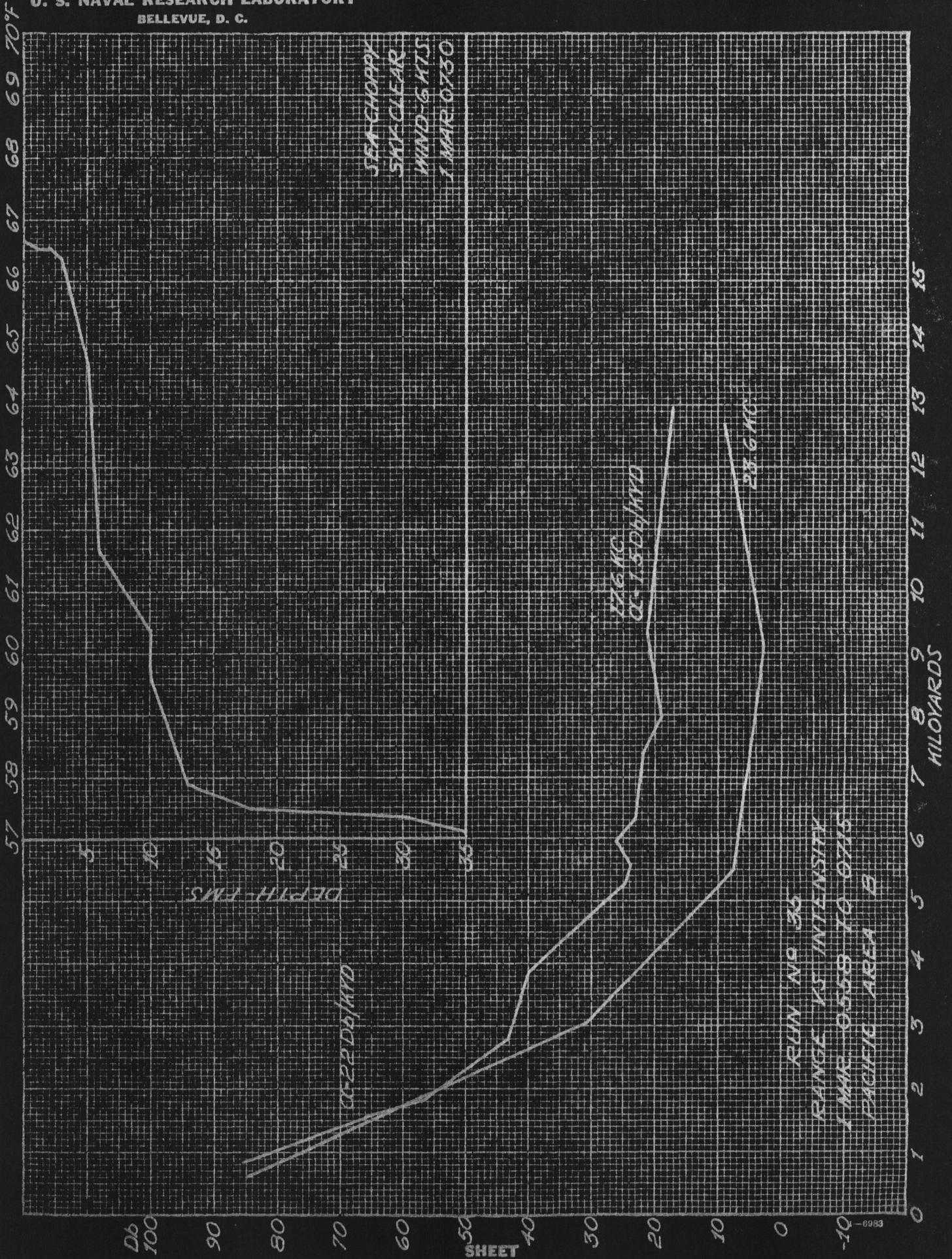




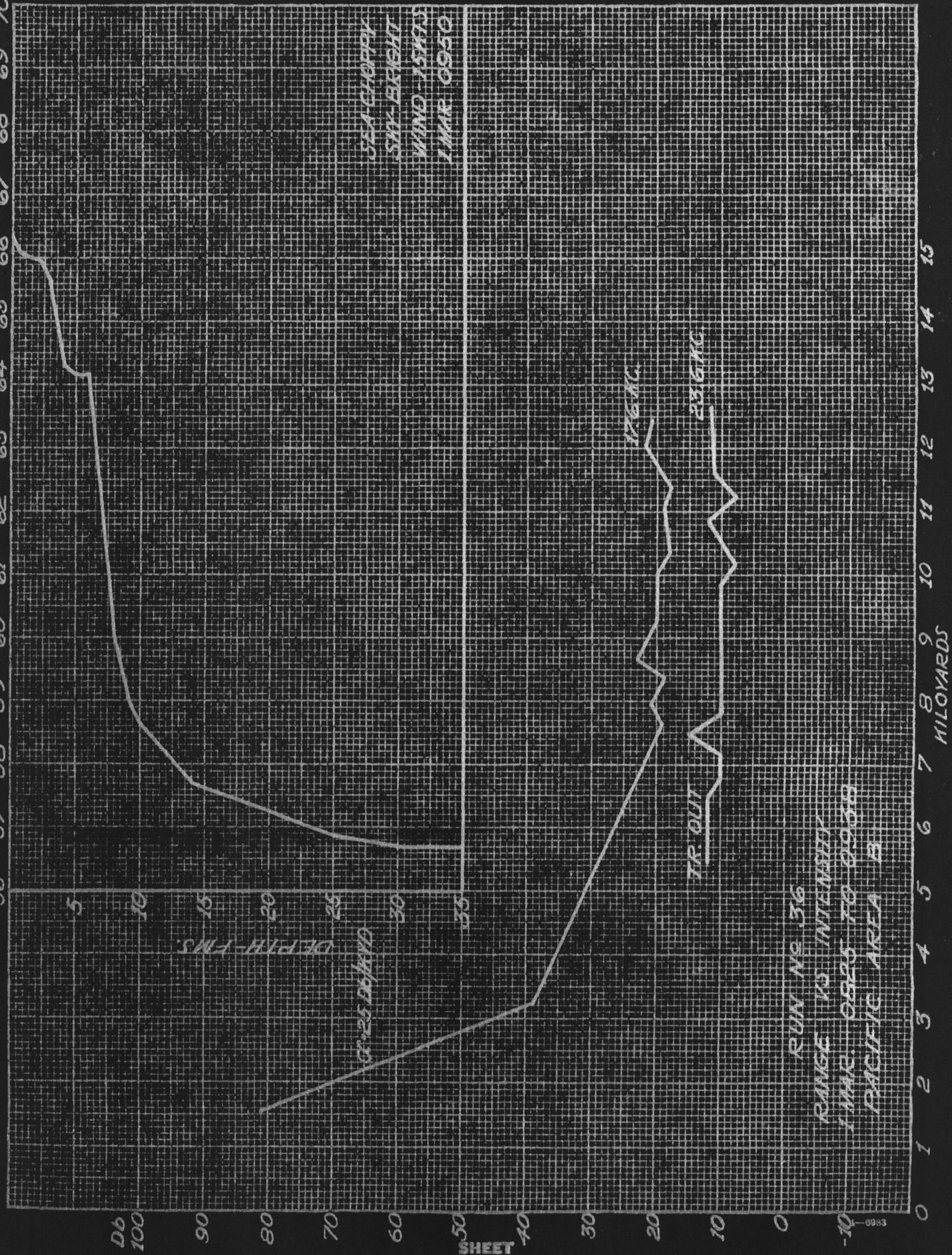
SHEET







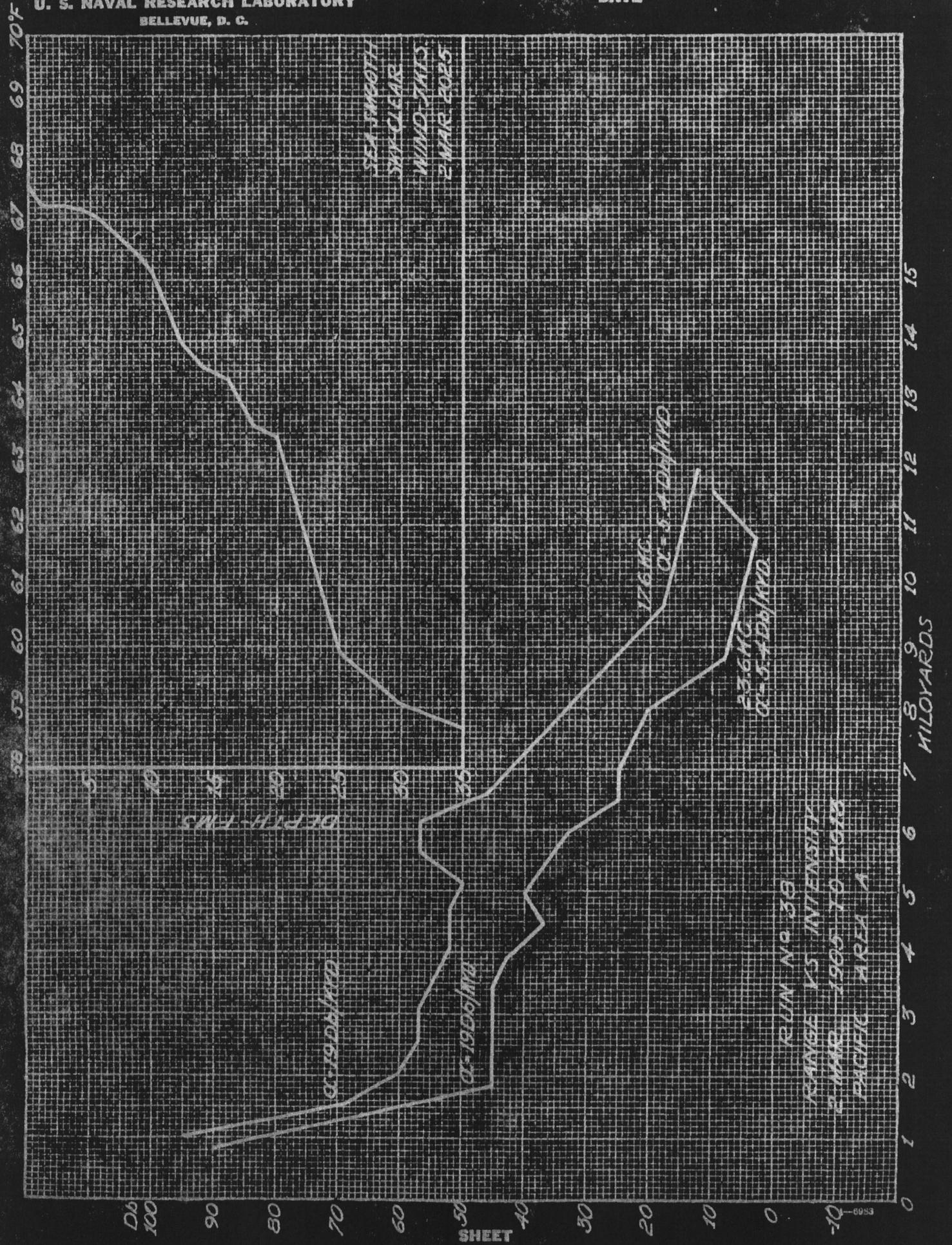
56 57 58 59 60 61 62 63 64 65 66 67 68 69 70°F



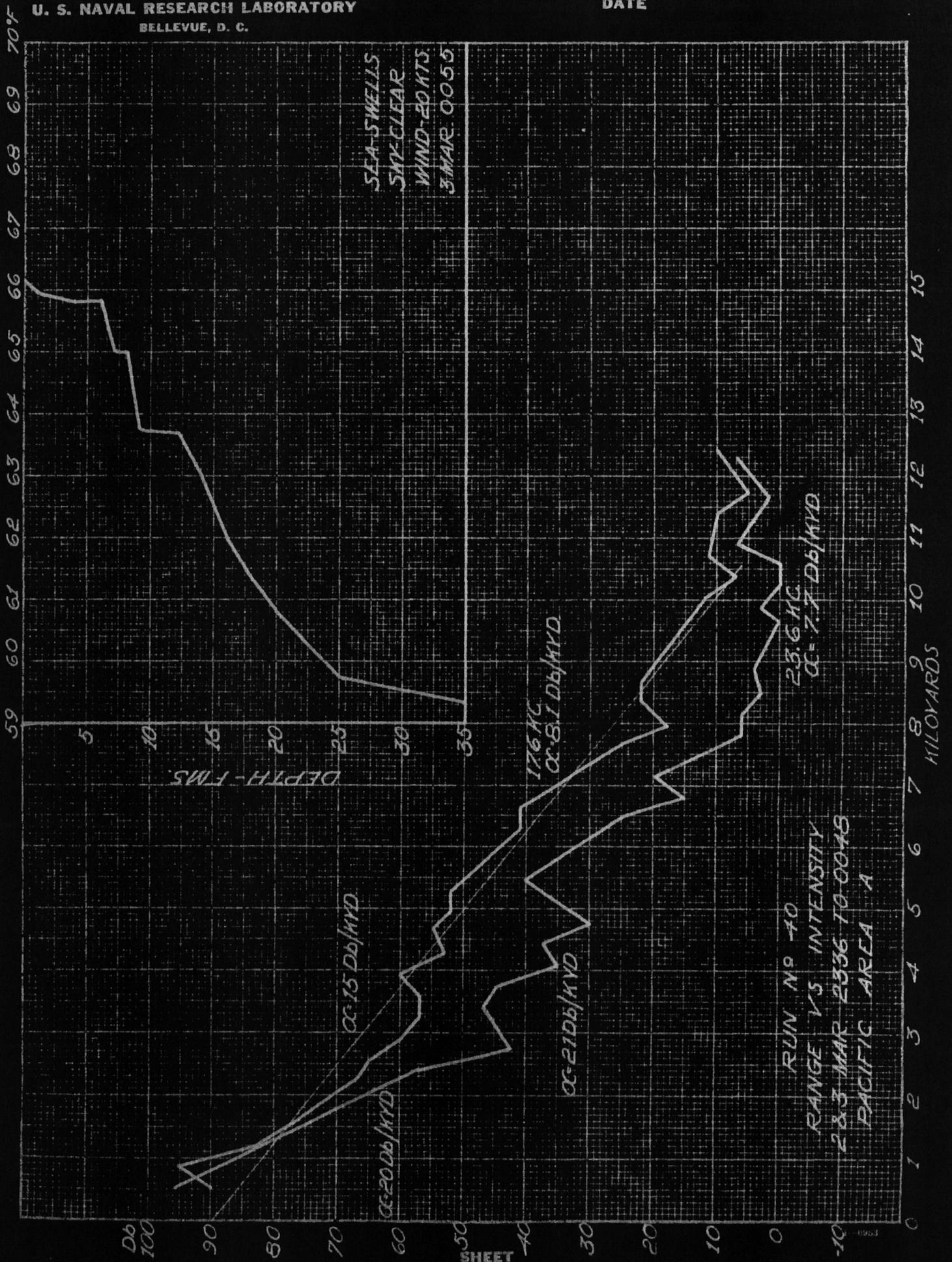
SHEET

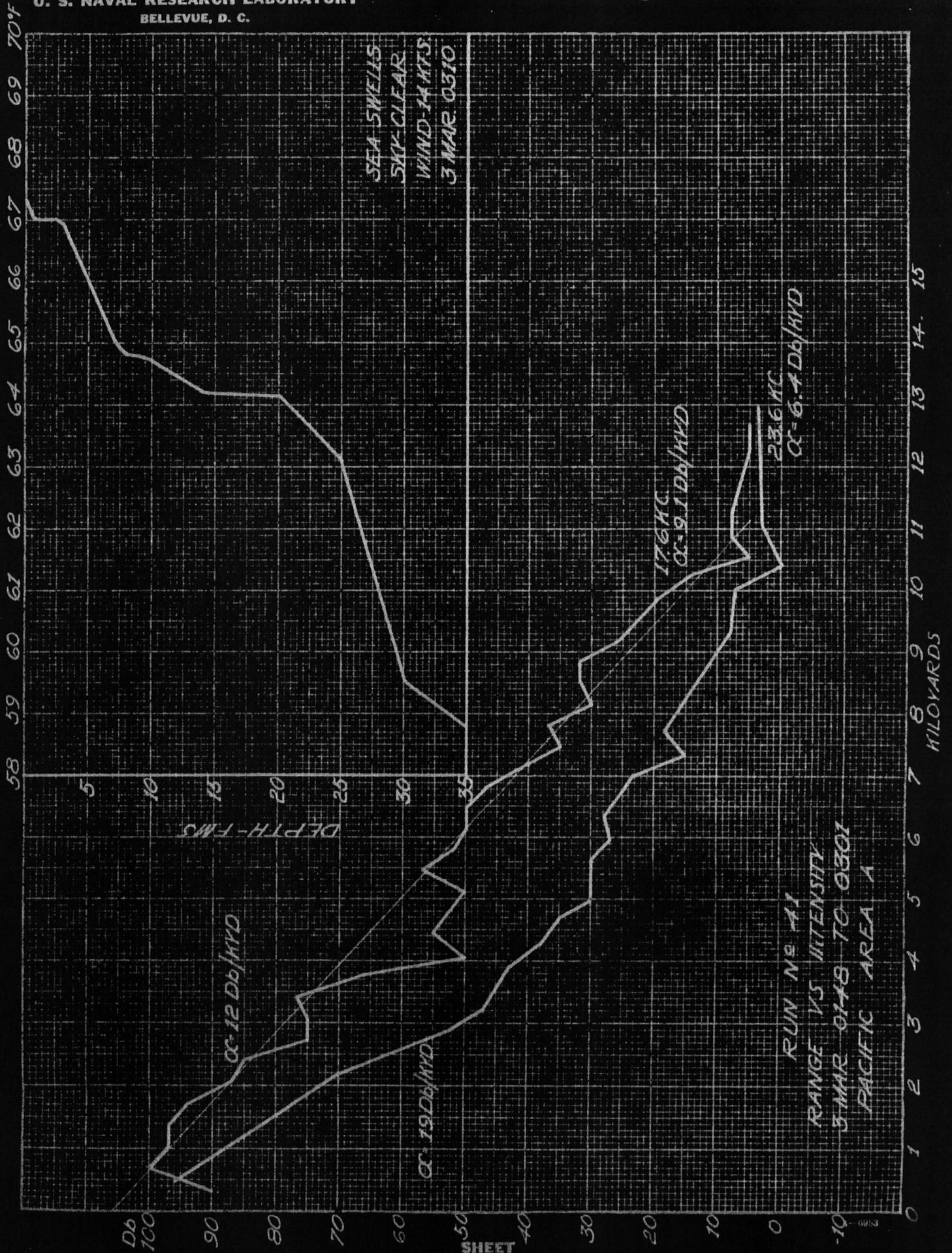
9983

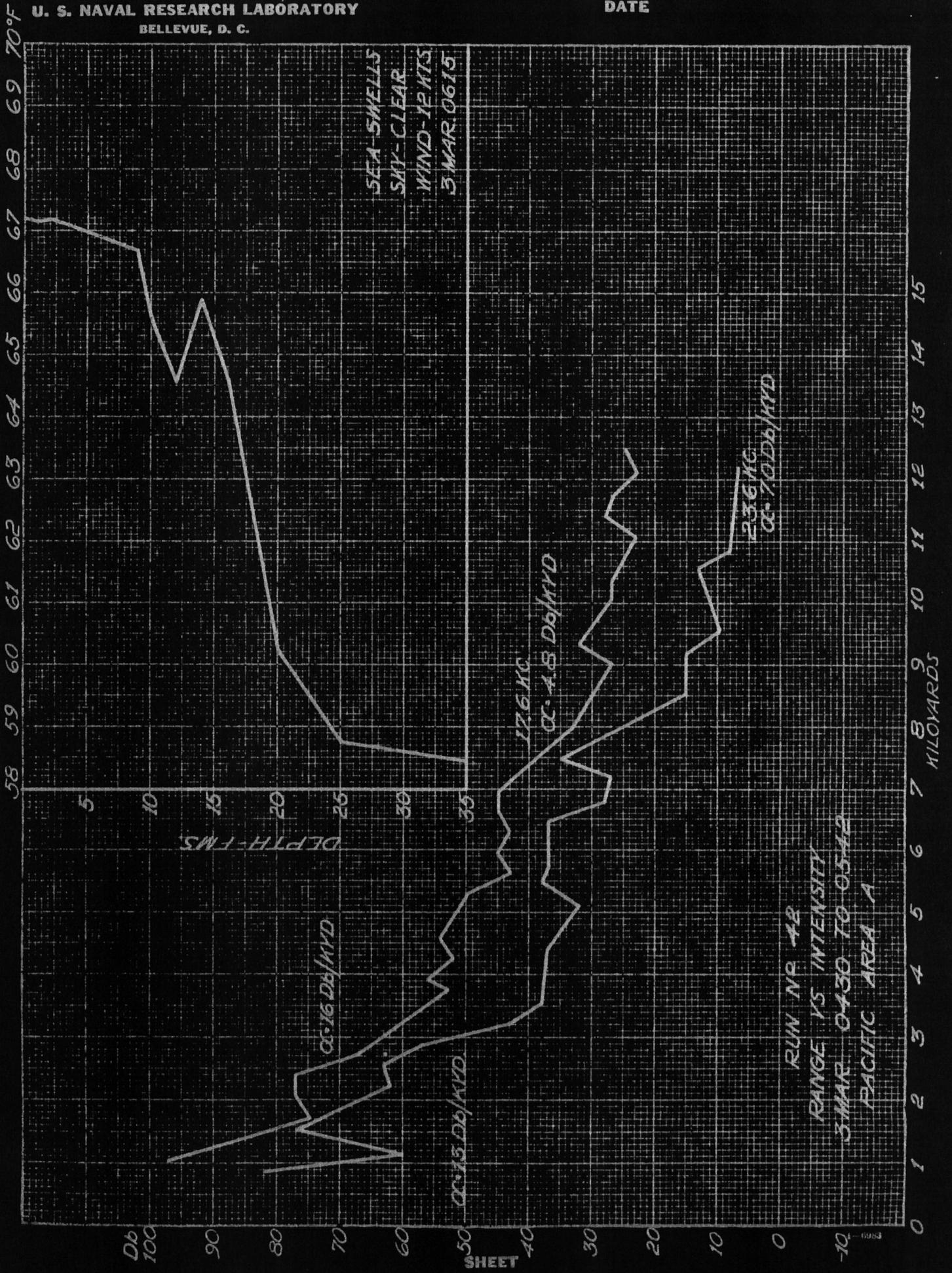




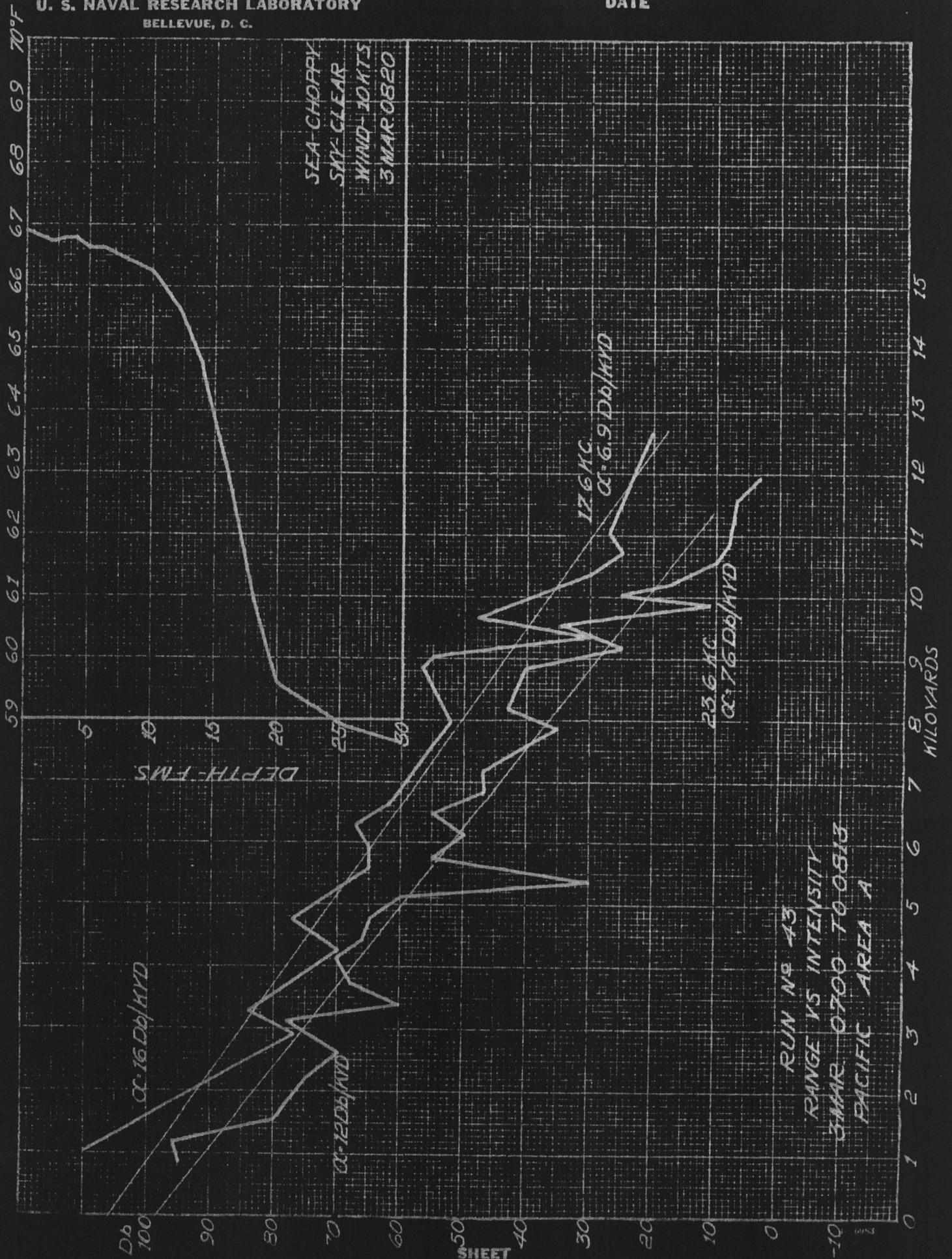


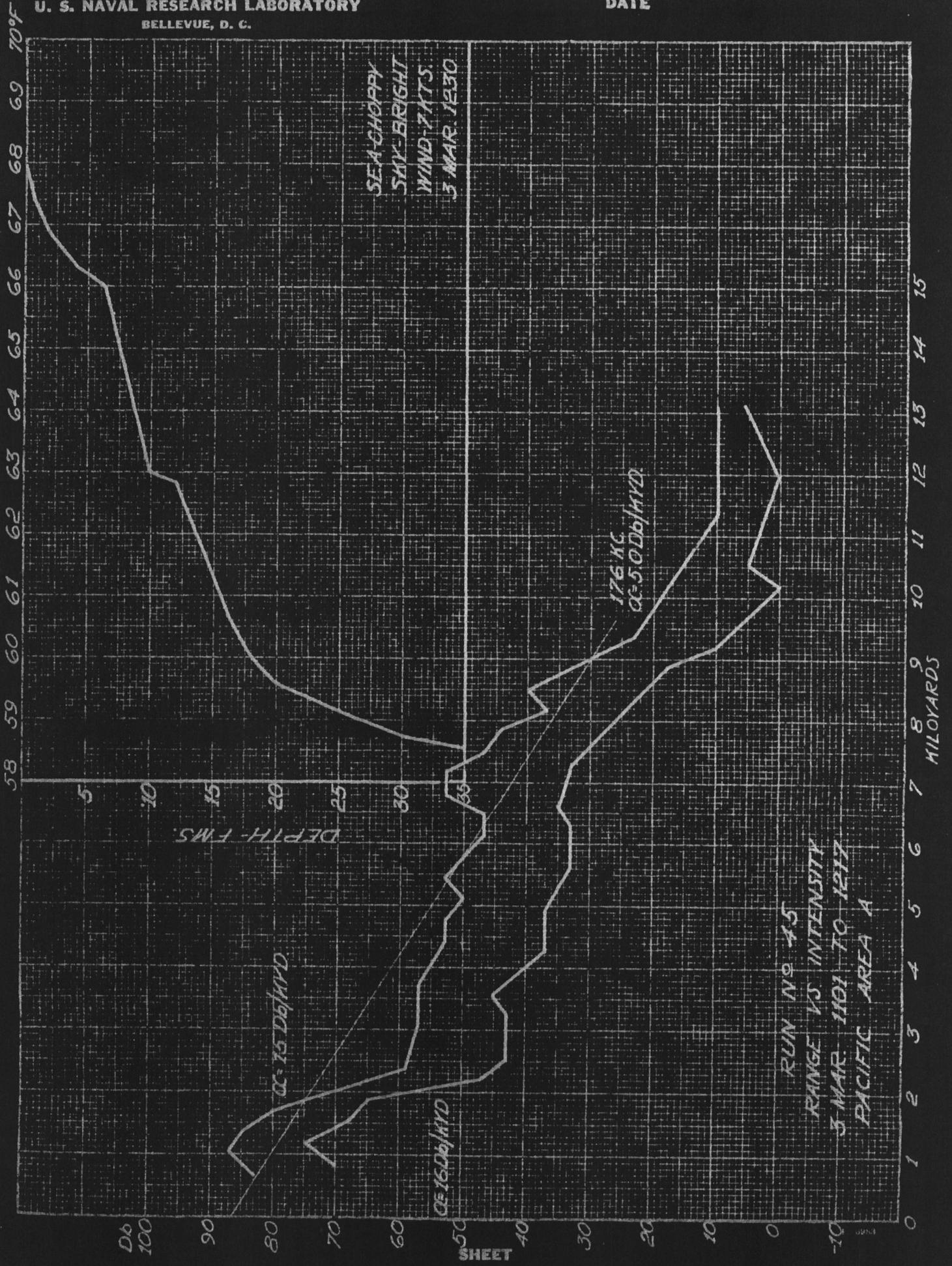


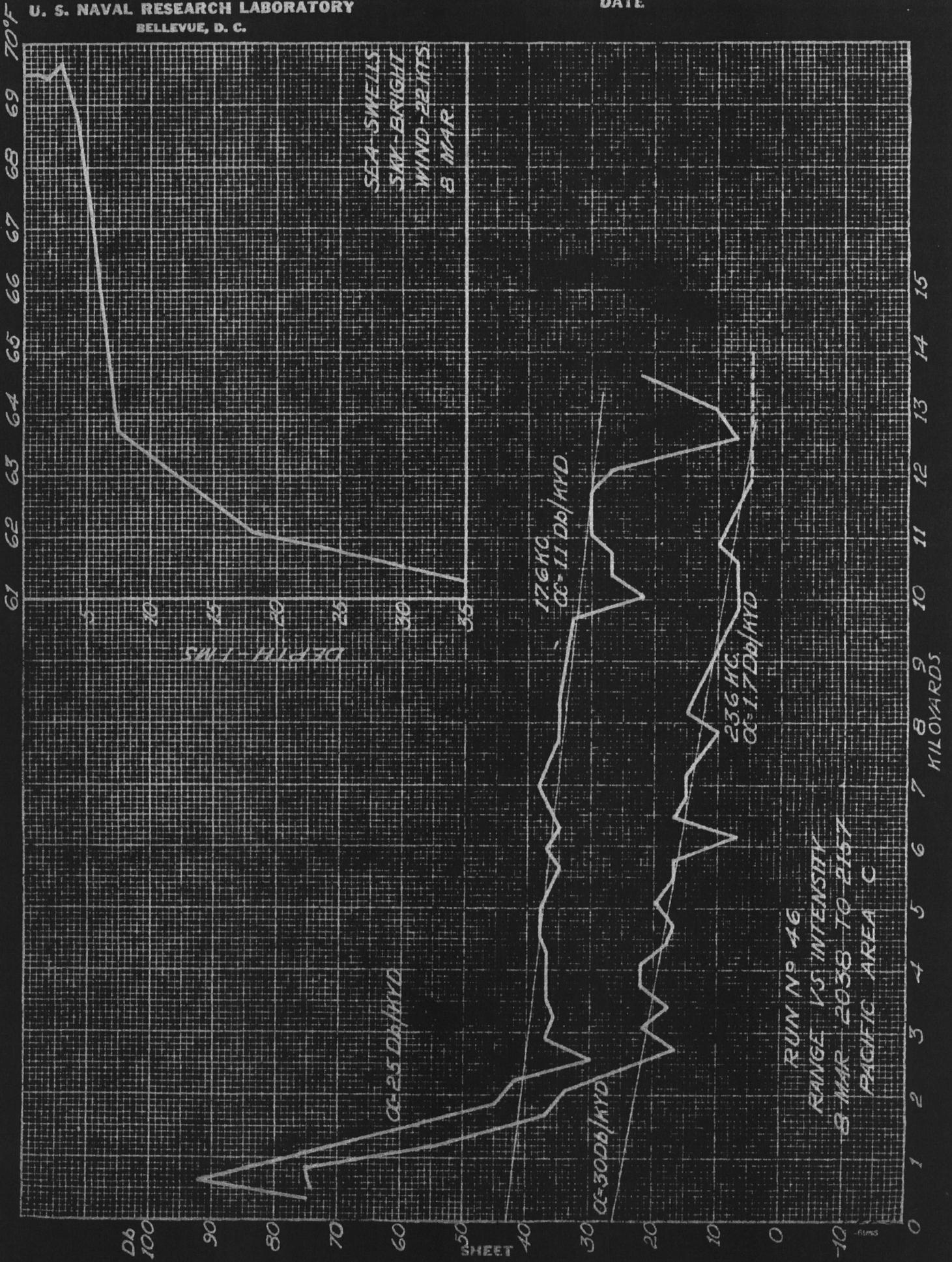




RUN NR 42  
RANGE 15 INTENSITY  
3 MAR 0430 TO 0542  
PACIFIC AREA A







SEA-SWELLS  
SKY-BRIGHT  
WIND-22 KTS  
8 MAR.

DEPTH - FMS

$\alpha=2.5 \text{ Db/Myd}$

17.6 Kt  
 $\alpha=1.1 \text{ Db/Myd}$

23.6 Kt  
 $\alpha=1.7 \text{ Db/Myd}$

$\alpha=30 \text{ Db/Myd}$

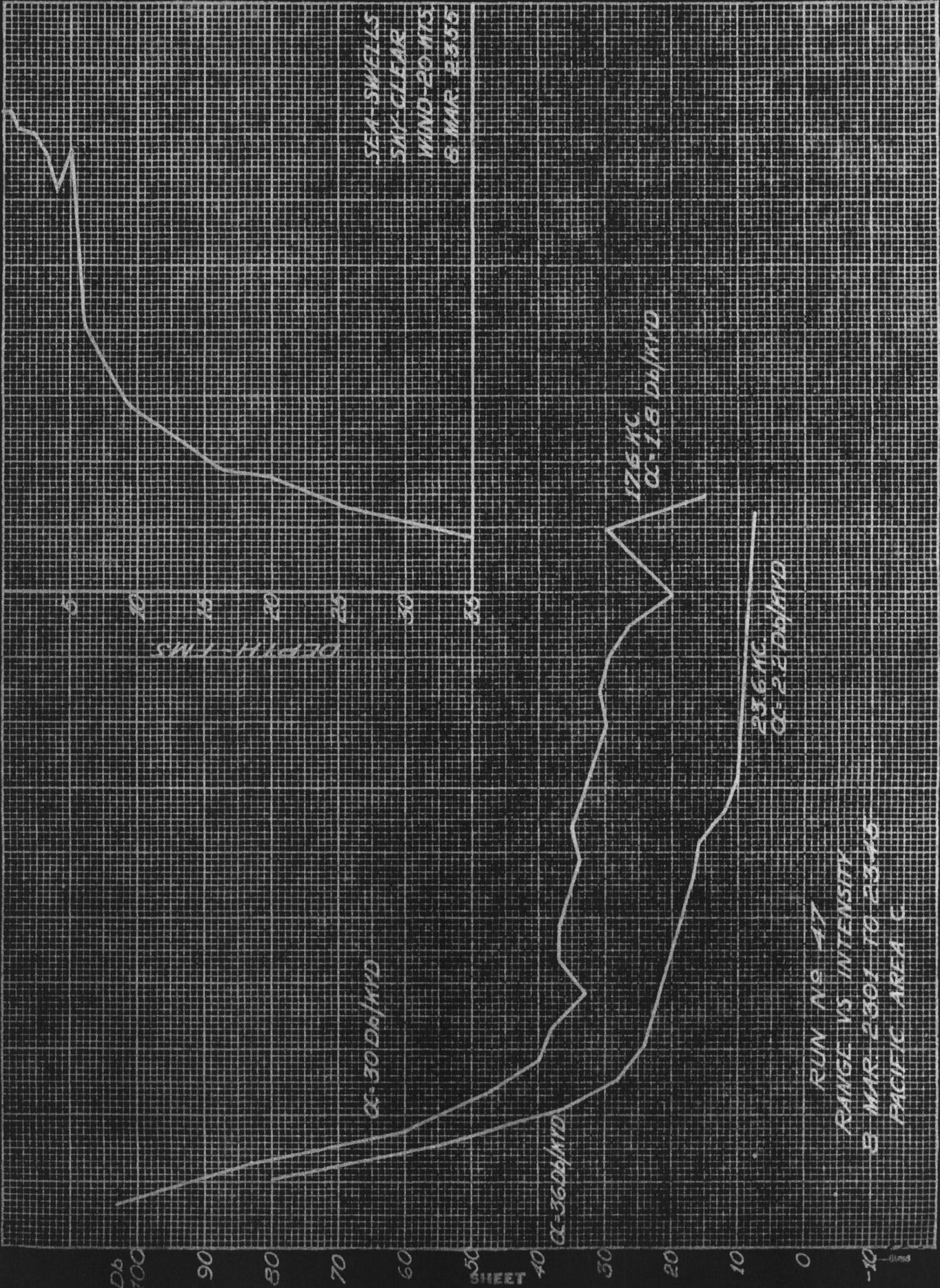
RUN No 46  
RANGE 1/5 INTENSITY  
8 MAR 2038 TO 2157  
PACIFIC AREA C

SHEET

PLATE 46

70°F  
69  
68  
67  
66  
65  
64  
63  
62  
61  
60  
59  
58  
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55  
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49  
48  
47  
46  
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14  
13  
12  
11  
10  
9  
8  
7  
6  
5  
4  
3  
2  
1  
0

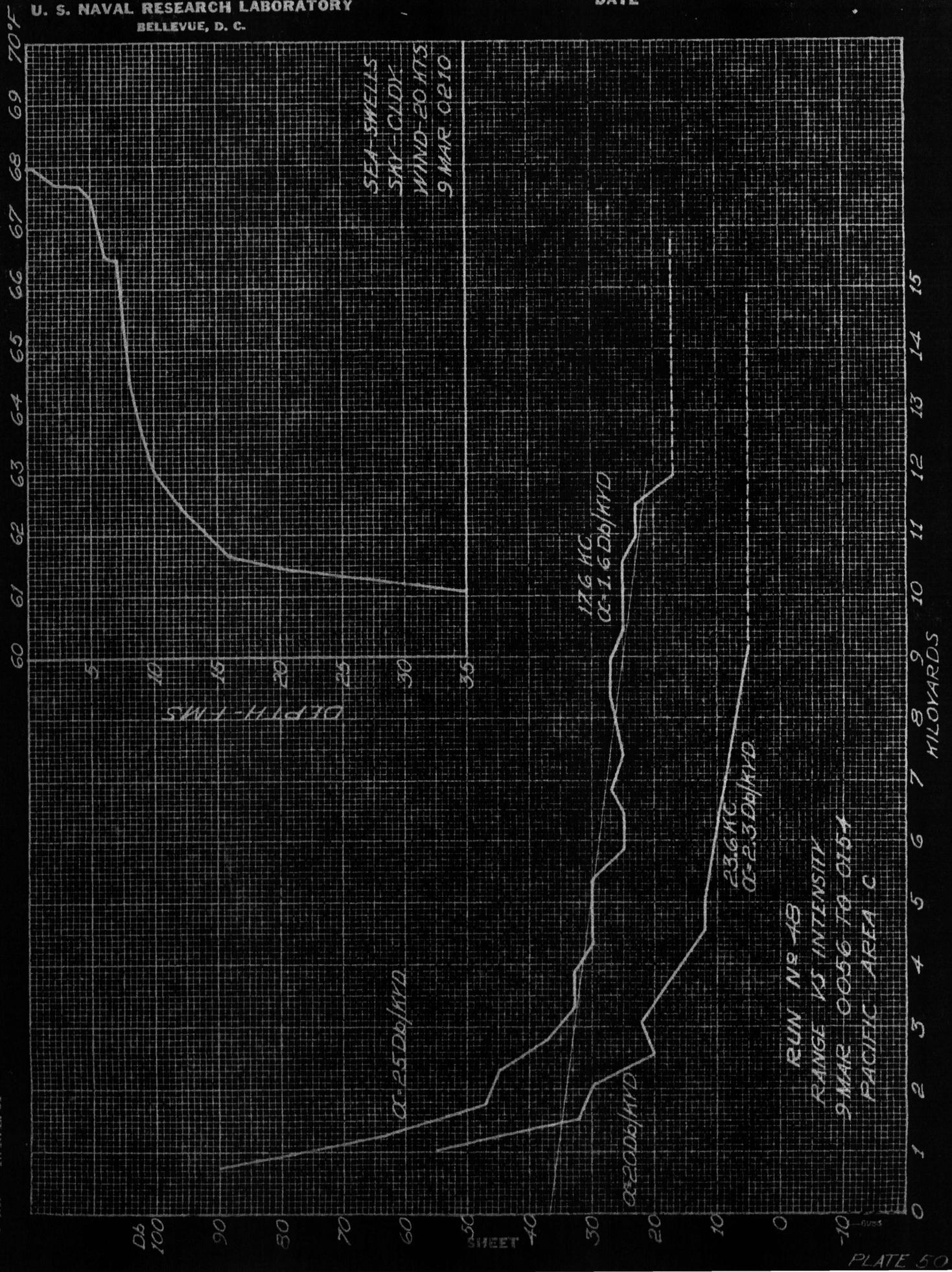
MILOYARDS  
15  
14  
13  
12  
11  
10  
9  
8  
7  
6  
5  
4  
3  
2  
1  
0

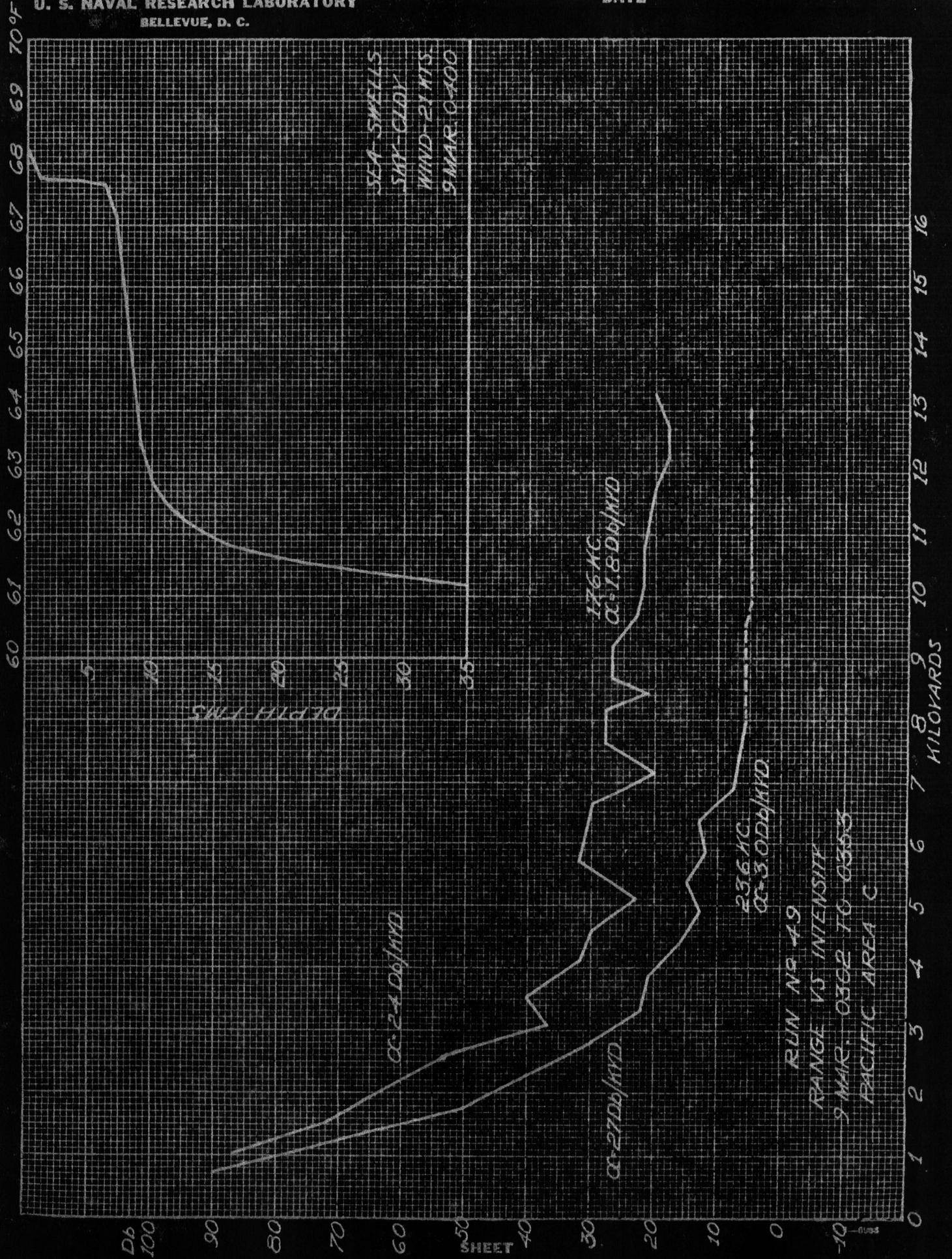


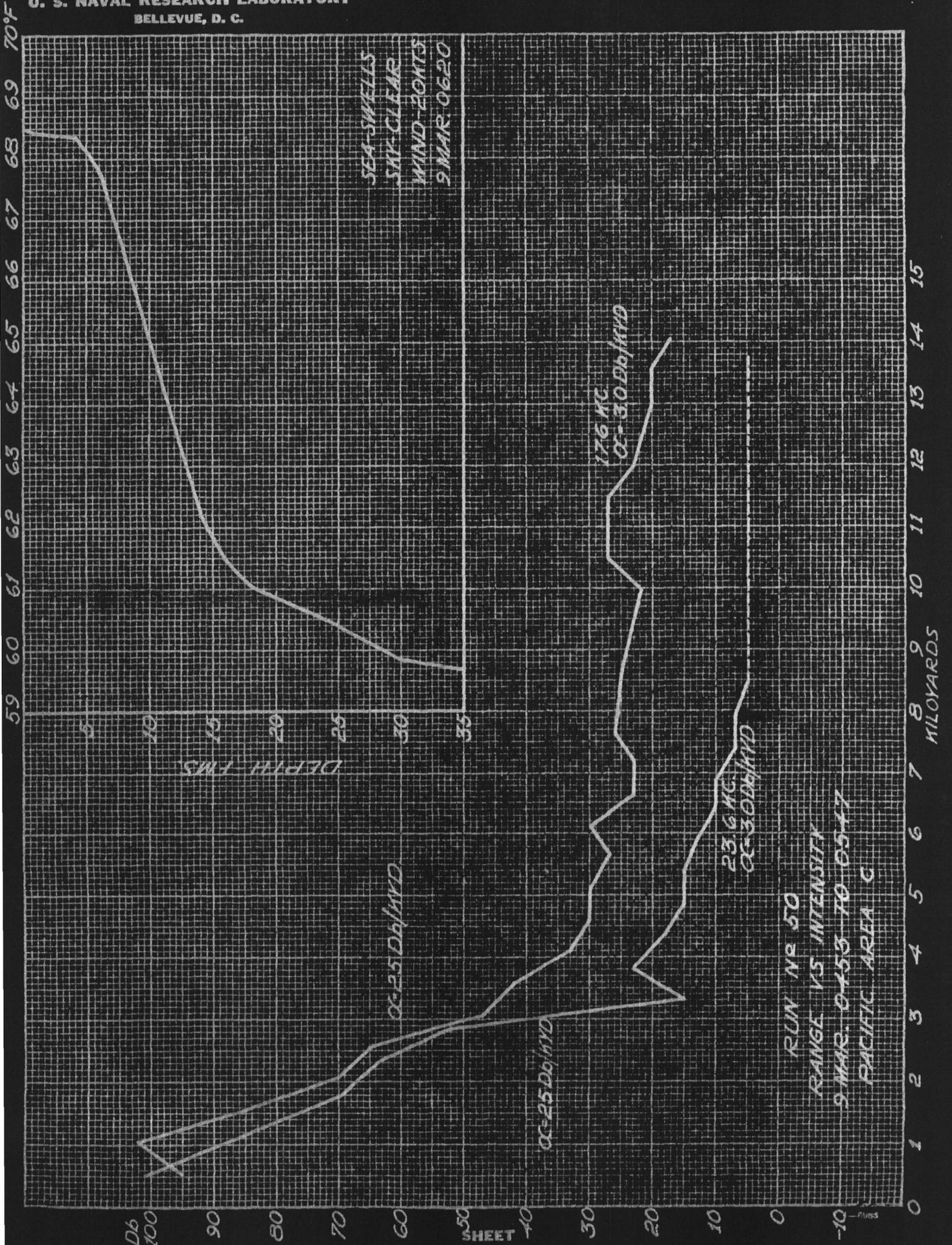
RUN No 47  
RANGE VS INTENSITY  
8 MAR 2301 TO 2345  
PACIFIC AREA C

Db  
100  
90  
80  
70  
60  
50  
40  
30  
20  
10  
0  
-10

SHEET



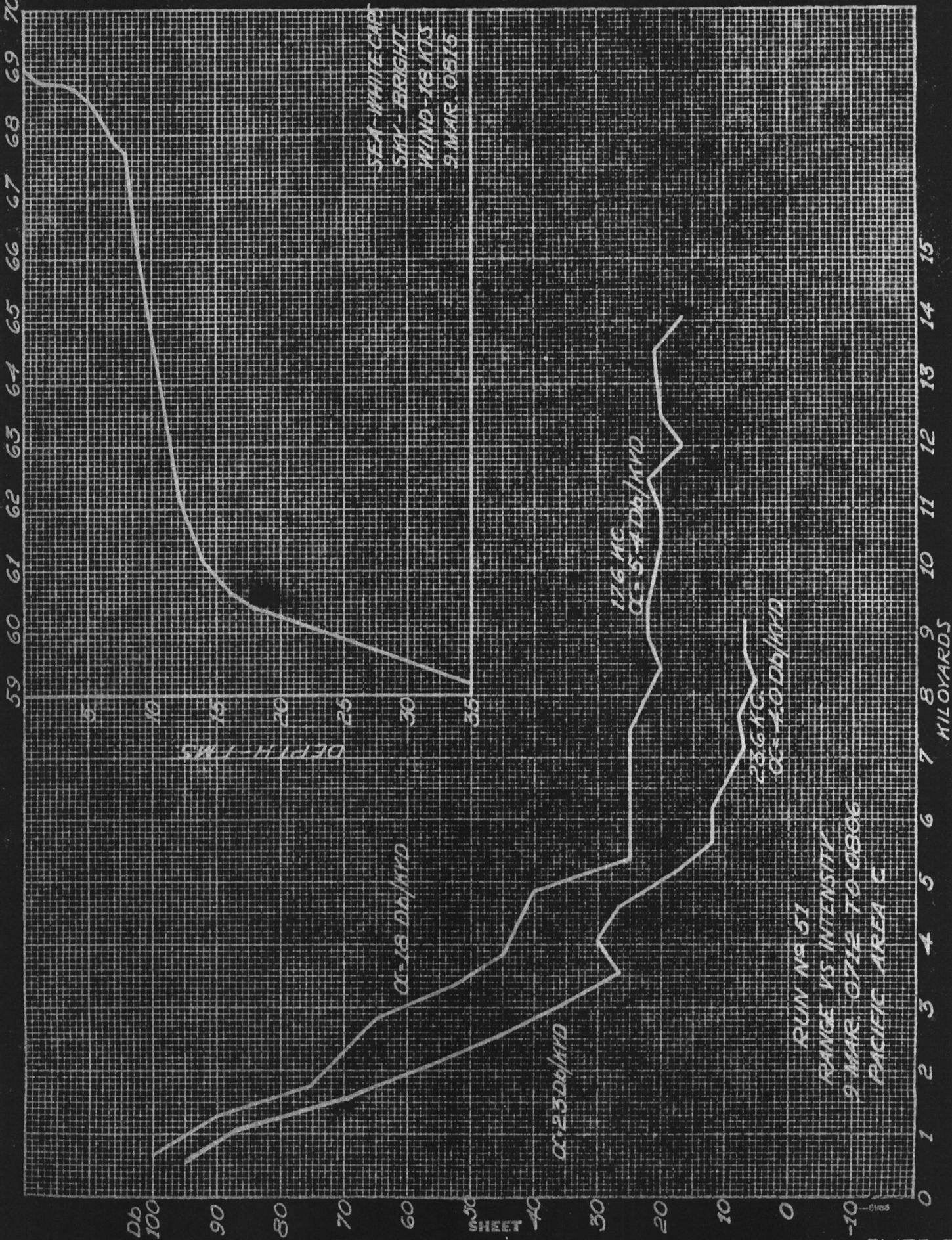


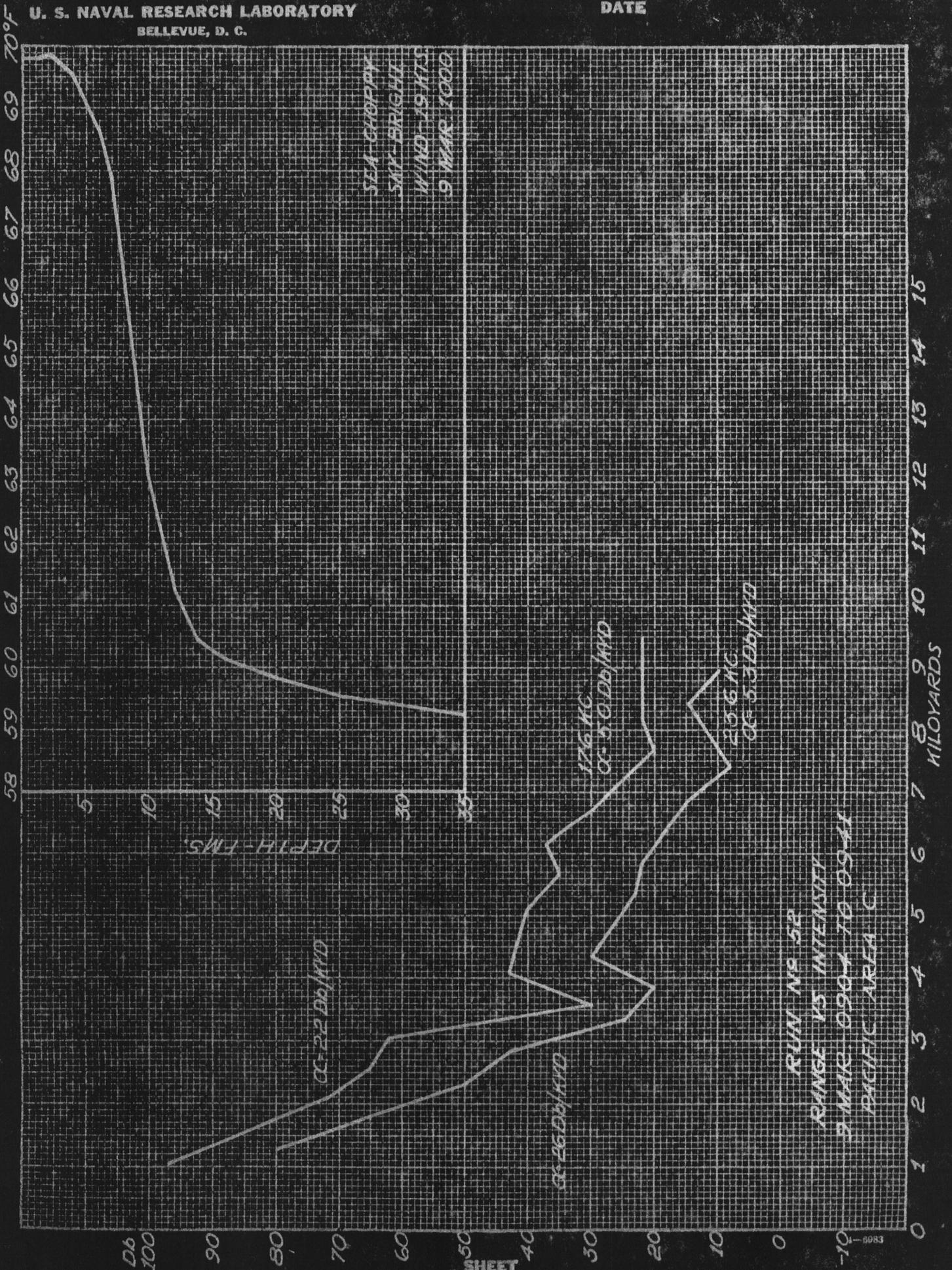


SEA SWELLS  
SKY CLEAR  
WIND 20 KTS  
9 MAR 0620

RUN No 50  
RANGE VS INTENSITY  
9 MAR. 0453 TO 0547  
PACIFIC AREA C

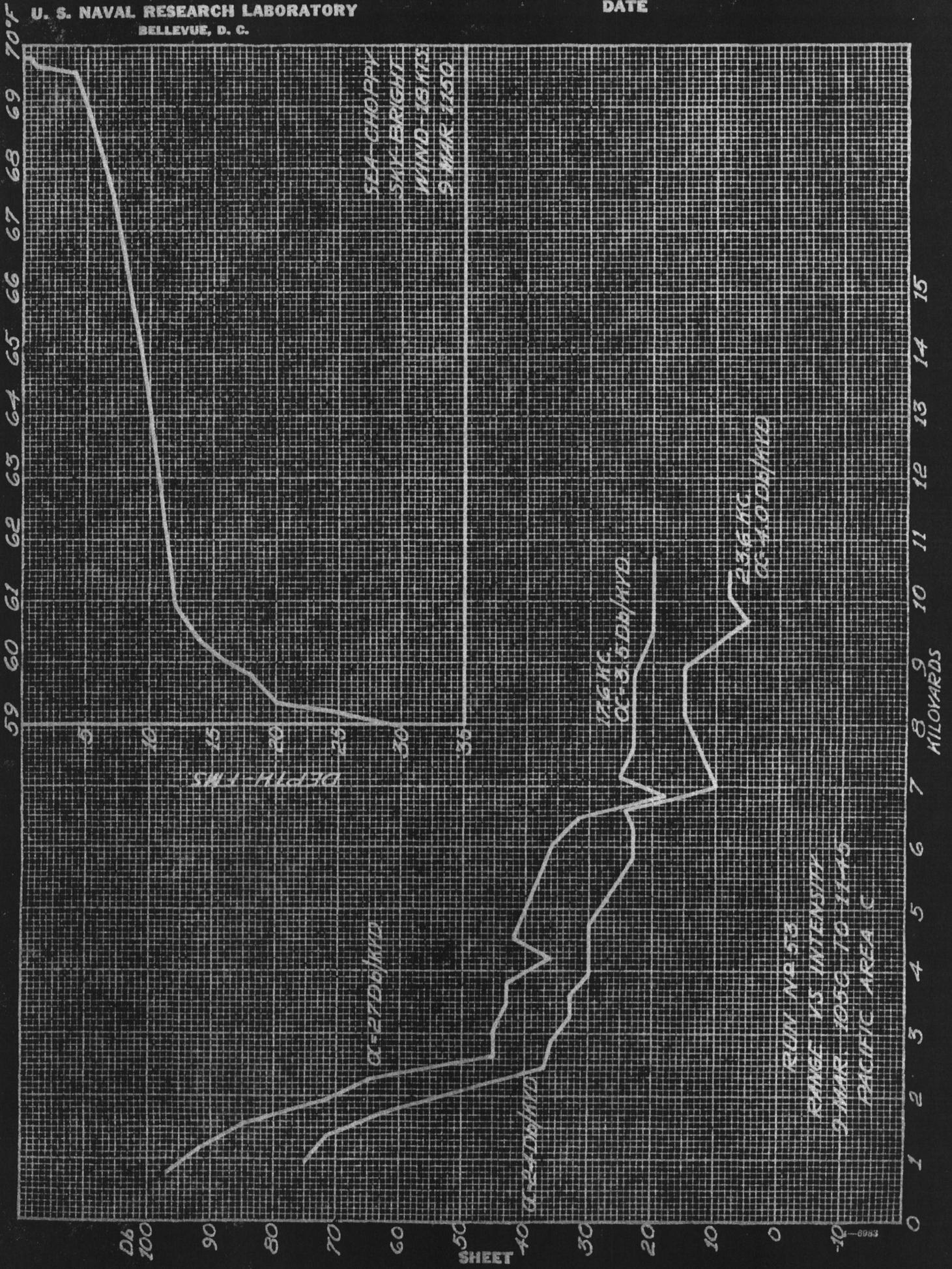
100 90 80 70 60 50 SHEET 40 30 20 10 0 -10





SEA CHOPPY  
SKY BRIGHT  
WIND 19 KTS  
9 MAR 1960

RUN No 512  
RANGE 15 INTENSITY  
9 MAR 0904 TO 0941  
PACIFIC AREA C



SEA-CHOPPY  
SKY-BRIGHT  
WIND-18 KTS  
9 MAR 1950

DEPTN - FMS

MILYARDS

$\alpha=27\text{Db}/\text{mid}$

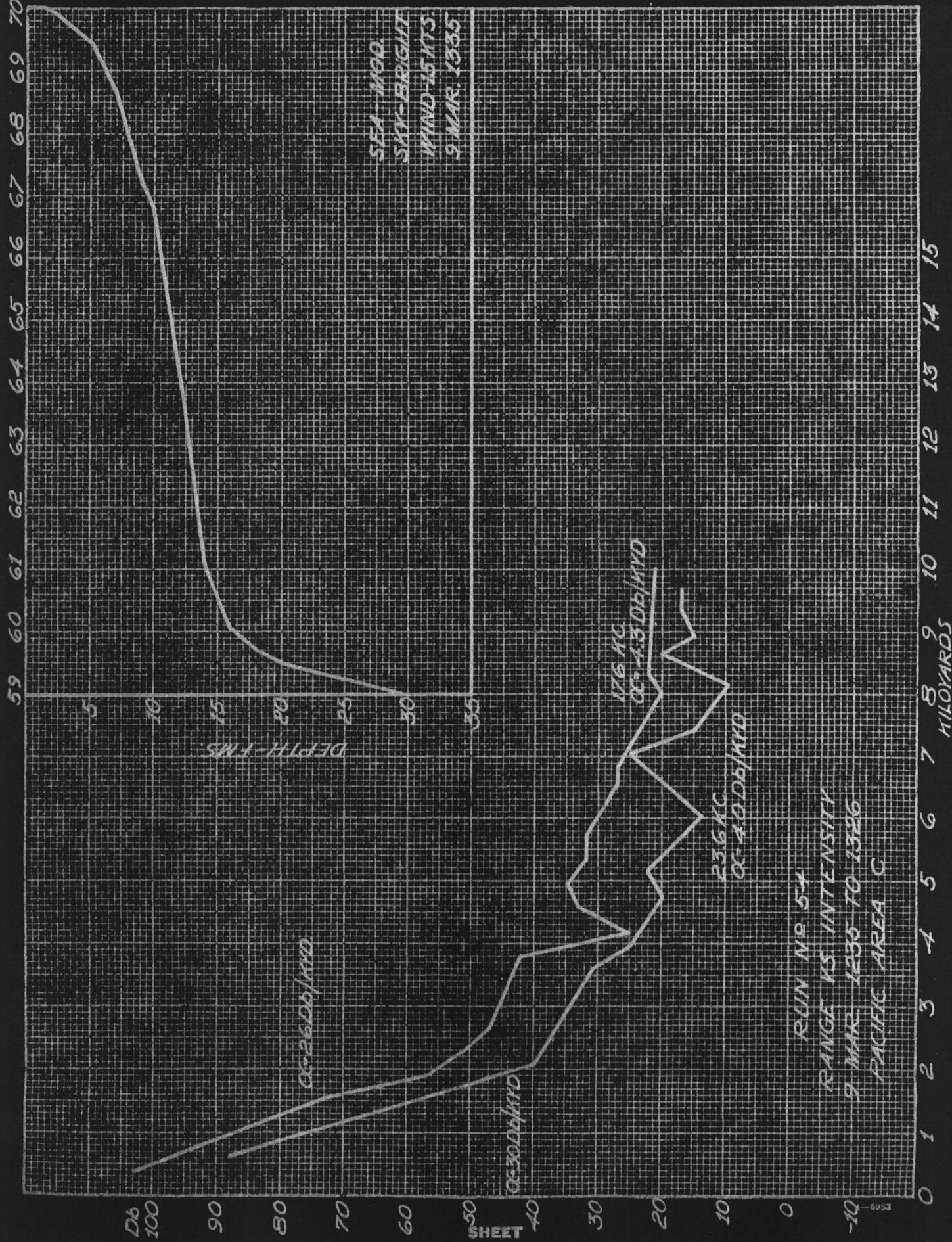
$\alpha=24\text{Db}/\text{mid}$

17.6 Kc  
 $\alpha=3.5\text{Db}/\text{mid}$

23.6 Kc  
 $\alpha=4.0\text{Db}/\text{mid}$

RUN NO 53  
RANGE 15 INTENSITY  
9 MAR 1950 TO 1715  
PACIFIC AREA C

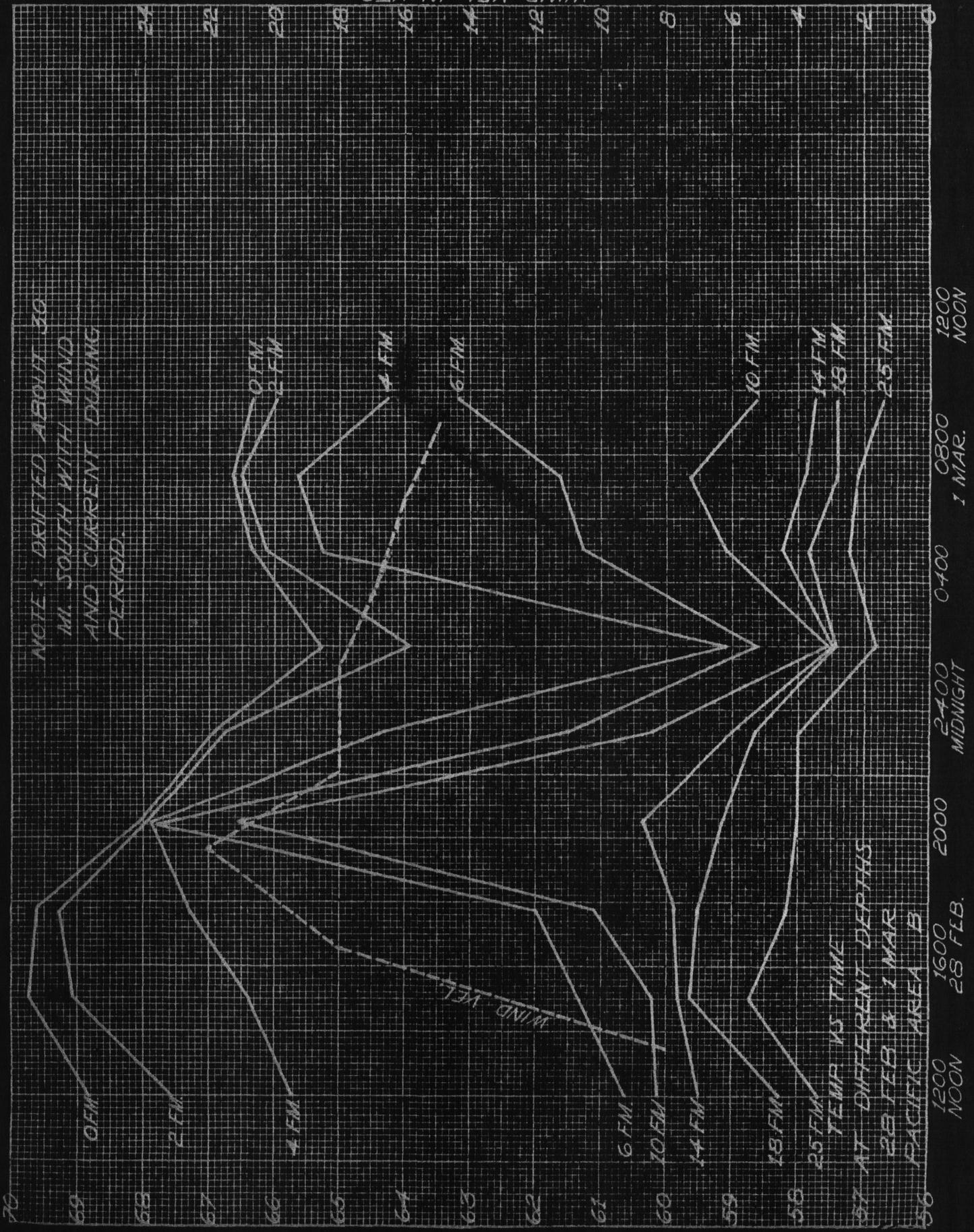
1-6983



SEA - MOD.  
SKY - BRIGHT  
WIND - 15 KTS  
9 MAR 1946

RUN No 54  
RANGE 15 INTENSITY  
9 MAR 1946 TO 1346  
PACIFIC AREA C

WIND VEL. IN KTS.



TEMP - °F  
 SHEET

8599-4

58 59 60 61 62 63 64 65 66 67 68 69 70°F

PANAMA BAY  
 LAT - 8°23'  
 LONG - 79°26'  
 SEA - CALM  
 SKY - BRIGHT  
 WIND - CALM  
 DEPTH - 27 FMS

24 FEB 1950

59 60 61 62 63 64 65 66 67 68 69 70°F

PANAMA BAY  
 LAT - 8°23'  
 LONG - 79°27'  
 SEA - CALM  
 SKY - BRIGHT  
 WIND - CALM  
 DEPTH - 27 FMS

20 FEB 1950

5-0022 N.R.L. 54

SHEET

PLATE 5E

